

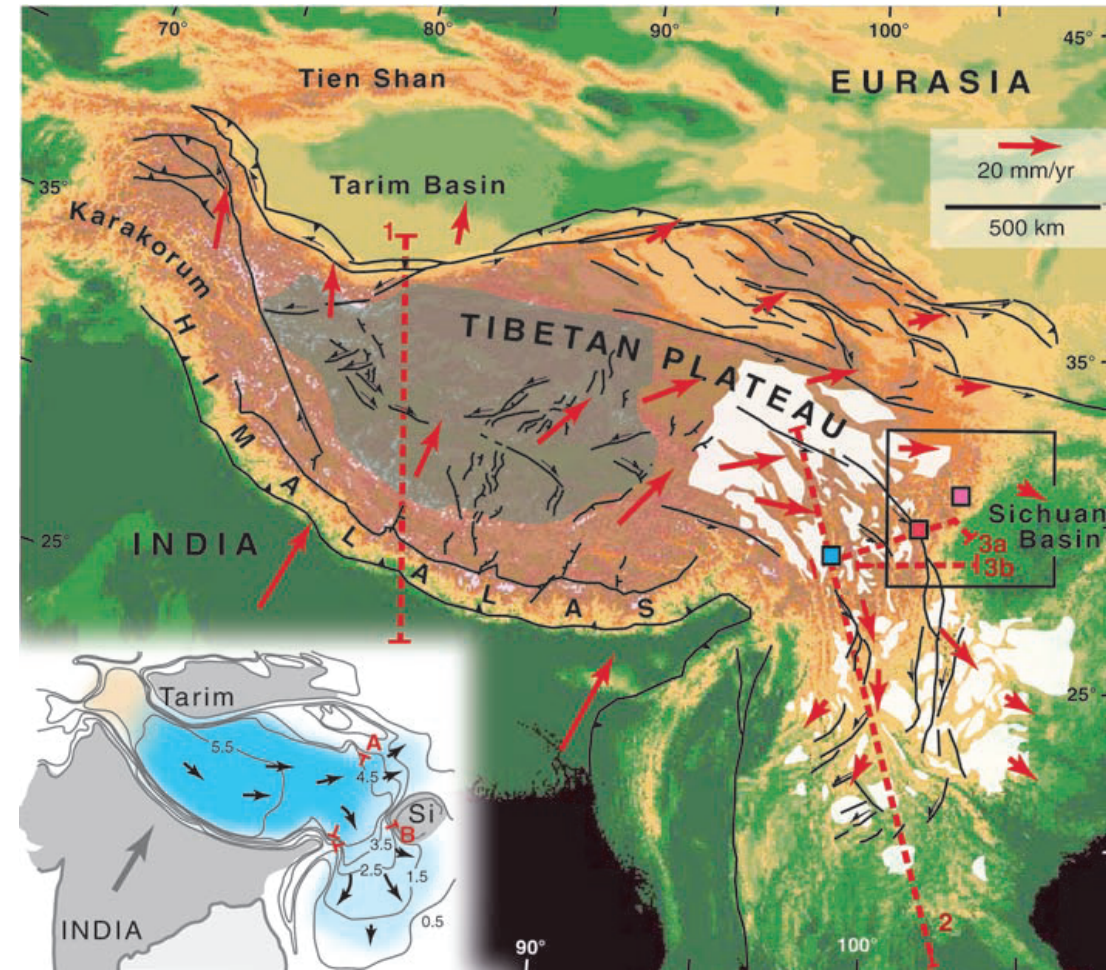
The development of topographic plateaus in an India-Asia-like collision zone using 3D numerical simulations

Adina E. Püsoök, Boris J.P. Kaus, Anton Popov

puesoek@uni-mainz.de

INTRODUCTION

- The India-Asia collision zone formed around 50 million years ago, when the Indian continent collided with Eurasia.
- Dynamics not yet understood: the rise of the abnormally thick Tibetan plateau, the deformation at its Eastern and Western syntaxes, the transition from subduction to collision and uplift and the interaction of tectonics and climate.



"Double thickness" problem:

- Wholesale underthrusting (Argand model)
- Distributed homogeneous shortening (England and Houseman, 1986);
- Slip-line field model (Tapponier and Molnar, 1976);
- Lower crustal flow model (Royden et al., 2008, Beaumont et al., 2004).

