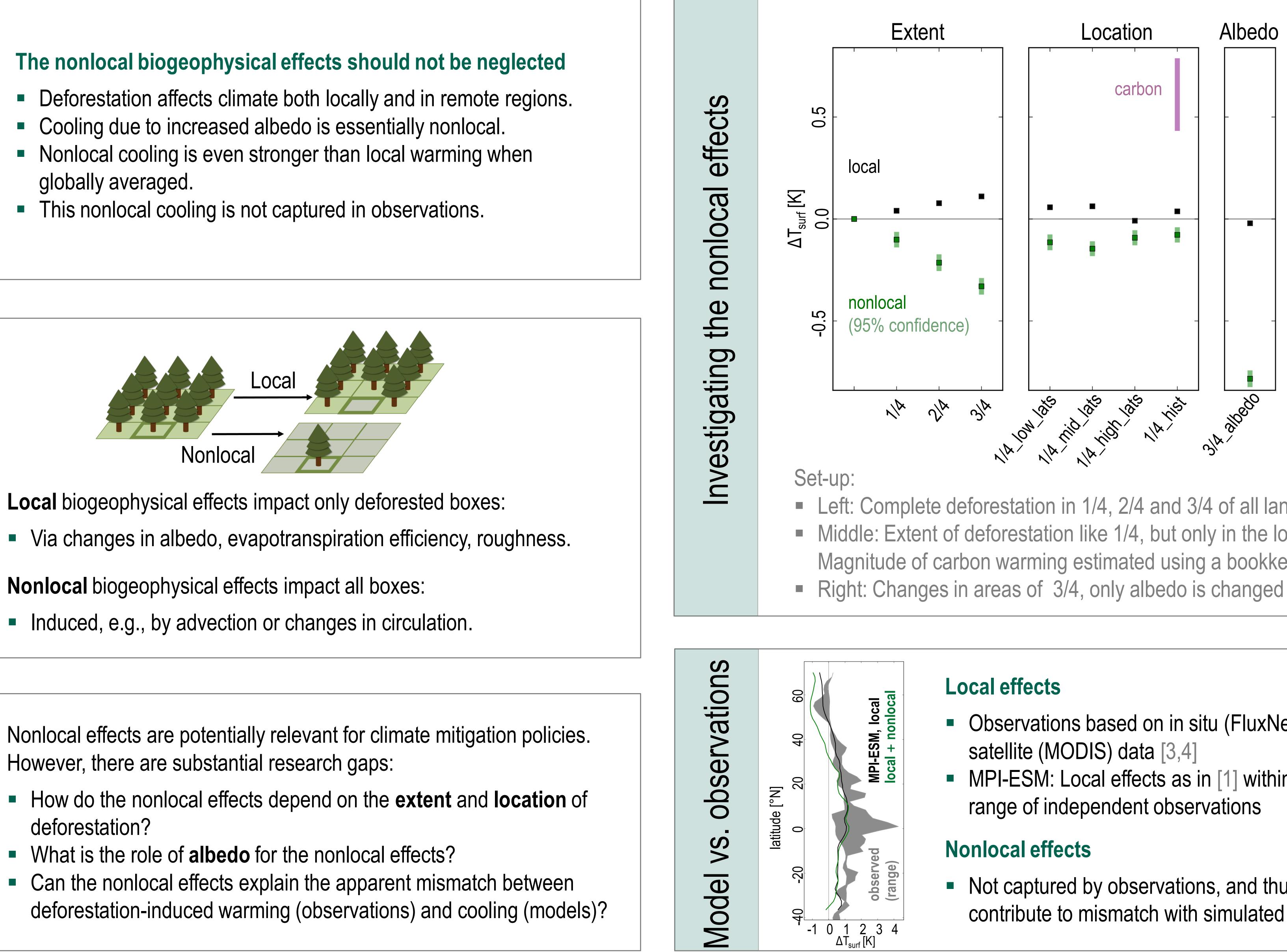




- globally averaged.



Motivation

However, there are substantial research gaps:



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References [1] Winckler, J., Reick, C. H., & Pongratz, J. 2016. Robust identification of local biogeopysical effects of land cover change in a global climate model. J. Clim., 30(3), 1159–1176, doi: 10.1175/JCLI-D-16-0067.1 [2] Bright, R. M., Davin, E. L., O'Halloran, T. L., Pongratz, J., Zhao, K., & Cescatti, A. 2017. Local surface temperature response to land cover and management change driven by non-radiative processes. Nat. Clim. Chang., doi:10.1038/NCLIMATE3250 [3] Li, Y., Zhao, M., Motesharrei, S., Mu, Q., Kalnay, E., & Li, S. 2015. Local cooling and warming effects of forests based on satellite observations. Nat. Commun., 6, 6603, doi:10.1038/ncomms7603 [4] Alkama, R, & Cescatti, A. 2016. Biophysical climate impacts of recent changes in global forest cover. Science, 351(6276), 600–604, doi:10.1126/science.aac8083 [5] Hansis, E., Davis, S. J., & Pongratz, J. 2015. Relevance of methodological choices for accounting of land use change carbon fluxes. Global Biogeochem. Cycles, 29, 1230–1246, doi:10.1002/2014GB004997 [6] Gillett, N., Arora, V., Matthews, D., and Allen, M. 2013. Constraining the ratio of global warming to cumulative CO₂ emissions using CMIP5 simulations. J. Climate, 26, 6844–6858, doi:10.1175/JCLI-D-12-00476.1

The neglected nonlocal climate effects of deforestation

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Nonlocal effects depend on extent and location of deforestation

- Nonlocal effects scale linearly with number of deforestation boxes.
- Nonlocal cooling occurs also for low-latitude deforestation and a more realistic distribution of deforestation ('hist').
- The overall effect of deforestation is likely still warming because of carbon release.

