



# The effects of rheological and tectonic parameters on the preservation of primordial reservoirs in Earth's lower mantle: a numerical study

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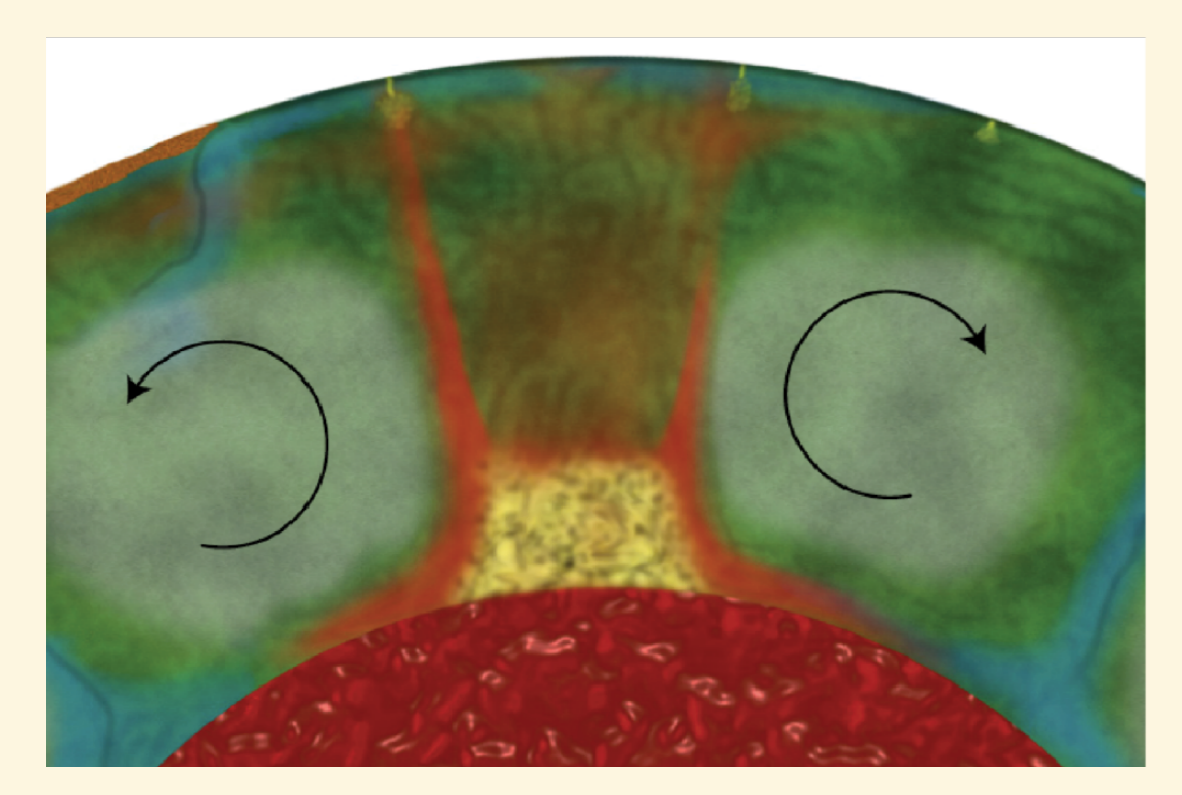
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## Motivation

Is the Earth's lower mantle heterogeneous in composition?

- Homogeneous, well-mixed mantle**
  - Seismic imaging of subducted lithosphere<sup>[1]</sup> and deep-rooted plumes<sup>[2]</sup> that pervade the whole mantle
  - Shallow plumes can spatially be related to LLSVPs<sup>[3]</sup>
- vs. Mid-mantle chemical heterogeneities**
  - Stagnating slabs at ~1000 km depth<sup>[4]</sup>
  - Sharp seismic impedance contrasts at a similar depth range<sup>[5,6]</sup>
  - Primordial <sup>182</sup>W/<sup>184</sup>W<sup>[7,8]</sup> and <sup>3</sup>He/<sup>4</sup>He<sup>[9]</sup> signatures in basalts

These discrepancies are incorporated in a recently proposed convective regime<sup>[10]</sup>:



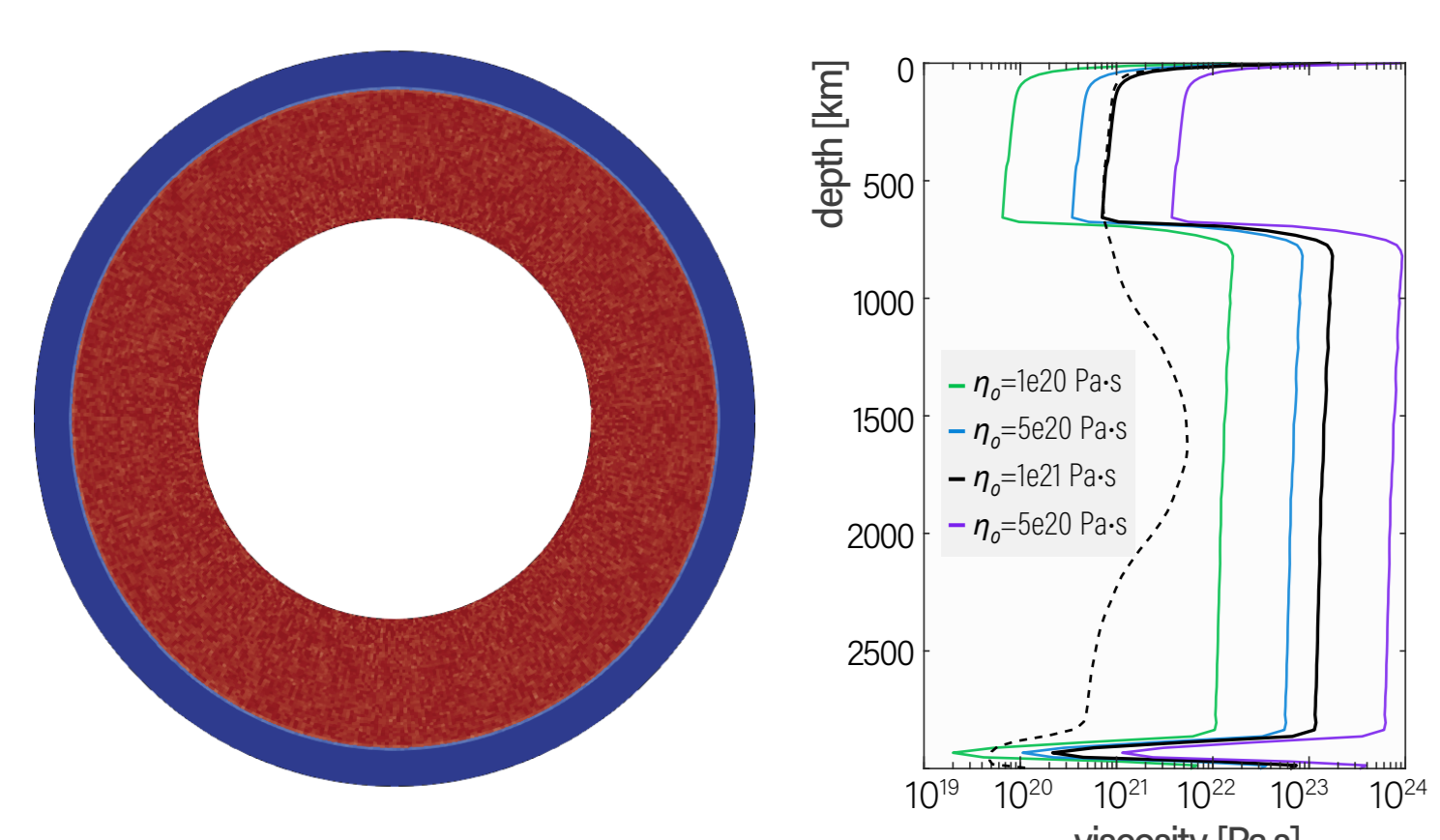
- Large, intrinsically **viscous domains** persist in the mid-mantle, with whole-mantle circulation being accommodated around them
- Domains believed to be enriched in the strong lower-mantle mineral **bridgmanite** (Fe,Mg)SiO<sub>3</sub>, resulting from a fractional crystallizing magma ocean

This regime has successfully been reproduced in 2D spherical annulus geometry models using composition-dependent rheology<sup>[11]</sup>, but lacks a thorough assessment of tectonic/rheological parameters.

**What are the effects of rheological and tectonic parameters on the style of mantle mixing and heterogeneity preservation?**

## Methods

- Thermomechanical convection** with 2D spherical annulus geometry using StagYY<sup>[12]</sup>
- Initial two-layered set-up** with **composition-dependent rheology**<sup>[11]</sup>: compared to **pyrolite**, **primordial material**
  - is more **viscous** in the lower mantle (SiO<sub>2</sub> enrichment)
  - is slightly more **dense** (FeO enrichment)
  - has a higher **bulk modulus** in the lower mantle<sup>[13]</sup>