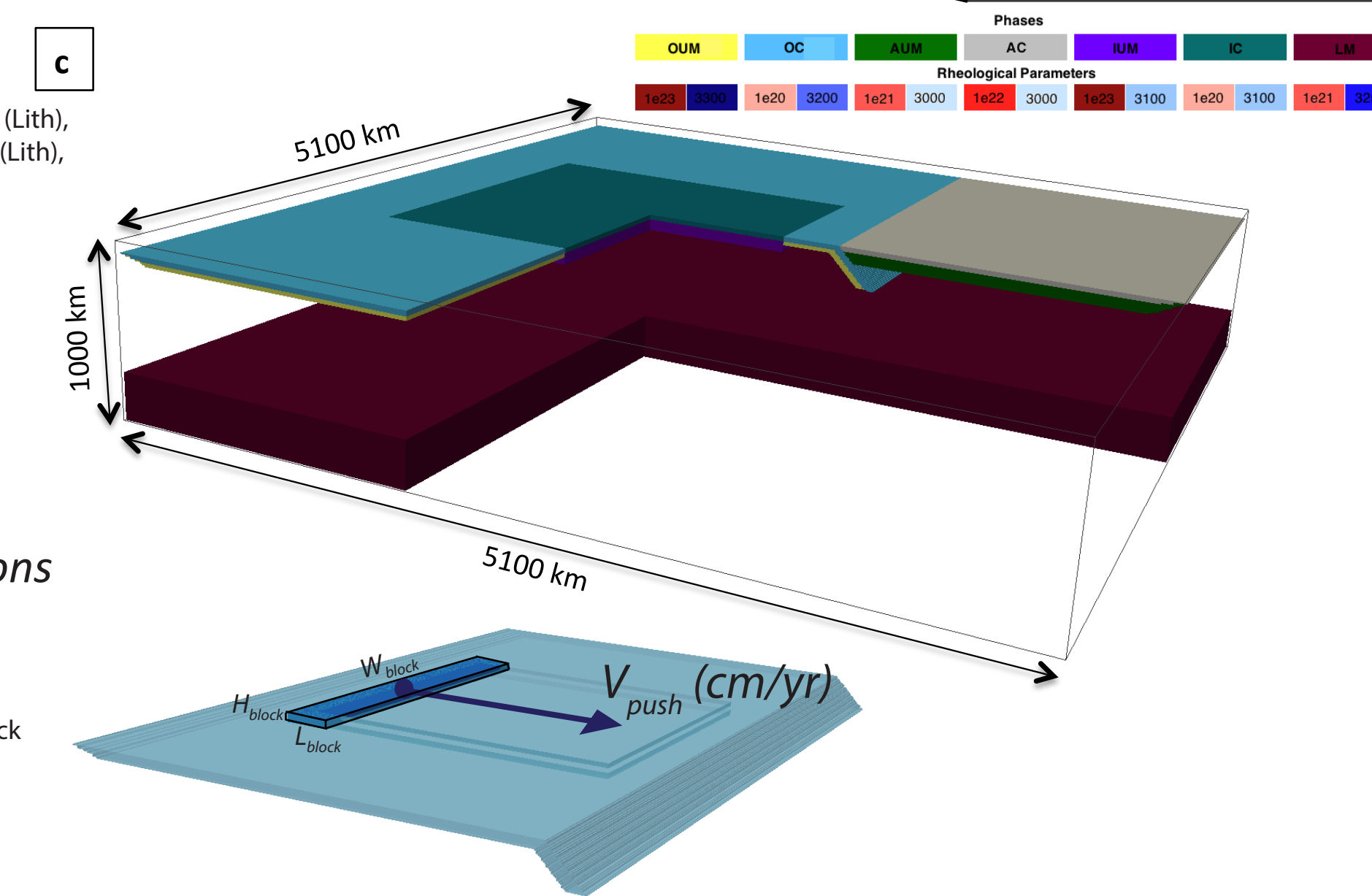
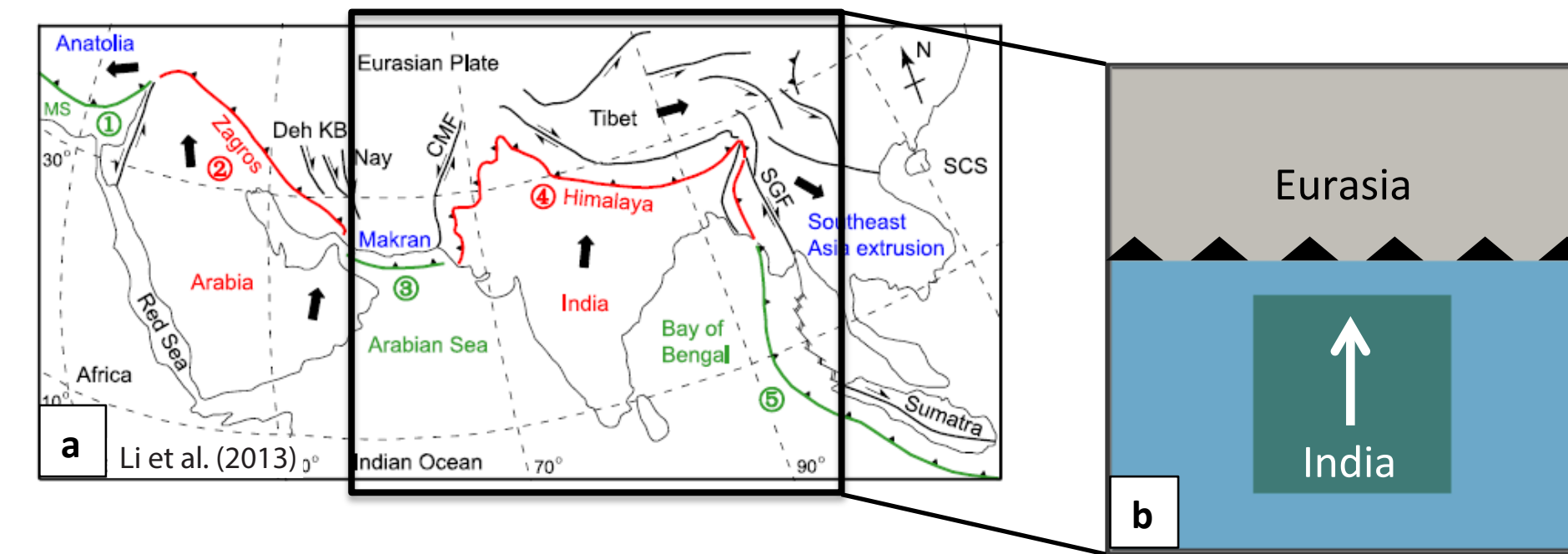
Fig. 1. From Royden et al. (2008). Topography shade map for the Tibetan region. Red arrows indicate generalized GPS measurements.

