Current level and rate of warming determine emissions budgets under ambitious mitigation Nicholas Leach^{1,2} Richard Millar² Karsten Haustein² Stuart Jenkins^{1,2} Myles Allen^{1,2}

Based on a Nature Geoscience Article of the same name: rdcu.be/2Use

1) Dept. of Physics, University of Oxford

2) Environmental Change Institute, University of Oxford



We can estimate the remaining **carbon budget** for ambitious mitigation targets using observations of the **current level and rate** of warming

Our calculation is **simple**, **transparent**, and both model and response independent





Navigation





The Budget Problem

- Over the last few years, there have been a wide range of estimates of the remaining carbon budget to key temperature targets.
- Here is a graphic from CarbonBrief showing just how large the difference between many of the estimates for the budget to 1.5°C is.



Remaining carbon budget for a 66% chance of less than 1.5C warming



CB



The Budget Problem

- These discrepancies are due to a number of geophysical constraints, such as:
- The current level of warming
- The definition of pre-industrial
- The current and future level of non-CO₂ forcing
- We hope to provide a simple framework for calculating budgets that can explain these differences.



Remaining carbon budget for a 66% chance of less than 1.5C warming



CB



• This shows one CO₂ emission scenario from the AR5 database.







• Here is the temperature anomaly within the same scenario.









- If we integrate annual emissions to get cumulative CO₂ emissions...
- ... we find that they are almost proportional to temperature anomaly, so we can write:

$$T=r\int E\,\mathrm{d}t$$







 Clearly the rate of change of temperature anomaly is also almost proportional to annual CO₂ emissions, so we have:

$$T = r \int E \, \mathrm{d}t$$

 $\mathrm{d}T$ _, _



dt





- The constant of proportionality, *r*, is an emergent property of ESMs consistent with observations of the climate system called the Transient Response to Cumulative Carbon Emissions (TCRE).
- It is only strictly true for CO₂ emissions and the warming that they directly cause:

$$T_{CO_2} = r \int E_{CO_2} \, \mathrm{d}t$$







 Here we show observations of global mean temperature anomaly from HadCRUT4, relative to 1850-1900.







- We decompose this using an OLS regression into anthropogenic and natural signals based on best estimates of the corresponding forcing components.
- This allows us to calculate the current warming rate by reducing the noise without changing the current level (present day natural warming is ≈ 0.02°C).







 Here we show the rate of change of anthropogenic warming, currently around 0.22°C / decade.







- We now extend the warming trajectory in two ways:
- A linear reduction in warming rate (quadratic stabilisation)
- Keeping a constant warming rate (linear temperature increase)







• Focusing on the interesting time period...







- Marking three key times on the diagram:
- 1. Present day
- 2. The time to 1.5°C for a constant rate of future warming
- 3. The time to stabilisation for a linear reduction in the warming rate







• To find the crossing time, simply divide the warming remaining, ΔT (the shaded area under the constant rate curve), by the current warming rate, T_0 :

 $\rho = \Delta T/T_0'$







We call the crossing time, *ρ*, the mitigation timescale, since it is also half the time to stabilisation – the shaded areas under the red and blue lines are equal (proof here).

$\rho = \Delta T/T_0'$

Note that for every year's delay in reducing emissions, as long as warming continues at the current rate, *p* falls by one year, and hence the time remaining to reduce the warming rate linearly to zero to meet any given temperature stabilization goal falls by two years.







- The TCRE then allows us to treat these warming trajectories as if they were CO₂forcing-equivalent emission timeseries.
- This means we can calculate the remaining CO₂ budget to 1.5°C:

 $ho = \Delta T/T_0'$

$$\int E \, \mathrm{d}t = E_0 \cdot \rho$$







- As mentioned before, \bullet this is only strictly true for CO₂-forcing-equivalent emissions, or using CO_2 induced warming.
- However, the equations ٠ below do still hold generally for many ambitious scenarios - see next section...

 $\rho = \Delta T / T'_0$ $\int E \, \mathrm{d}t = E_0 \cdot \rho$







- *ρ* is also equal to the time constant for an exponential decay in emissions (or equivalently warming rate).
- Click **here** for all the mathematical derivations behind this section.







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Model Scenario Experiments

- This shows the budget predicted using our equation (below) versus the actual budget within a subset of the AR5 scenarios (subset details here).
- We have calculated the budget with *total* warming and CO_2 emissions:

 $\int E \, \mathrm{d}t = E_0 \cdot \Delta T / T'_0$







contents

Model Scenario Experiments

- Two questions arise:
- Why do most of the 1. lines lie below the 1:1 line?
- Why does the 2. scenario shown in the first section (circled) lie so close to the 1:1 line?





- 430-480 ppm
- 480-530 ppm
- 530-580 ppm

580-650 ppm 650-720 ppm



Model Scenario Experiments

- These are both due to the future non-CO₂ forcing fraction.
- If the non-CO₂ forcing fraction increases into the future, it "takes up" more warming than predicted by our equation, therefore reducing the CO₂ budget, i.e.:

$\int E\,\mathrm{d}t\approx (E_0+\alpha F_0')\cdot\rho-\alpha\Delta F,$

where ΔF is the net change in non-CO₂ forcing between the present day and peak warming, and α is approximately 1200 GtCO₂ / Wm⁻².