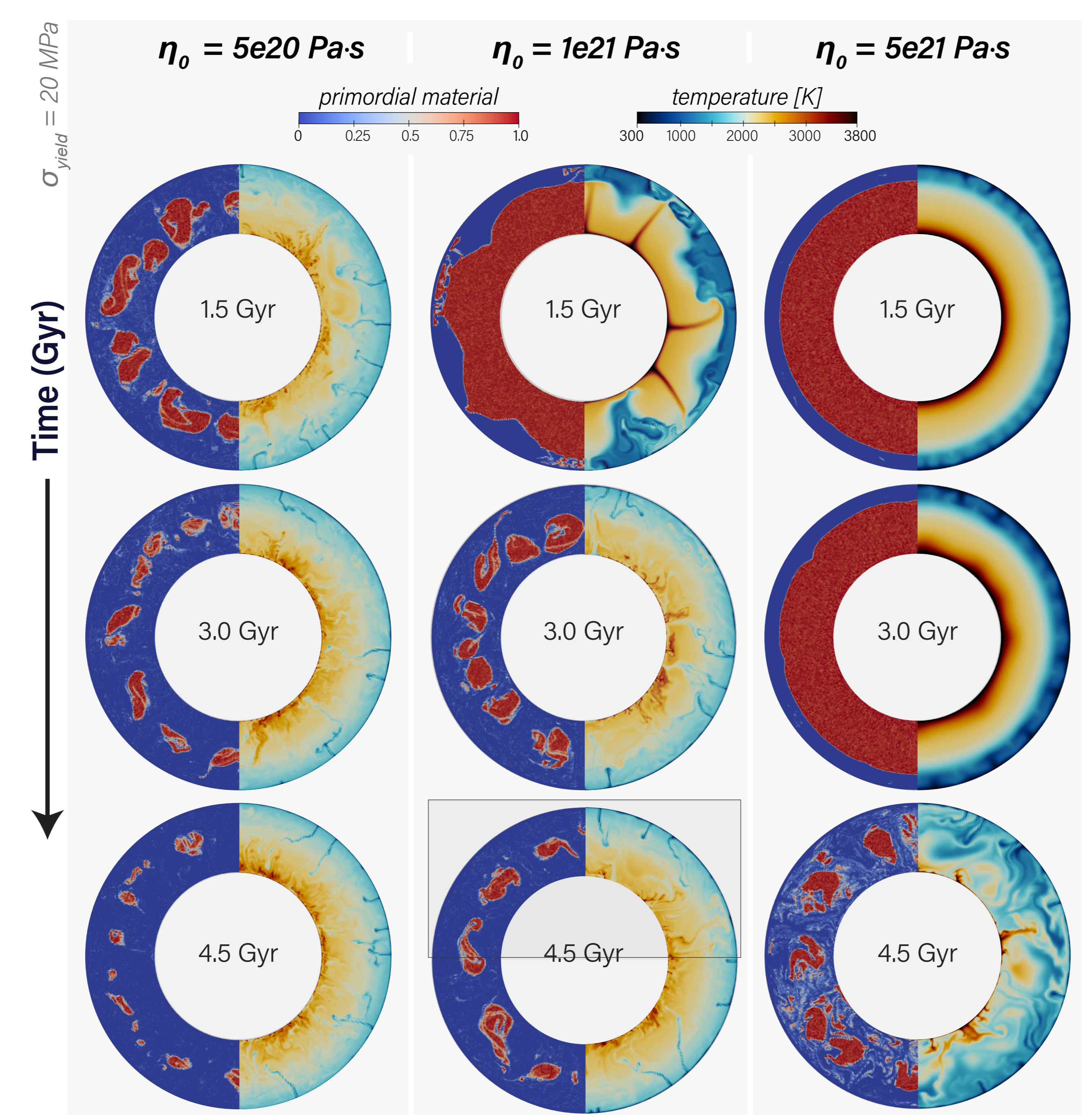


(Top left) Initial model set-up: 2230 km-thick primordial layer in the lower mantle and pyrolytic material in the upper mantle, resolved by a grid of 512x96 cells. (Top right) Initial viscosity profiles of models with different reference viscosities, dotted profile is the reference model at 4.5 Gyr.