METHOD

- We use the parallel 3D code LaMEM (Lithosphere and Mantle Evolution Model), with a finite difference staggered grid (FDSTAG) solver, viscous rheology and an internal free surface, which allows for the development of topography
- We simulate both free subduction-collision models, and with imposed velocity boundary conditions

Model Setup

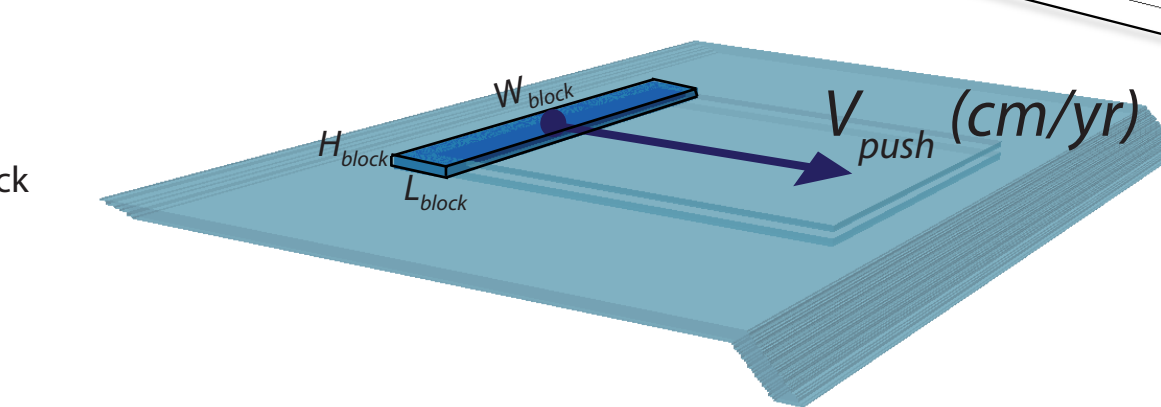
Fig. 2. a-b) Initial model setup was built using a simplified tectonic map of the region. c) 3D perspective of the model setup. Domain: 5100 km x 5100 km x 1000 km. Computational resolution: 170 x 170 x 33 grid cells (3rd particles/grid cell). Boundary conditions: Free slip on all boundaries and Sticky Air layer on top. Stokes Solver: Fully Coupled Velocity Solver: HyPre + Multigrid (AML) Phases: OUM - Oceanic Upper Mantle (Lith), OC - Oceanic Crust, AUM - Asian Upper Mantle (Lith), AC - Asian Crust, IUM - Indentor Upper Mantle (Lith), IC - Indentor Crust, LM - Lower Mantle. The rheological parameters used for the reference model are displayed in the colorbar. Reference values for the mantle: $\eta = 1e20$ Pa.s, $\rho = 3200$ kg/m³.