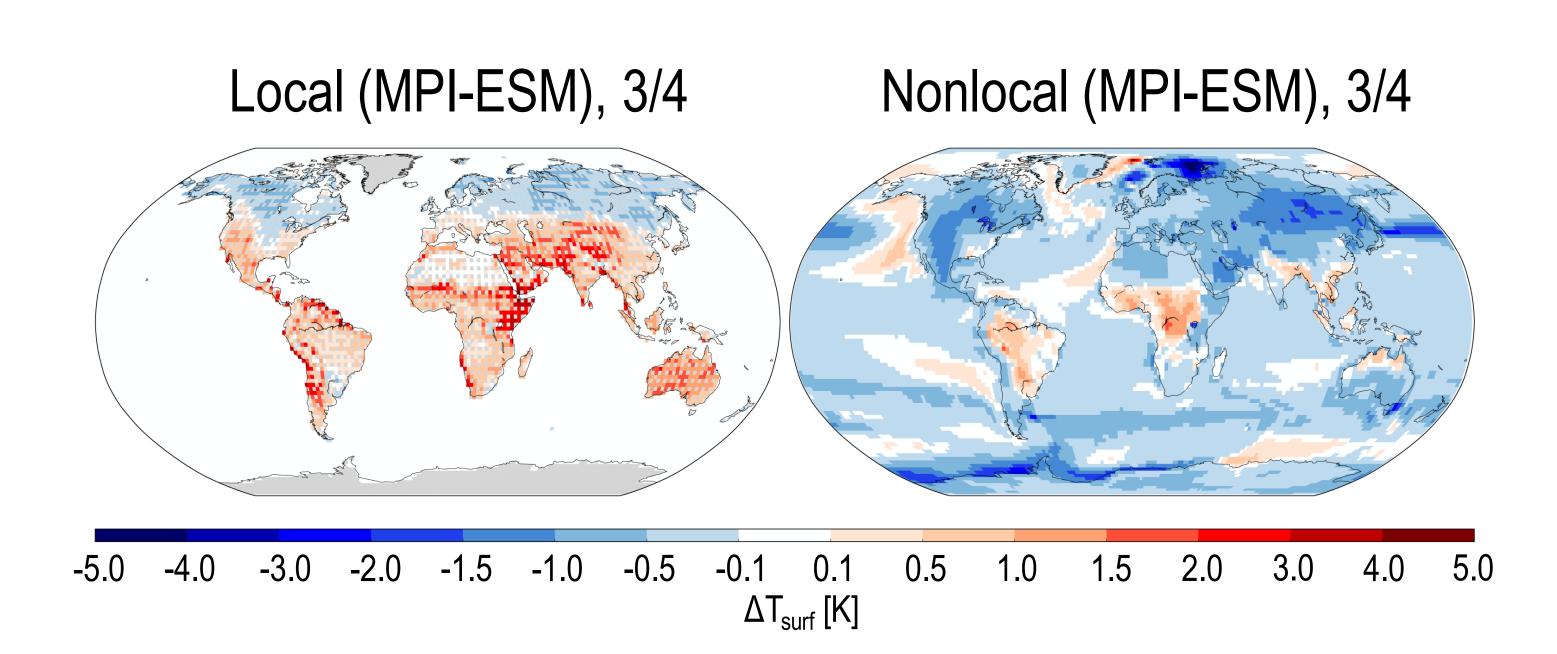
Role of changes in surface albedo

- The increase in albedo is responsible for the deforestation-induced biogeophysical cooling.
- Locally, this cooling is partly compensated for by changes in the sensible and latent heat fluxes (not shown).
- The albedo-induced cooling is mainly included in the nonlocal effects. Thus models may seem to cool stronger than observations.

Left: Complete deforestation in 1/4, 2/4 and 3/4 of all land boxes, spatially homogeneous distribution of deforestation. Middle: Extent of deforestation like 1/4, but only in the low, mid, or high latitudes and near historically deforested regions. Magnitude of carbon warming estimated using a bookkeeping approach [5] and the transient response to cumulative emissions [6]. Right: Changes in areas of 3/4, only albedo is changed to grass while other surface properties of trees remain.

- Observations based on in situ (FluxNet) [2] or
- MPI-ESM: Local effects as in [1] within the

Not captured by observations, and thus contribute to mismatch with simulated effects.







Nonlocal cooling is stronger than local warming

(local : nonlocal = 1 : -3 for homogeneous spatial deforestation).



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