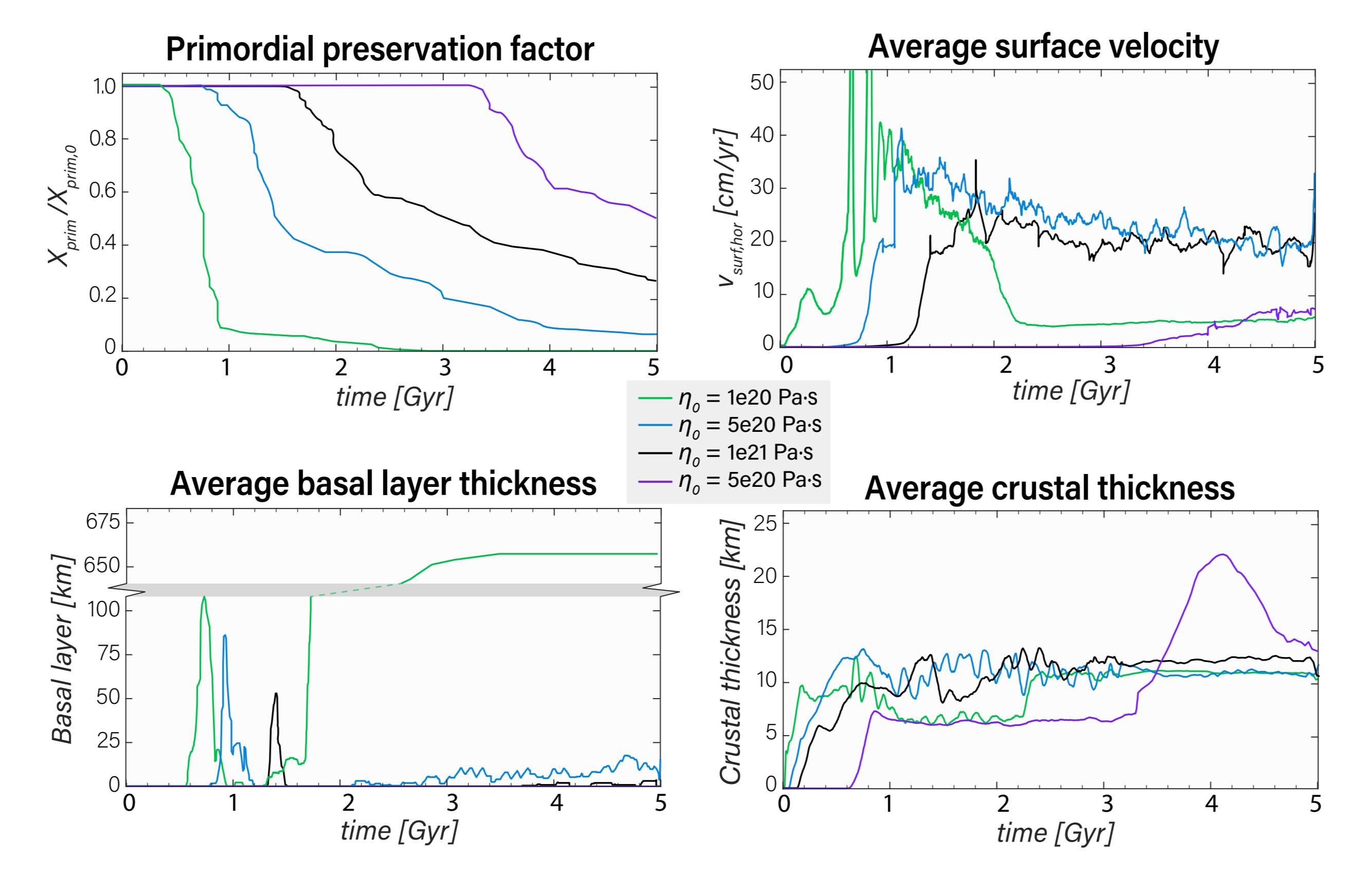
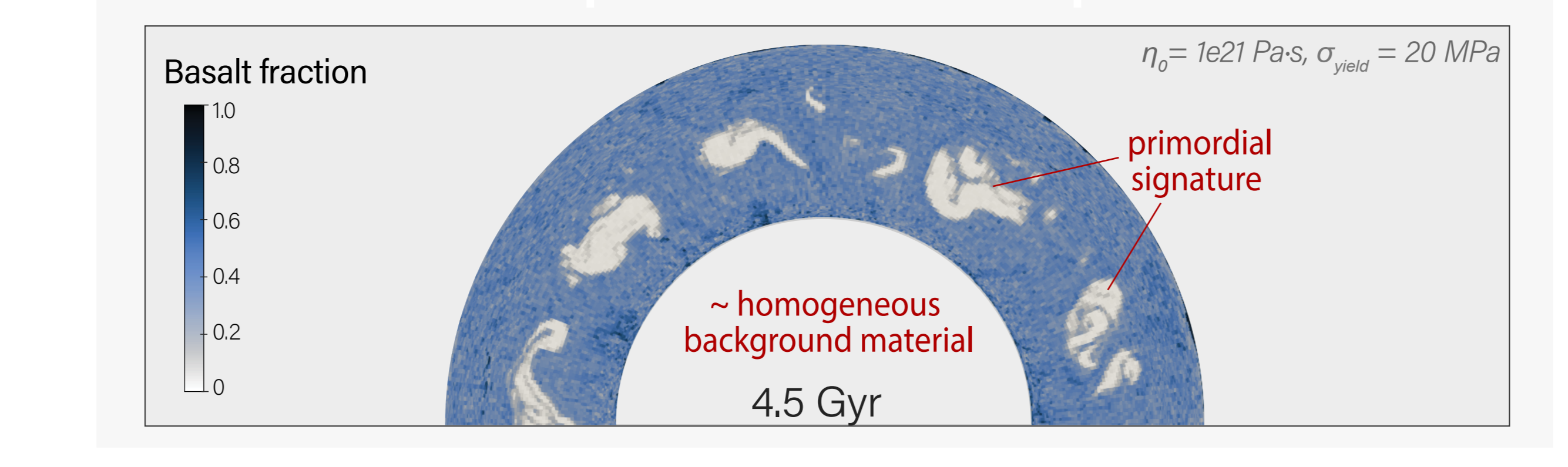
Symbol	Definition	Value
$h$	Mantle domain thickness	2980 km
$g$	Gravitational acceleration	9.81 ms <sup>-2</sup>
$T_s$	Surface temperature	300 K
$T_{CMB}$	CMB temperature	4000 K
$T_a$	Reference temperature	1600 K
$\sigma_{yield}$	Yield stress	20e6 MPa*
$\sigma'_{yield}$	Yield stress depth derivative	0.01
$\eta_0$	Reference viscosity	1e21 Pa·s*
$E$	Activation energy	35 KJ/mol
$\lambda_{pm}$	Lower mantle viscosity contrast	300
$\Delta\rho_{pm}$	Surface density contrast	0.4%
$K_{a,pm}$	Lower mantle bulk modulus (primordial)	230 GPa
$K_p$	Lower mantle bulk modulus (pyrolite)	210 GPa