External forcing

Pushing is imposed as Dirichlet boundary conditions

Fig. 3. Schematic of how the pushing BCs are incorporated into the model. The pushing block is defined by its center coordinates: $(x_{block}, y_{block}, z_{block})$ and by its dimensions: $L_{block}, W_{block}, H_{block}$. The block moves with constant velocity V_{push} .



ASSUMPTIONS and APPROACH

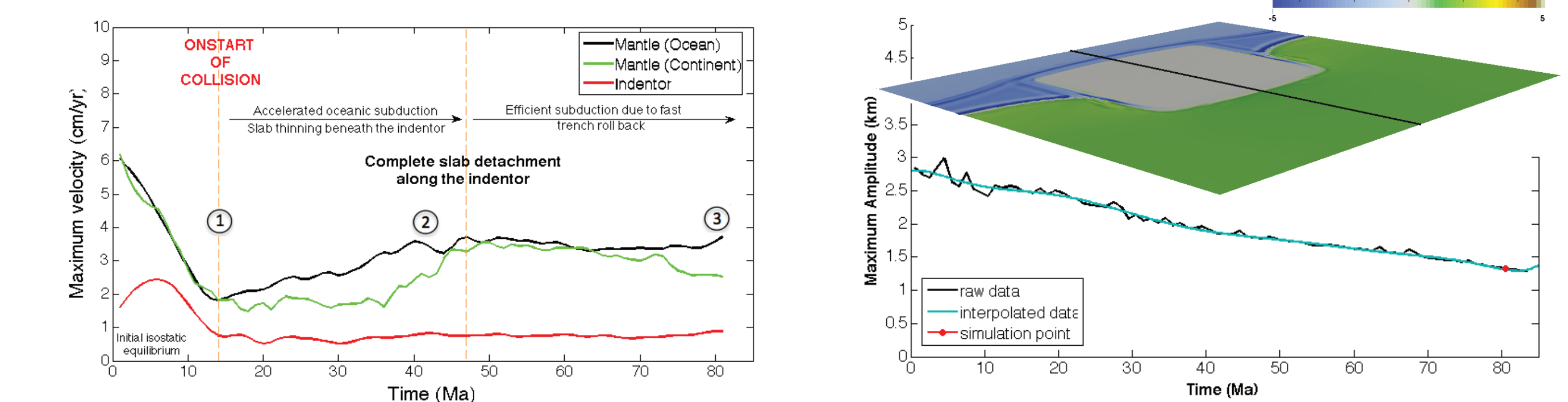
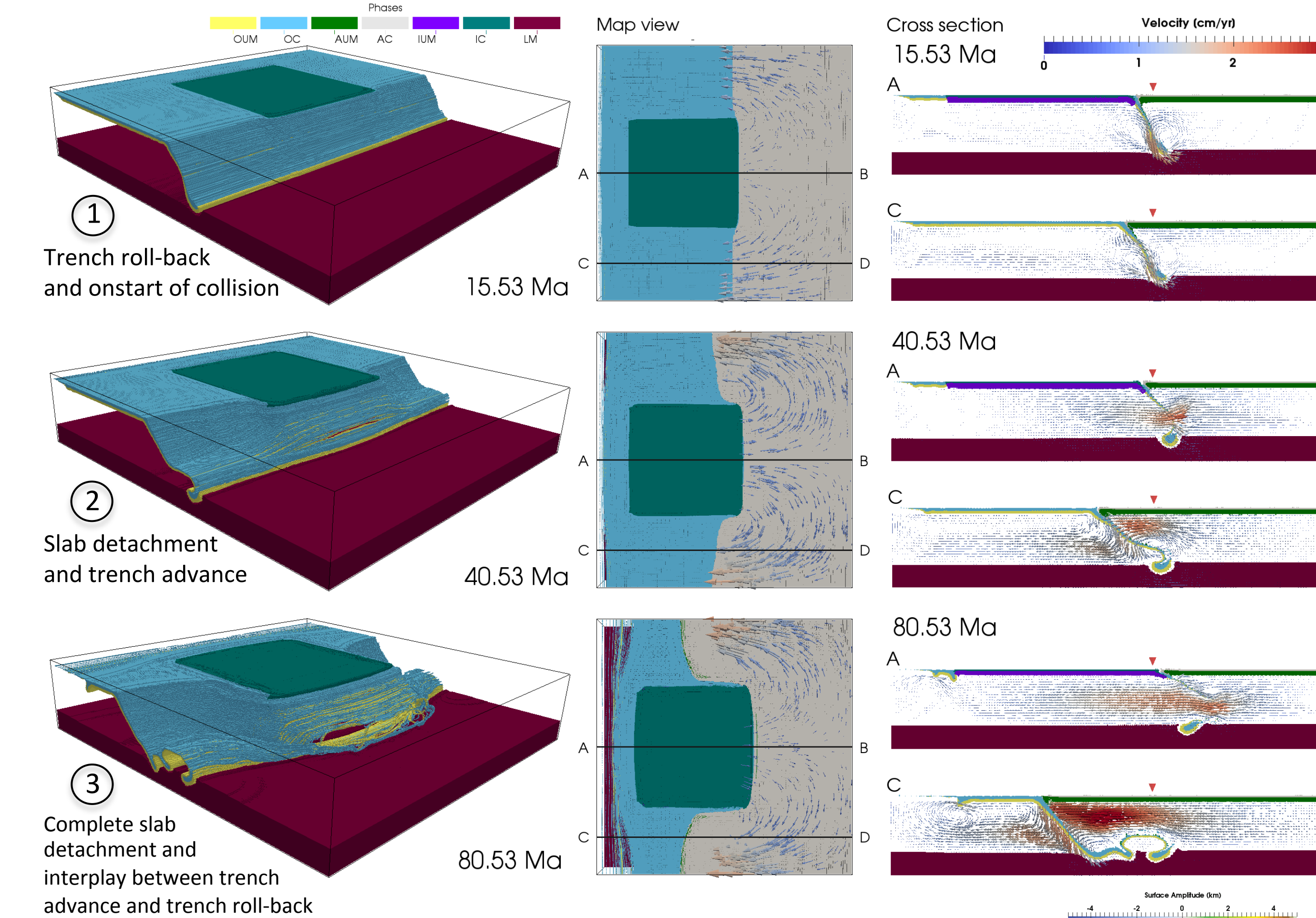
- The model has 16 free parameters, and to test the basic response of the system to these parameters, it would require 2¹⁶ = 65536 runs.
- We assume that the properties of the lithosphere are the focus of this study, and choose reference values for the mantle.
- We use weak/strong or low/high values for the parameters, because we are interested in understanding the importance of the varied parameters to the system, rather than focus on their true value. See table below for more reference.
- Quinteros et al. (2010) showed that a viscosity contrast of 10 and density contrast of 50 kg/m³ at the transition zone between the upper and lower mantle gives the best results of slab behaviour in numerical models compared to tomography observations.
- A viscosity contrast between slab and mantle less than 10³ Pa.s would result into an unstable, fragmented slab, contrary to what is observed in tomographic studies.