- 430-480 ppm
- 480-530 ppm
- 530-580 ppm

580-650 ppm 650-720 ppm



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Model Scenario Experiments

Since this fraction either increases or remains approximately constant in the future for most scenarios, our predictions tend to overestimate the budget for these scenarios as they do not account for the extra warming taken up by non-CO $_2$ forcers:

 $\int E \, \mathrm{d}t \leq E_0 \, \rho$





- 430-480 ppm
- 480-530 ppm
- 530-580 ppm

580-650 ppm 650-720 ppm



Model Scenario Experiments

- Here is the CO₂ forcing fraction for the circled scenario.
- The predicted budget for this scenario is almost exactly equal to the actual number since this fraction is very nearly constant after 2020.







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Model Scenario Experiments

If we incorporate all anthropogenic climate forcers using CO_2 -fe emissions, which behave exactly like CO₂ emissions, we find the prediction matches the actual budgets almost exactly.

 $\int E^{\rm fe}\,\mathrm{d}t\approx E_0^{\rm fe}\,\rho$





- 430-480 ppm
- 480-530 ppm
- 530-580 ppm

- 580-650 ppm
- 650-720 ppm
- Response to FAIR SCM



Model Scenario Experiments

Similarly, if we use CO₂-induced warming rather than total warming, our prediction is close to the 1:1 line.

 $\int E \, \mathrm{d}t \approx E_0 \cdot \rho^{CO_2}$







• In $(\Delta T, T'_0)$ space constant values of ρ lie on straight lines, from our mitigation timescale equation:

 $\rho = \Delta T/T_0'$

- We prefer this presentation of the budget as a timescale since it removes the uncertainties associated with the present-day annual emission rate.
- The associated exponential emissions reduction rates are also shown here.







- We compute anthropogenic and natural temperature responses to 200 best-estimate forcing ensemble members for 20 different model parameterisations.
- We then perform OLS regressions against 100 HadCRUT4 temperature observations and 50 CMIP5 PiC internal variability members to derive the anthropogenic signal.
- This results in the 20 million member ensemble of remaining warming to 1.5°C and corresponding present-day rate* shown here.

*taken as the trend of the best fit line of warming over 2013-2017







• The median of the distribution lies at a mitigation timescale of 22.5 years.







• A reduction rate of 5.5% per year would likely avoid 1.5°C.







• It is likely that we will cross 1.5°C somewhere between 2030 and 2050 if the current warming rate continues.







- Here we show results that use Berkeley Earth rather than HadCRUT4 temperature observations in the computation of the ensemble.
- This illustrates the importance of the current level of warming, a factor that has caused some of the discrepancies between previous estimates of the budget.
- However, even if the levels were identical, the higher warming rate in Berkeley Earth still reduces the mitigation timescale by roughly 3 years (see here for the impact of defining warming relative to 0.87°C in 2006-2015 as in the Paris Agreement Structured Expert Dialogue).







Comparison

- Richardson et al. (2018) used CMIP5 data, blended and interpolated identically to the observational datasets to match their coverage, to compute observationally consistent budgets to likely remain below 1.5°C, finding:
- 800/488 GtCO2 [HadCRUT4 / Berkeley]
- Combining our estimates of the timescale with emissions from the Global Carbon Project, we find:

744 / 407 GtCO2
 [HadCRUT4 / Berkeley]







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Summary and key messages

- With a simple, transparent and model independent method, we can calculate:
 - $_{\odot}$ the CO_2-fe budget with present day emissions and total warming
 - $\,\circ\,$ the CO_2 budget with present day emissions and CO_2-induced warming
 - $\,\circ\,$ An upper bound on the CO_2 budget with present day emissions and total warming
- The current rate of warming is as important as the level, and reducing its uncertainty would help to constrain estimates of the budget
- This can explain where some previous estimates disagree even if the models used predict similar levels of warming at present day, if some models are warming faster than others, they will naturally have a reduced budget in comparison.





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Acknowledgements

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For more information

• Link to paper: rdcu.be/2Use

- Correspondences to nicholas.leach@stx.ox.ac.uk
- Find me on twitter @nickleach0
- Thank you!



See next slide for appendices...



Appendices











Derivations of the mitigation timescale Linear warming rate reduction from t_0 to t_1 :

$$T'(t) = T'(t_0) - T'(t_0) \cdot \frac{t - t_0}{t_1 - t_0}$$
$$T(t) = T_0 + T'(t_0) \cdot (t - t_0) - T'(t_0) \cdot \frac{(t - t_0)^2}{2(t_1 - t_0)}$$
$$T(t_1) - T(t_0) \equiv \Delta T = T'(t_0) \cdot \frac{t_1 - t_0}{2}$$
$$\rho = \frac{\Delta T}{T'(t_0)} = \frac{t_1 - t_0}{2}$$

 \Rightarrow mitigation timescale is half the time to peak warming.

General exponential stabilisation:

$$T'(t) = T'(t_0) \cdot e^{-\frac{(t-t_0)}{\lambda}}$$

$$T(t_1) - T(t_0) \equiv \Delta T = \int_{t_0}^{t_1} T'(t) dt = \lambda T'(t_0)$$

$$\frac{\Delta T}{T'(t_0)} = \rho = \lambda$$

 \Rightarrow mitigation timescale is the decay timescale, and the associated reduction rate is

 $(1 - e^{-\frac{1}{\rho}}) \times 100\%$ per year





Scenario Selection

 We preselected a subset of AR5 database scenarios based on consistency with present-day observations of CO₂ emissions and non-CO₂ forcing gradient.