- Visco-plastic rheology** in which viscosity is T-dependent following the Arrhenius law:
 
$$\eta(T, P) = \eta_0 \exp\left(\frac{E}{RT} - \frac{E'}{RT_0}\right)$$
- "Plastic yielding"** breaks stagnant-lid to give plate-like behavior:
 
$$\sigma_{yield}^{eff} = \sigma_{yield} + \sigma'_{yield} P$$
- Phase changes** at 410, 660 and pPv boundaries
- Primordial-to-basalt/harzburgite** tracer conversion at 125 km depth (depth of pyroxenite melting<sup>[14]</sup>). A **melting law** produces basalt and residu harzburgite from mantle material.

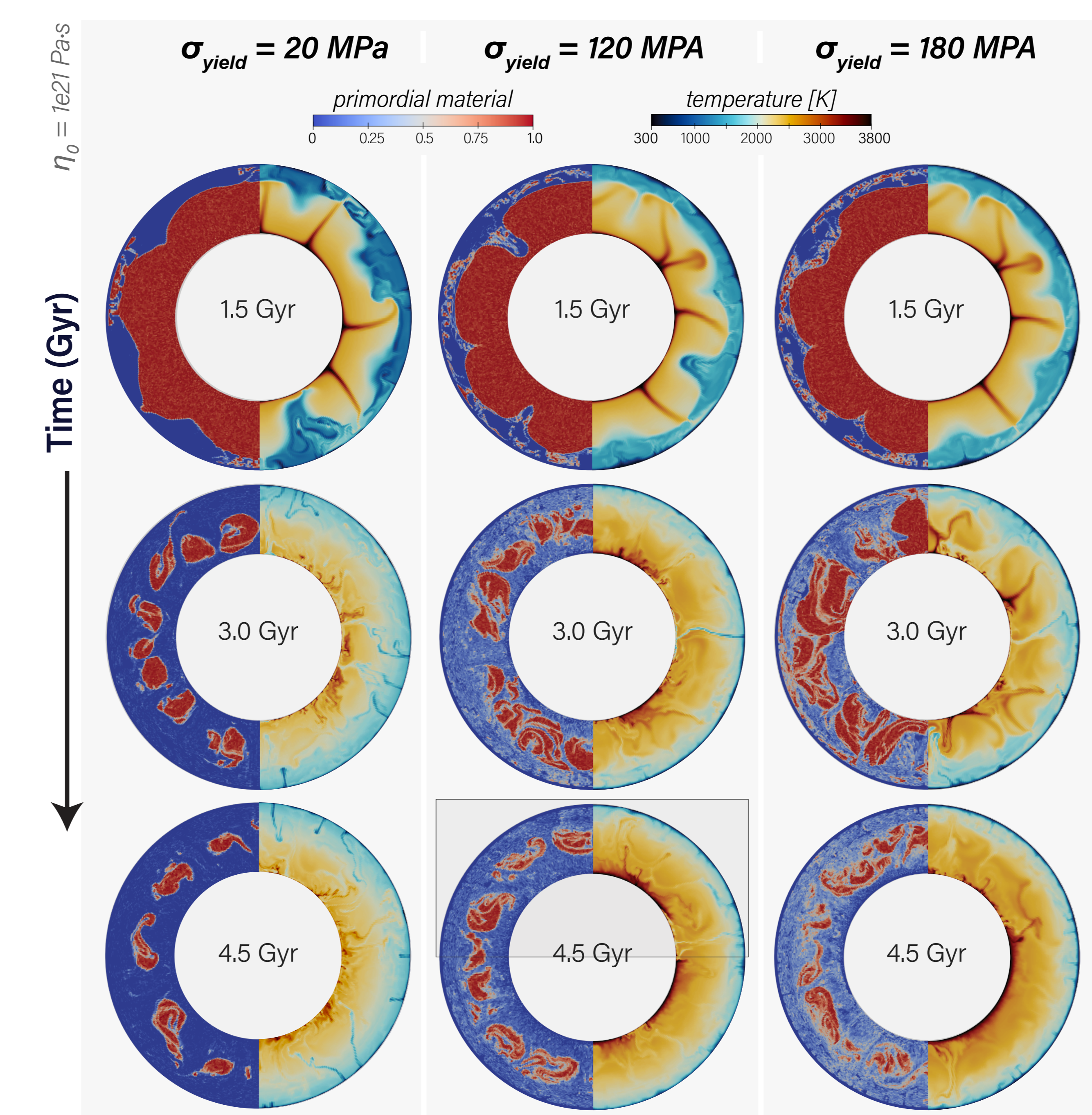
## Results: reference viscosity



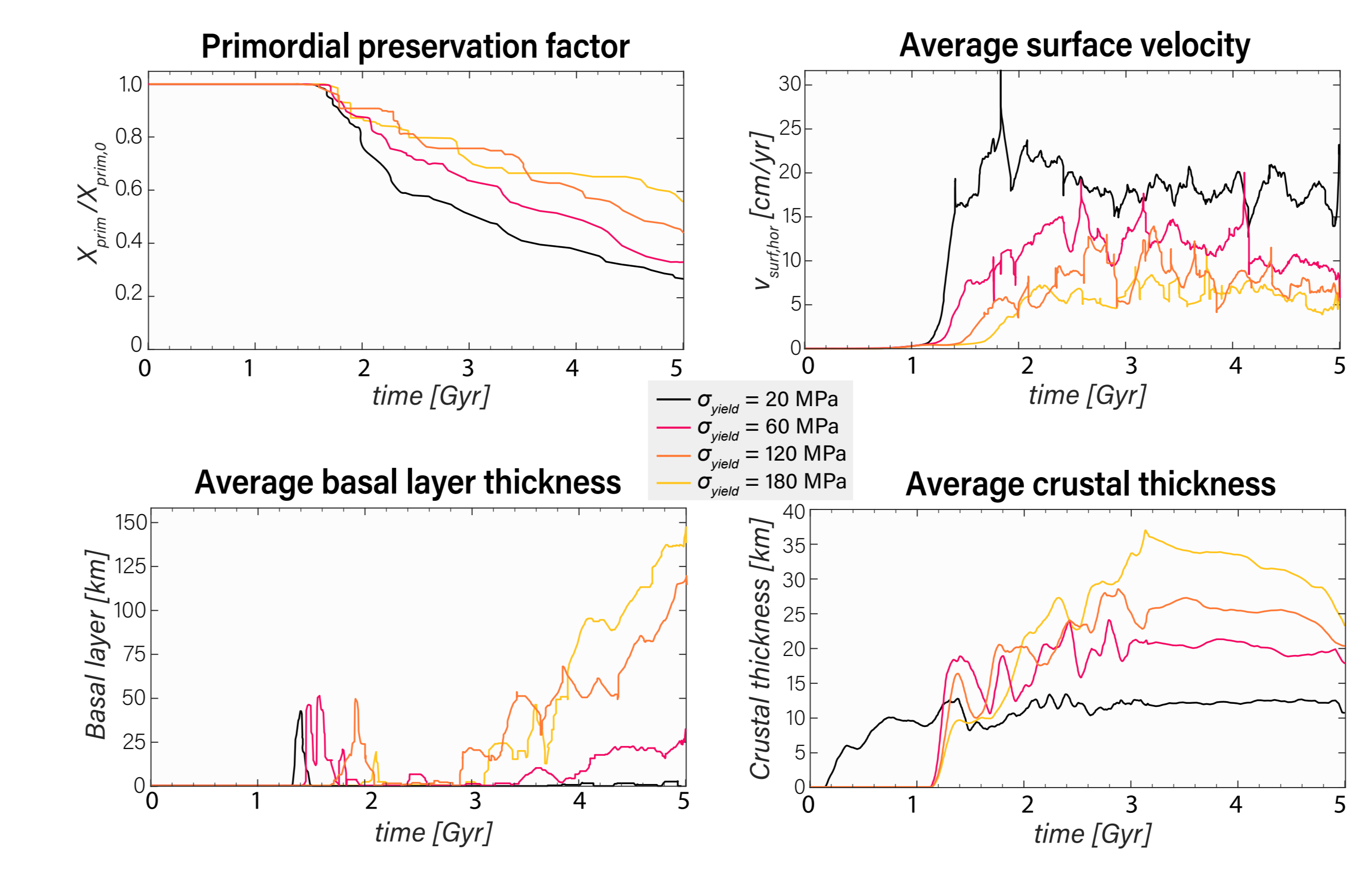
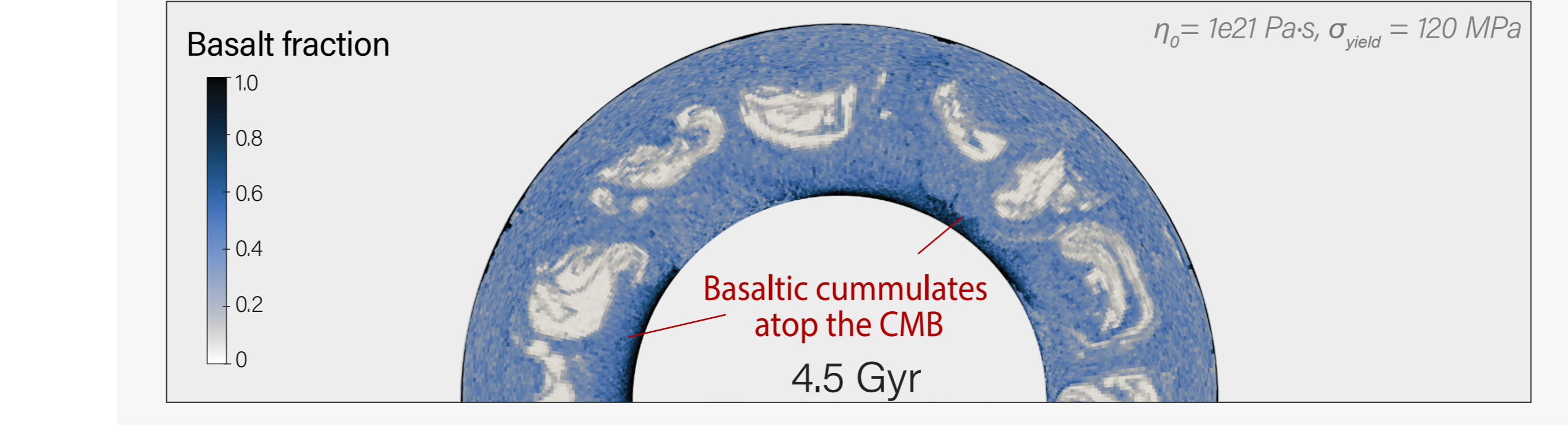
- 5e20 Pa·s**: Early overturn, Higher efficiency of mantle mixing
- 1e21 Pa·s**: Moderate heterogeneity preservation as primordial blobs in the mid-mantle (*reference model*)
- 5e21 Pa·s**: Very late overturn, Inefficient mixing of mantle material