Factor	Case 1	Case 2	Case 3	Case 4
No. of runs	16	48	72	68
Parameters	η_{AUM} 1e21-1e22 Pa.s η_{AC} 1e21-1e22 Pa.s ρ_{OC} 3100-3200 kg/m ³ ρ_{AUM} 3000-3100 kg/m ³	η_{IUM} 1e22-1e23 Pa.s η_{IC} 1e20-1e21 Pa.s ρ_{AUM} 3100-3200 kg/m ³ ρ_{IC} 3000-3100 kg/m ³	η_{AUM} 1e21-1e23 Pa.s η_{AC} 1e21-1e23 Pa.s ρ_{AUM} 3000-3100 kg/m ³ ρ_{AC} 3000-3100 kg/m ³	pushing BC strong blocks pushing BC + strong blocks

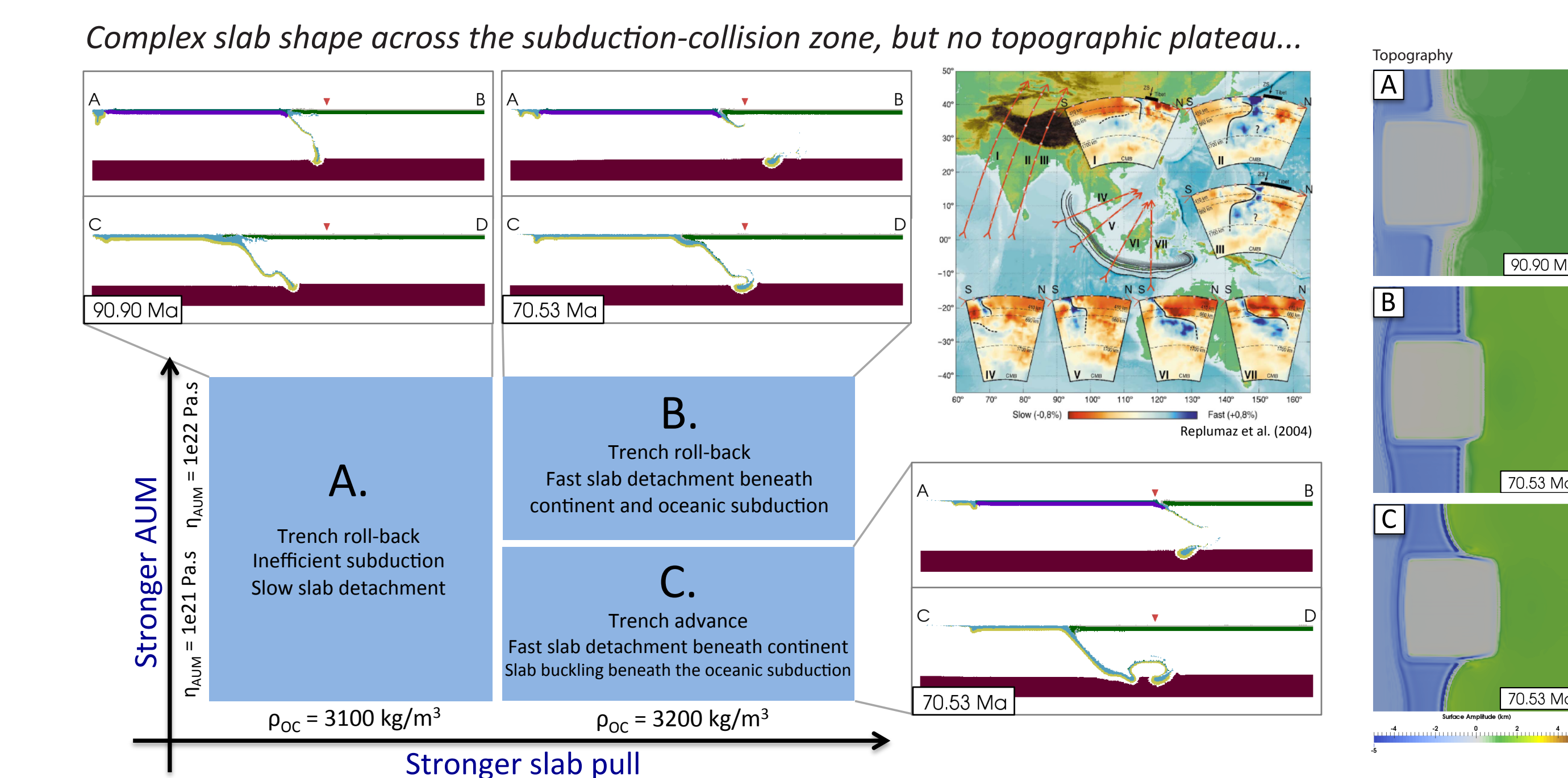
Table 1. Summary of cases and parameters varied.

