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### . Study area & rationale

- The study area is located in the Lahti region where moraine ridges, Salpauselka I and Salpauselka II, "bend"
- The ridges were deposited during sillstands o the Fennoscandian ice sheet linked to the Younger Dryas (Rinterknecht et al., 2004; Saarnisto and Saarinen, 2001)
- Ice margin was, at that time, grounded in a large proglacial lake, the Baltic Ice Lake.
- The sediments were deposited as ice-contact deltas, over a very short period (180-250 years)

# 2. Morphology & metrics

- High resolution LIDAR data (2m horizontal, 0.3m vertical) was used to map the morphology
- The deltas, at present, form flat topped hills standing proud of the topography
- Distributary channels network still visible on the flat delta top
- Multiple kettle holes formed where buried ice blocks

Study area

- Ground penetrating radar(GPR) profiles and outcrops analyzed and interpreted providing information on formation mechanism, sedimentary environments and reservoir properties of ice-contact deltas
- This study aims to provide insights for analogous deposits (Pleistocene and older) which can be prospective for hydrocarbons or water reservoirs.





- Sediments supplied to the ice margin via meltwater conduits and/or bedrock troughs
- narrow, steep ridge was formed
- smaller ones • Evidence for ice margin oscillations and partial







## 4. Ground penetrating radar survey results

- GPR profiles acquired using 80 MHz antenna
- Penetration depth ranges between 15 m and 25 m
- Vertical resolution 0.2m-0.5m
- Multiple steeply-dipping reflectors interpreted as delta foresets
- Discontinous and/or chaotic reflectors in the ice proximal part of the delta interpreted as ice contact/ ice proximal facies
- Flat reflectors in the upper part of the delta interprete as delta topset facies
- Bedrock highs often blanketed by glaciofluvial, deltaic sediments
- Reflections are mostly discontinuous implying erosional contacts within the strata
- Clear angular unconformity between topset and foreset facies
- GPR penetration depth not enough to image the whole thickness of strata



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Ground penetrating radar (GPR) profile B-B' across an ice-contact delta (perpendiculat to the depositionl dip) Hyrtiälänkangas, Salpausselka 2, Finland



Ice - proximal slope Ground penetrating radar (GPR) profile A-A' across an ice-contact delta (ice-proximal to ice-distal) Hyrtiälänkangas, Salpausselka 2, Finland **Delta topsets** Bedrock high





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<u>Site 2</u>





Site 5





### 5. Conclusions

- World-class example of ice-contact, Gilbert-type deltas which can be used as a direct analogue for similar deposits in the subsurface
- Mostly good, reservoir-quality sediments in the medial and distal part of the delta
- Topsets often composed of boulders and cobbles which may cause potential drilling hazards
- Ice contact (Gilbert type) deltas rather than "typical" frontal moraines
- Rapid deposition controlled by the ice margin position as it retreats (backfilling) and depth of the proglacial lake
- Abrupt facies change laterally and distally from the efflux point at a scale of 10s/100s of meters
- Kettle hole facies can be misinterpreted ice distal sediments

- Large ice contact deltas are an amalgamation of several lobes (bajada-
- The largest and but also the most complex deltas formed in the interlobate zone (a rule?)
- Ice contact facies are considered as nonreservoir facies due to their poor sorting.
- GPR profiles provide good understanding of the structure of the delta which can be ground-proved in isolated, moderate quality outcrops.
- Another technique (electrical resistivity tomography?) required to image the full thickness of sediments.

### References

https://doi.org/10.1016/S0921-8181(01)00131-X

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