

The Accommodation Coefficient of Water Molecules on Ice and its Role for Cirrus Clouds

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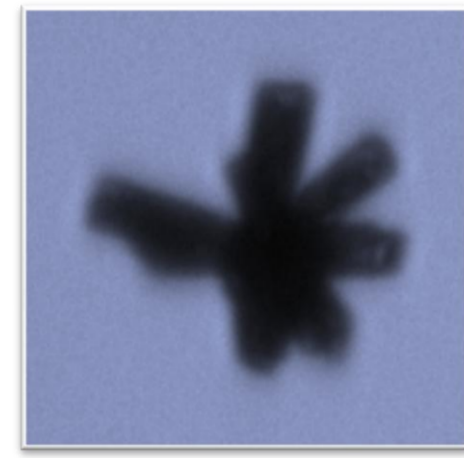
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I. MOTIVATION

One of the parameters governing the growth of ice crystals in cirrus clouds is the **accommodation coefficient of water molecules on ice** α . α describes the sticking probability of water molecules colliding with the surface of an ice particle. It is relevant for the ice growth in the kinetic regime, i.e. for sub-micron ice crystals.



However, the **magnitude of α is still unclear**. Experimental results vary between unity and values below 0.01 [Haynes *et al.*, 1992].

Model calculations suggest that **values for α between 0.1 and 1 do not have a significant impact on ice growth in cirrus clouds**. Lower values however could explain the observation of unexpectedly high ice number concentrations and supersaturations within cirrus clouds [Gierens *et al.*, 2003; Lohmann *et al.*, 2008].

II. THEORY OF ICE GROWTH

For **spherical ice crystals**, the mass increase per time can be described by the following formula [Pruppacher and Klett, 1997]:

$$\frac{dm}{dt} = \frac{4\pi r (S_{ice} - 1)}{\frac{RT}{e_{sat,i} D_v^* M_w} + LH}$$

where r is the ice particle radius, S_{ice} the saturation ratio, and $e_{sat,i}$ the saturation vapor pressure with respect to ice. R , T , and M_w are the gas constant, absolute temperature, and molar weight of water, respectively. LH describes the growth-impeding effect of the latent heat of deposition.

The **accommodation coefficient α** enters in the modified diffusivity of water vapor in air

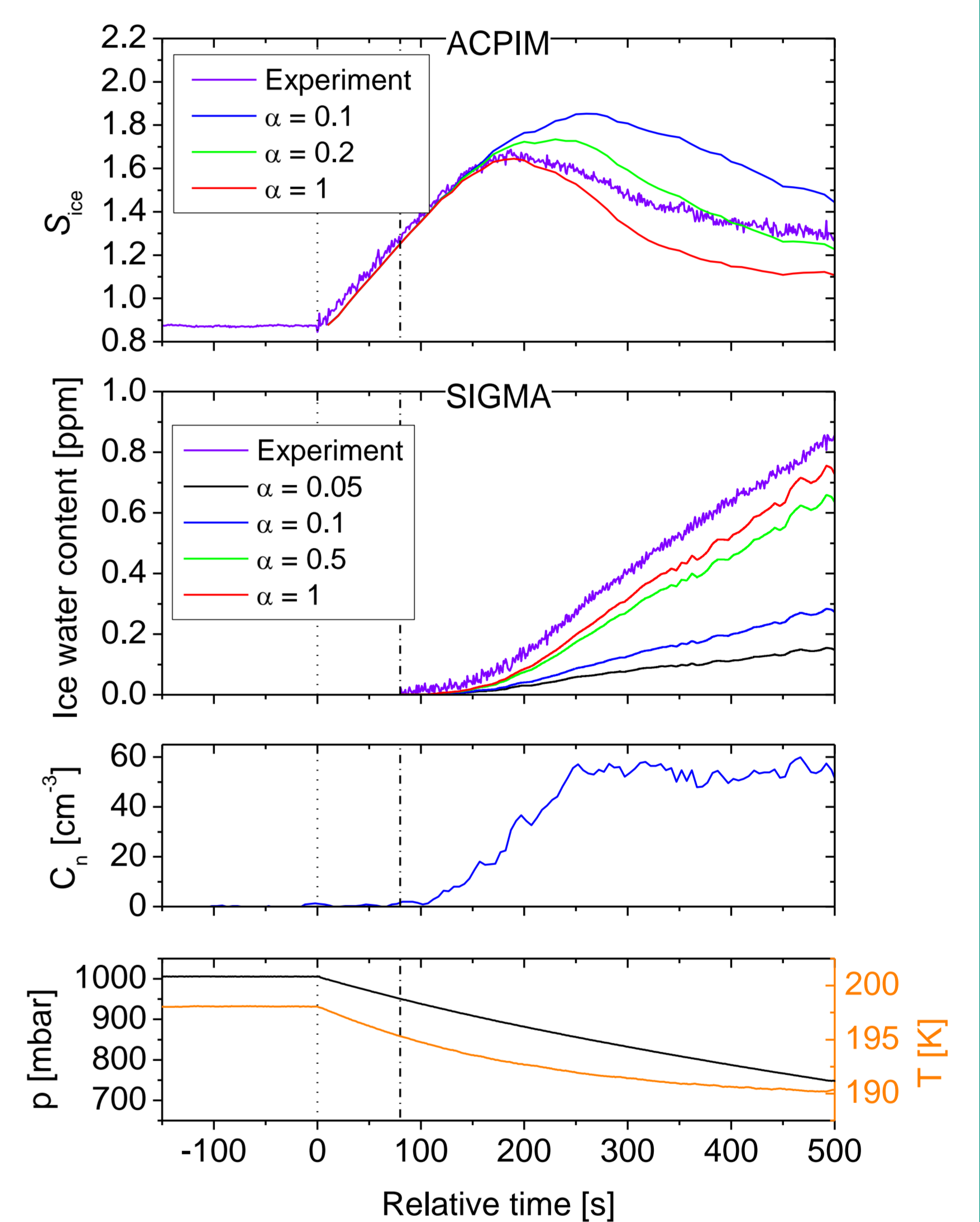
$$D_v^* = \frac{D_v}{\frac{r}{r+\Delta_v} + \frac{D_v}{r\alpha} \left(\frac{2\pi M_w}{RT}\right)^{1/2}}$$

