



Pyplis* | A Python based software package for the analysis of volcanic SO2 emissions using UV SO2 cameras

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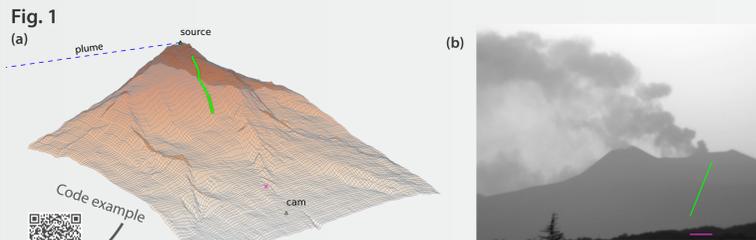


Main Features

- △ Easy setup of various data formats and camera specifications
- △ Detailed 3D geometrical calculations
- △ Routine for image based correction of the signal dilution effect
- △ Flexible options for retrieval of plume background intensities
- △ Engines for automatised cell and DOAS calibration
- △ Optical flow and cross correlation based plume velocity retrievals
- △ Simultaneous emission rate retrievals for multiple retrieval lines

* due to a naming conflict, the software was renamed from *piscope* to *pyplis*

1 Measurement geometry and signal dilution correction

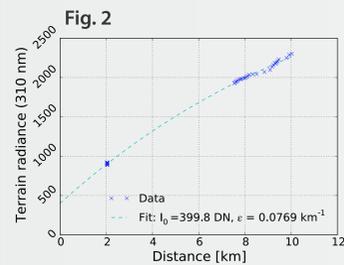


Infobox 1 | Geometrical calculations

Plume distances (required for emission rate retrieval) are retrieved on a pixel-level based on camera coordinates and plume azimuth. **More features include:**

- Retrieval of camera viewing direction using distinct features in the images (e.g. volcano summit)
- Retrieval of distances to local terrain features in the images on pixel-level (see Fig. 1)

Fig. 1a: 3D terrain map of the measurement setup including 2 topographic profile lines (green and magenta, defined in the image, cf. Fig. 1b) used for the signal dilution analysis (Fig. 2)
 Fig. 1b: Exemplary on-band (310 nm) image. Due to SO2 absorption the Etna plume appears darker than the sky background. Measured intensities along the two indicated lines are used to estimate the scattering extinction (cf. Fig. 2).



Equation 1 | Atmospheric scattering model

$$I_{\text{meas}}(\lambda) = I_0(\lambda)e^{-\epsilon(\lambda)d} + I_A(\lambda)(1 - e^{-\epsilon(\lambda)d})$$

$I_{\text{meas}}(\lambda)$: measured intensity of terrain feature
 $I_0(\lambda)$: initial intensity of terrain feature
 d : distance to terrain feature
 $I_A(\lambda)$: average ambient intensity
 $\epsilon(\lambda)$: scattering extinction coefficient

Fig. 2: Retrieval of extinction coefficients
 Exemplary dilution analysis (cf. Infobox 2) using the two topographic lines shown in Fig. 1, resulting in an on-band extinction coefficient of $\epsilon_{310} = 0.0769 \text{ km}^{-1}$.

Infobox 2 | Image based correction for the signal dilution effect
 Distant terrain features appear brighter (cf. lines in Fig. 1b). This is due to atmospheric scattering and can result in significant underestimations in the emission rates. A first order correction for this so called signal dilution effect can be performed using local terrain features in the images located at different distances [1]. The method yields atmospheric scattering extinction coefficients (Fig. 2) using an approximative scattering model (Eq. 1) and brightness values measured along suitable terrain features in the images (cf. Fig. 1). The required distances to these features can be retrieved automatically in Pyplis (cf. Infobox 1, Fig. 1).