RESULTS

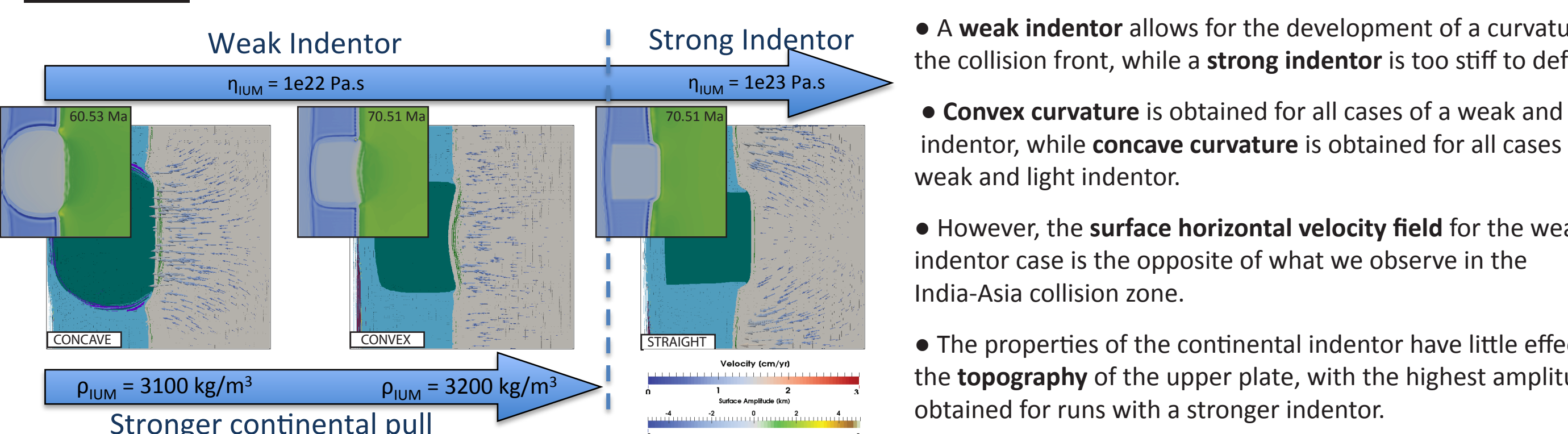
Reference Model - evolution and general aspects