where Δ_v is the vapor jump distance.

III. AIDA CIRRUS EXPERIMENTS

Dedicated experiments examining the ice crystal growth for **deposition nucleation** in the **temperature range from 190 K to 230 K** were carried out at the cloud simulation chamber AIDA [Möhler *et al.*, 2003].

As aerosols, **hematite** particles and **graphite-spark generator (GSG) soot** were used.



Example experiment of the AIDA studies

- Dotted line: start of the experiment, i.e. start of pumping
- Dashed-dotted line: ice onset

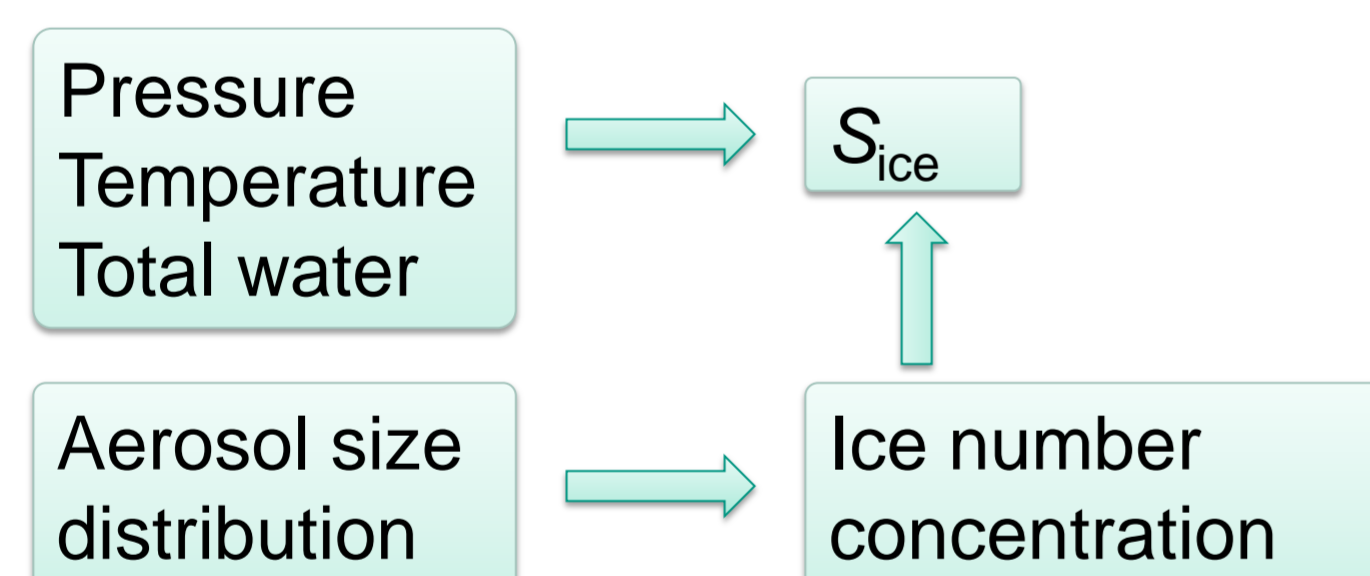
Water vapor and total water are measured by two **tunable diode laser (TDL) hygrometers** [Fahey *et al.*, 2009]. From the difference of these two measurements, the ice water content is derived.