2 Plume background analysis

Equation 2 | Optical density (OD)

$$\tau = \ln\left(\frac{I_0}{I}\right)$$

I : measured plume intensity
 I_0 : corresponding background intensity

Equation 3 | SO2 apparent absorbance (AA)

$$\tau_{AA} = \tau_{\text{on}} - \tau_{\text{off}}$$

Infobox 3 | Calculation of optical density (OD) images

The calculation of OD images (Eq. 2) requires knowledge of the sky background intensities behind the plume. These can either be retrieved from the plume images directly (by fitting a 2D polynomial to clear sky areas, cf. Fig. 3a) or using a sky reference image which is scaled using a gas free sky area in the plume image (cf. *scale_rect* in Fig. 3). The latter may include further corrections to account for variations in the sky background curvature between the two images (cf. Fig. 3b).

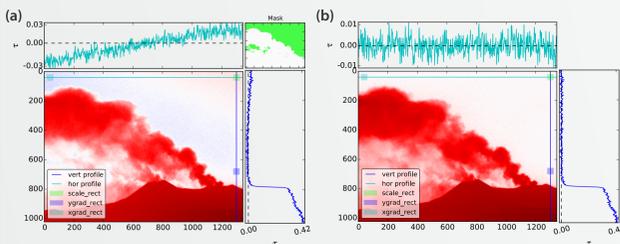
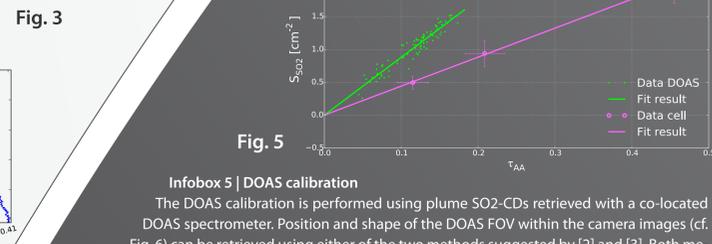


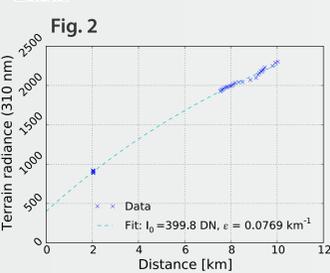
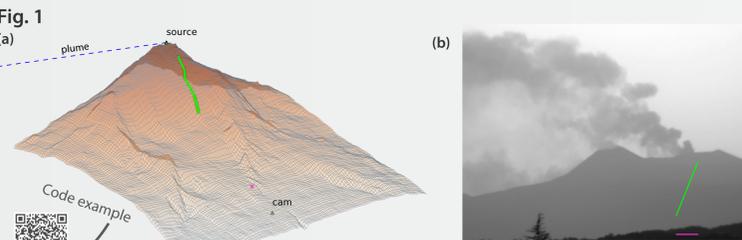
Fig. 3: Exemplary optical density images
 Two exemplary on-band OD images calculated using a masked 2D polynomial fit (a) and using an additional sky background image (b), including 2. order horizontal and vertical sky gradient corrections along the indicated vertical and horizontal profile lines.



Infobox 5 | DOAS calibration

The DOAS calibration is performed using plume SO2-CDs retrieved with a co-located DOAS spectrometer. Position and shape of the DOAS FOV within the camera images (cf. Fig. 6) can be retrieved using either of the two methods suggested by [2] and [3]. Both methods identify the pixel region showing highest correlation between a time series of DOAS SO2-CDs and camera images. The calibration curve (Fig. 5) is then retrieved from a time series of AA values (extracted within the DOAS FOV) and the corresponding DOAS SO2-CDs. The DOAS calibration is more accurate in case aerosols are abundant in the plume [2].

3 Camera calibration



Equation 4 | SO2 column density

$$S_{\text{SO}_2} = \int_{\mathcal{C}} \epsilon_{\text{SO}_2}(x, y, z) ds$$

S_{SO_2} : SO2 - column density (CD)
 \mathcal{C} : photon trajectory
 ϵ_{SO_2} : SO2 concentration distribution

Infobox 4 | Cell calibration
 The camera calibration (conversion of ODs into SO2-CDs (see Eq. 4) can be performed by consecutively placing a number of SO2 gas cells in front of the camera while pointing into a clear sky area. The calibration curve (cf. Fig. 5) is retrieved from cell AA images and the corresponding (known) SO2-CDs of the cells.