## Results: yield strength



- 20 MPa**: Frequent lithospheric drips into lower mantle, Efficient mixing of mantle material in background flow (*reference model*)
- 120 MPa**: Fewer downwellings and a warmer lower mantle, Primordial material entrained in background flow
- 180 MPa**: Hot lower mantle, Fragmented primordial domains in the mid-mantle, Inefficient mixing of mantle material in background flow



## Conclusions

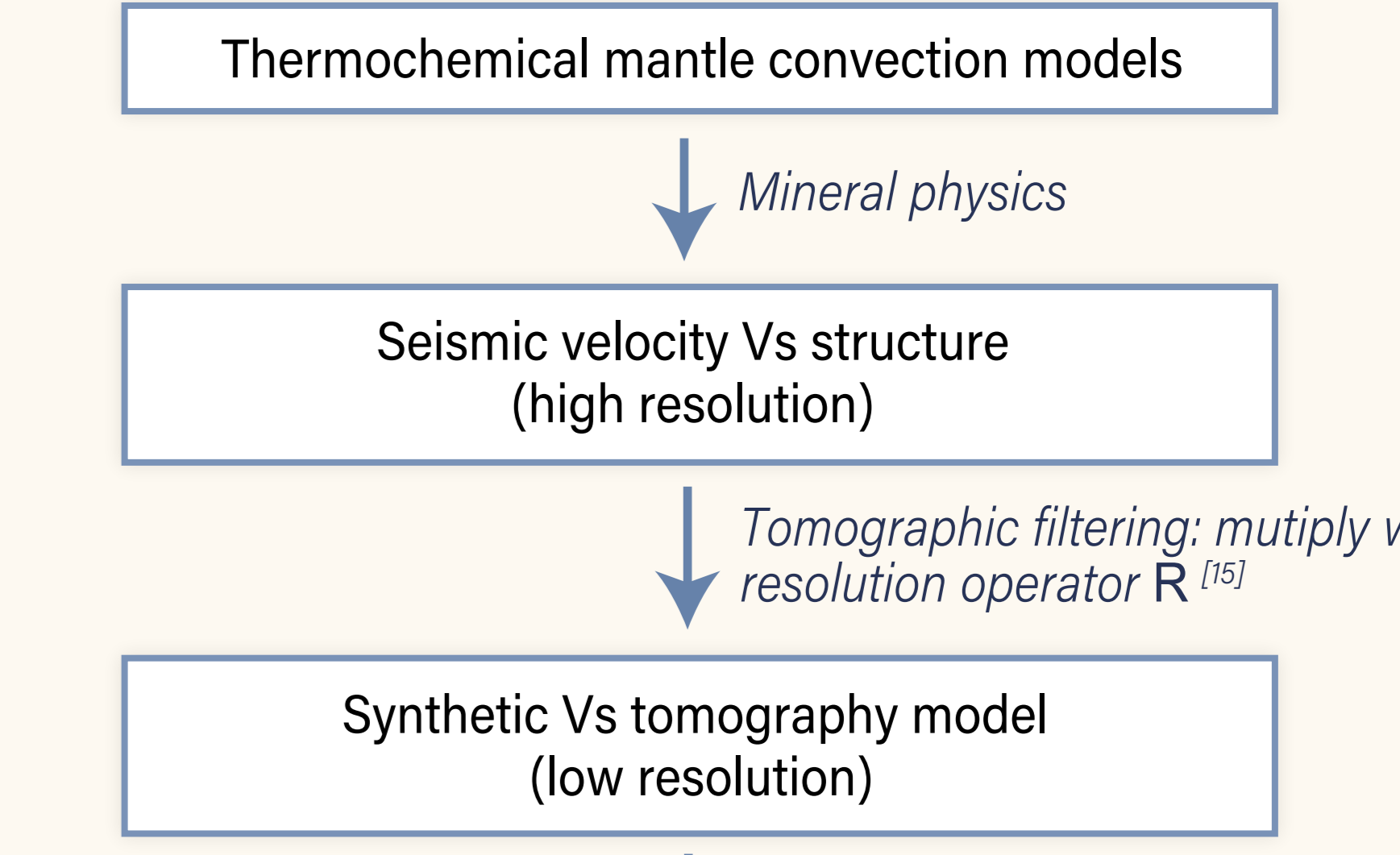
The style of mantle mixing and heterogeneity preservation is greatly influenced by the tectonic style and mantle rheology:

- The reference viscosity affects the **convective vigor** and thereby the **timescale** of model evolution; lower values promote efficient mixing of primordial material into the ambient mantle.
- The plate yielding strength controls the **abundance** and **stiffness** of subducting slabs that in turn interact with the primordial domains.
- Stronger slabs enhance primordial **fragmentation** but **reduce mixing efficiency** of primordial with pyrolytic material.
- In addition, greater yield strengths promote a **thicker basaltic crust** at the surface and **accumulation of basaltic material** atop the CMB, underlying the neutrally buoyant bridgmanitic-enriched domains in the mid-mantle.

## Outlook

- Future work will involve:**
- Expanding the rheological parameter space by including the effect of T- and P-dependence of viscosity (via the Arrhenius law), and assessing which models are most Earth-like
  - Integrating our numerical results with geophysical observations: (*suggestions welcome*)

How would our geodynamic models translate to seismic tomography?



- Exploring heterogeneity preservation in **3D numerical models**
- Incorporating strain-dependent rheology as a means to focus deformation in the weaker, relatively bridgmanite-depleted mantle pyrolite<sup>[16]</sup>

## References

- van der Hilst, R. S., Widiyantoro, and E. Engdahl (1997), Nature, 386, 578-584.
- French, S., and B. Romanowicz (2015), Nature, 525, 95-99.
- Burke, K., Steinberger, B., Torsvik, T. H., & Smethurst, M. A. (2008). Earth and Planetary Science Letters, 265, 49-60.
- Fukao, Y., & Obayashi, M. (2013). Journal of Geophysical Research, 118, 5920-5938.
- Waszek, L., N. Schmerr, and M. Ballmer (2018), Nature Communication, 10(3), 236-240.
- Jenkins, J., A. Deuss, and S. Cottar (2017), Earth and Planetary Science Letters, 459, 196-207.
- Rizo, H., R. Walker, R. Carlson, M. Horan, S. Mukhopadhyay, V. Manthos, D. Francis, and M. Jackson, (2016). Science, 352, 809-812.
- Mundl, A., M. Touboul, M. Jackson, J. Day, M. Kurz, V. Lekic, R. Helz, and R. Walker, (2017). Science, 356, 66-69.
- Graham, D., P. Michael, and T. Shea (2016), Earth and Planetary Science Letters, 454, 192-202.
- Ballmer, M., C. Houser, J. Hernlund, R. Wentzovitch, and K. Hirose, (2017), Nature Geoscience, 10, 236-240.
- Gebhardt, D., Gülcher, A.J.P., Ballmer, M.D. and Tackley, P.J. [in prep]
- Hernlund, J., and P. Tackley (2008), Physics of the Earth and Planetary Interiors, 171(1-4), 48-54.
- Li, B., and J. Zhang (2005), Physics of the Earth and Planetary Interiors, 255, 80-92.
- Petermann, M., and Hirschmann, M. M. (2003). Journal of Geophysical Research, 108(B2), 2125.
- Koelemeijer, P., Schuberth, B.S.A., Davies, D.R., Deuss, A. and Ritsema, J. (2018), Earth and Planetary Science Letters, 494, 226-238.
- Girard, J., Amulele, G., Farra, R., Mohiuddin, A. and Karato, S. (2016), Science, 351, 144-147.