The ice number concentration C_n is measured by an **optical particle counter (WELAS)**.

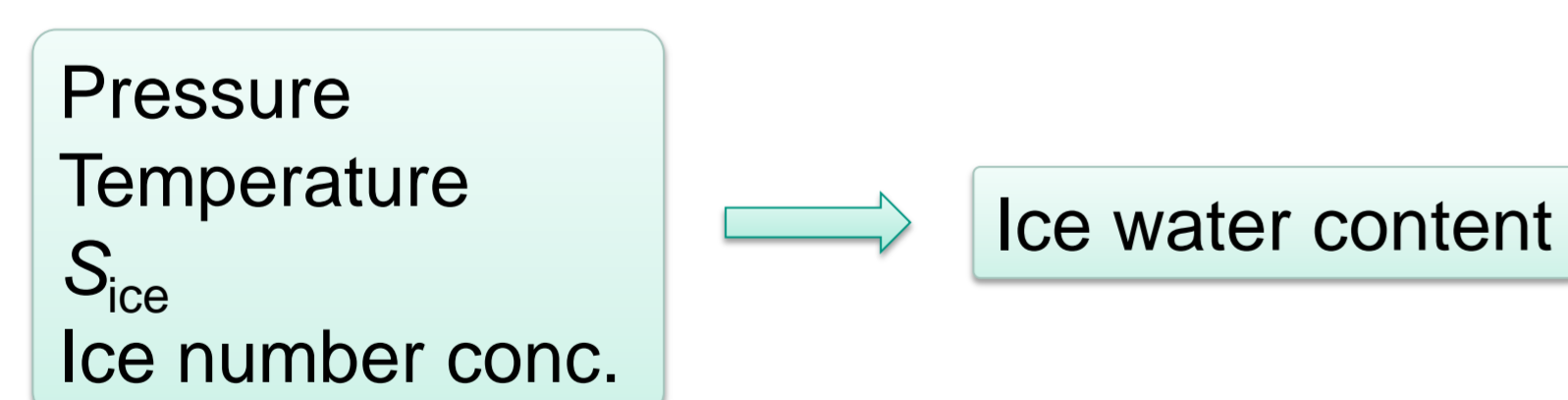
IV. MODELING

Two models are used to derive the accommodation coefficient α from experimental data:

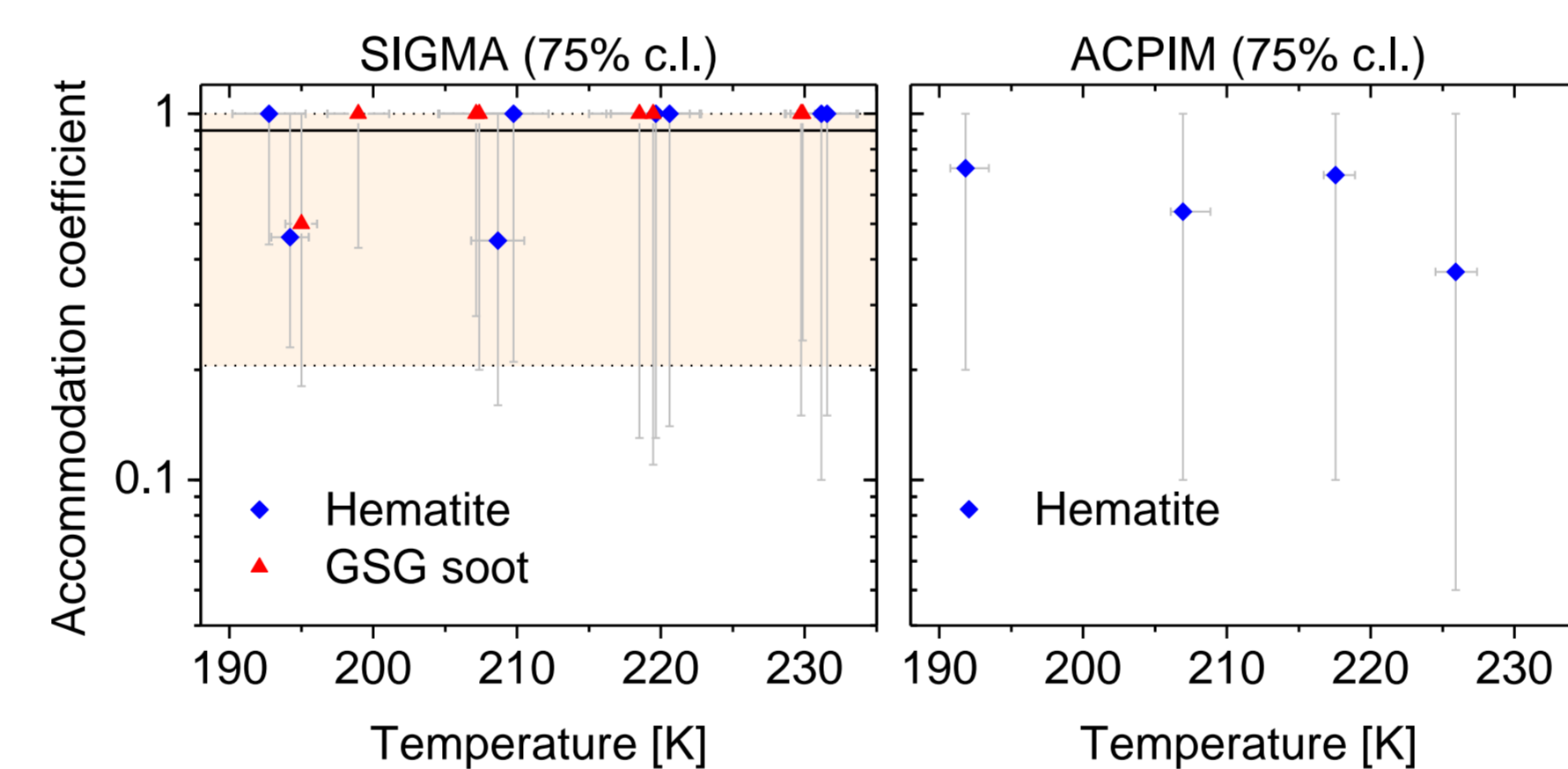
- The **Aerosol-Cloud-Precipitation Interaction Model (ACIPIM)** [Connolly *et al.*, 2009]



- The **Simple Ice Growth Model for determining Alpha (SIGMA)**



VI. RESULTS



Temperature dependent values of α for a set of 16 individual AIDA experiments. The error bars are obtained from the MC uncertainty estimate described previously.

- Both models in **good agreement** with each other with all best-fit points in the range 0.1-1
- **Lower values than 0.1 excluded** by the error estimate, independent of the aerosol type (SIGMA)
- **Temperature averaged value $\alpha = 0.9_{-0.7}^{+0.1}$** (SIGMA)

V. UNCERTAINTY ESTIMATE

Accuracies of the experimental data sets used for a **Monte Carlo (MC)** uncertainty analysis:

	Accuracy	ACIPIM	SIGMA
S_{ice}	5%	•	•
Total water	5%	•	•
Ice number concentration	20%	•	•
Aerosol number concentration	20%	•	
Aerosol size distribution	10%	•	
Temperature	0.3 K	•	

VII. CONCLUSIONS

- The ACIPIM and the SIGMA model are in **good agreement** despite their different approaches in determining α
- In the temperature range from 190 K to 230 K, **values between 0.1 and 1 for α** are preferred by both models
- The **uncertainty analysis** excludes α -values below 0.1 with a **temperature averaged value $\alpha = 0.9_{-0.7}^{+0.1}$** (SIGMA)
- These results suggest that **α does not have a significant impact on ice growth** in cirrus clouds

ACKNOWLEDGEMENTS

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REFERENCES

- D. R. Haynes *et al.*, J. Phys. Chem. **96**, 8502 (1992)
 K. M. Gierens *et al.*, J. Geophys. Res., **108**, D2, 4069 (2003)
 U. Lohmann *et al.*, Environ. Res. Lett. **3**, 045022 (2008)
 H. R. Pruppacher and J. D. Klett, 2nd ed., Kluwer Acad. (1997)
 O. Möhler *et al.*, Atmos. Chem. Phys. **3**, 211 (2003)
 D. W. Fahey *et al.*, AquaVIT White Paper (2009), available at <https://aquavit.icg.kfa-juelich.de/AquaVit/AquaVitWiki>
 P. J. Connolly *et al.*, Atmos. Chem. Phys. **9**, 2805 (2009)