4 SO2 emission rate retrieval

Equation 5 | Emission rate integral

$$\Phi(\ell) = f^{-1} \sum_{m=1}^M S_{\text{SO}_2}(m) \cdot \langle \mathbf{v}(m) \cdot \hat{\mathbf{n}} \rangle \cdot d_{\text{pl}}(m) \cdot \Delta s(m)$$

$\Phi(\ell)$: emission rate through PCS line ℓ
 f : camera focal length
 $S_{\text{SO}_2}(m)$: SO2 - CD at pixel pos. m on ℓ
 $\mathbf{v}(m)$: gas velocity vector at pixel pos. m on ℓ
 $\hat{\mathbf{n}}$: normal vector of ℓ
 $d_{\text{pl}}(m)$: plume distance at pixel pos. m on ℓ
 $\Delta s(m)$: discrete integration step at pixel pos. m on ℓ

Infobox 8 | SO2 emission rate analysis

Emission rates are retrieved from calibrated SO2-CD images using Eq. 5. The integration can be performed along linear plume intersections (see e.g. Fig. 6). Plume distances are calculated from the measurement geometry (see Infobox 1). Gas velocities can either be retrieved using an optical flow algorithm or using signal cross correlation (cf. Infoboxes 6 & 7).

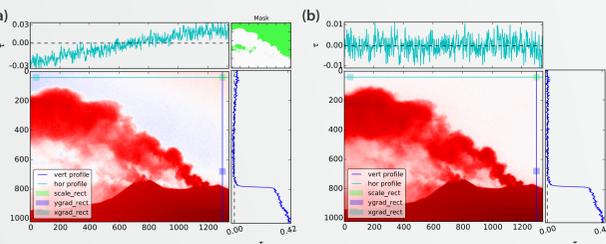
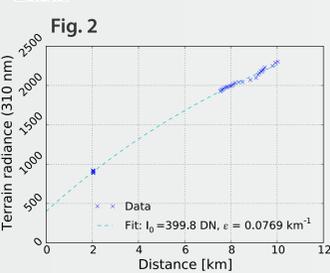
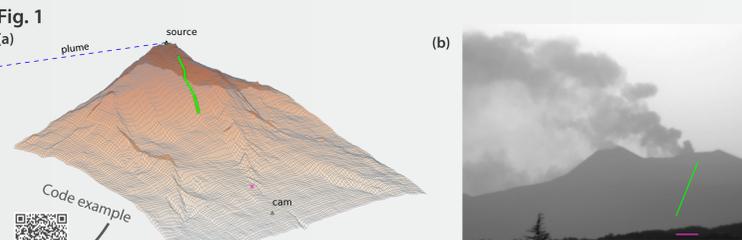


Fig. 6: Calibrated SO2 CD image
 Figure showing a calibrated SO2-CD image including two exemplary PCS lines and the position of the DOAS FOV.

5 Gas velocities



Infobox 6 | Cross correlation
 The plume velocity is retrieved from a time-series of images based on the cross correlation lag between the integrated-column-amount (ICA) of two plume intersections (e.g. [4], cf. Fig. 7).

Infobox 7 | Optical flow
 Optical flow (OF) algorithms detect motion in consecutive frames allowing for the retrieval of gas velocities on pixel-level (e.g. [5]). An estimate of the gas velocity within low contrast image regions (where the OF algorithm fails to detect motion) can be inferred based on distinct peaks in histograms calculated from the optical flow field (Fig. 8). Pyplis includes an implementation of the Farneback optical flow algorithm [6].

6 Links and Downloads

Website: <http://pyplis.readthedocs.io/en/latest/index.html>

Source code (Github): <https://github.com/jgliss/pyplis>

PyPi: <https://pypi.python.org/pypi/pyplis>

Web

Get the poster

More plots

Watch video

Code example

7 References

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[4] McGoigle, A. J. S., Hilton, D. R., Fischer, T. P., Oppenheimer, C., 2005. Plume velocity determination for volcanic SO2 flux measurements. *Geophysical Research Letters* 32 (11), 111302

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[6] Farneback, G., 2003. *Two-Frame Motion Estimation Based on Polynomial Expansion*. Springer Berlin Heidelberg, Berlin, Heidelberg, pp. 363-370