Case 1 - Slab pull and dynamics of the slab

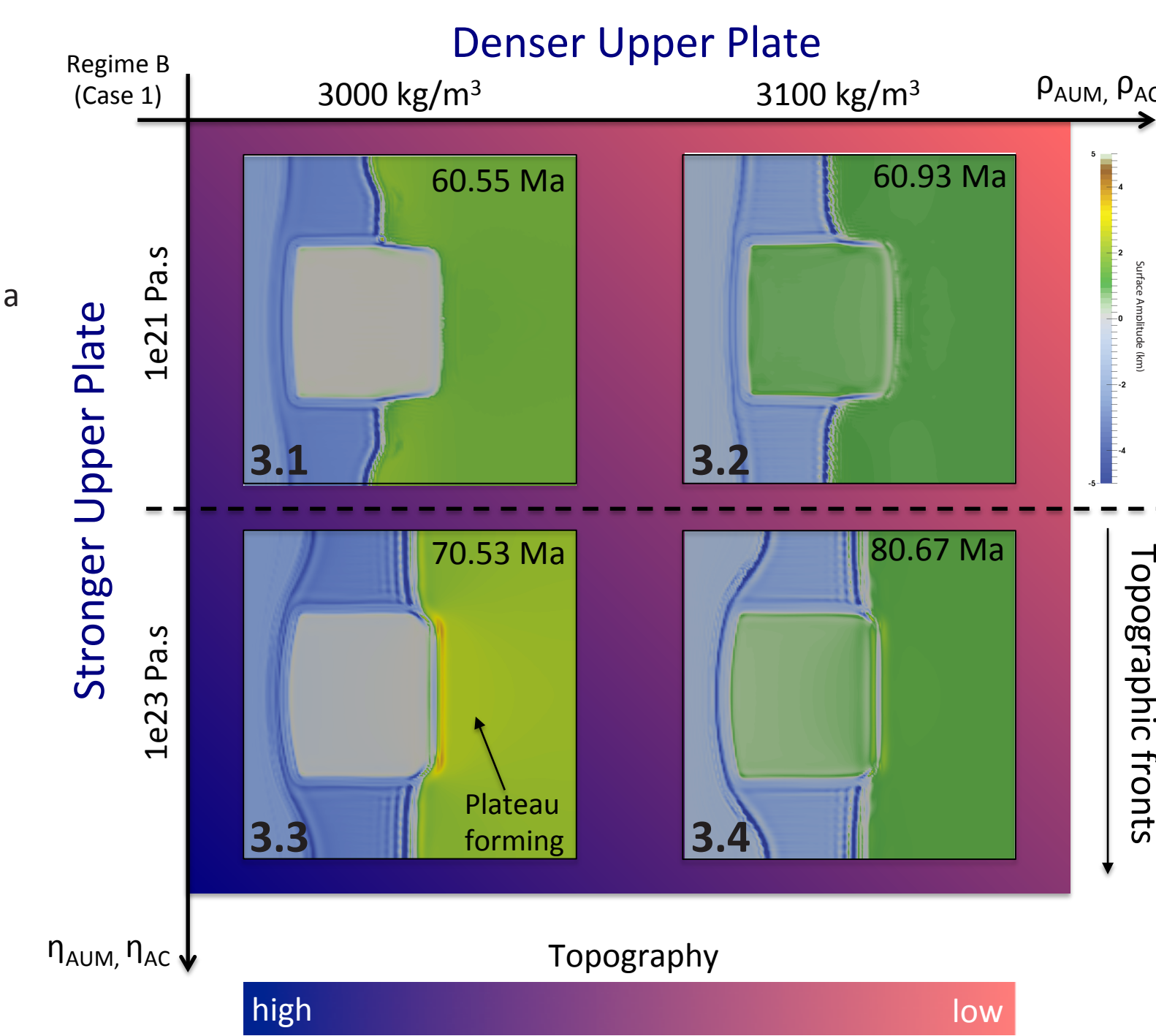
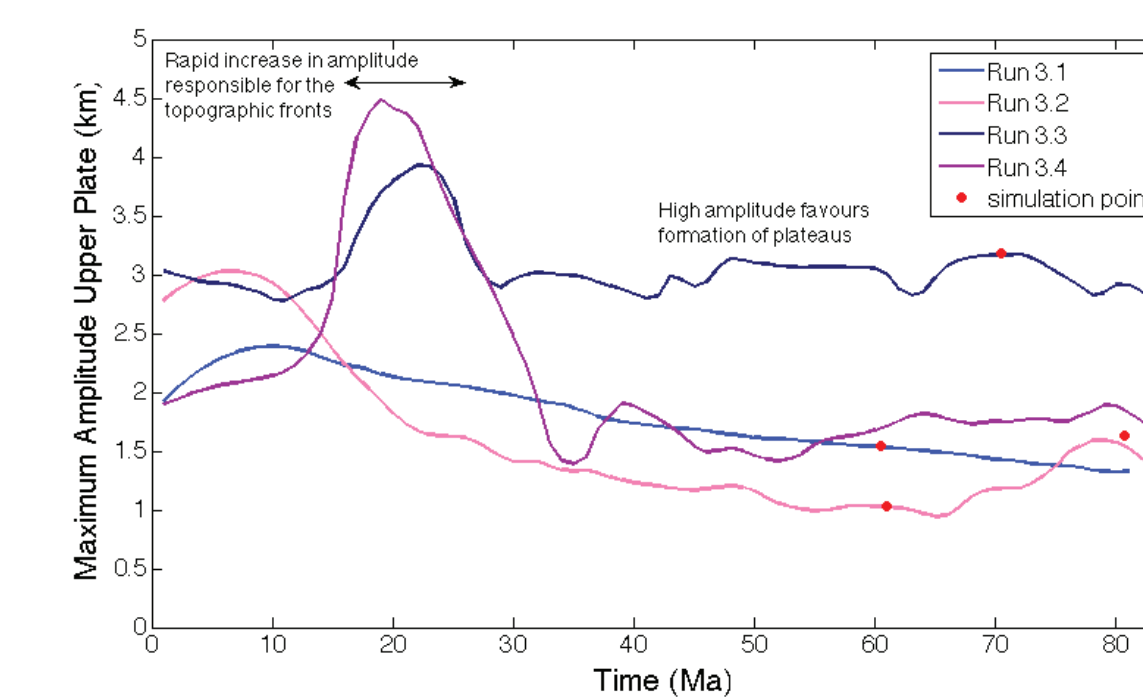


Case 2 - Deformable continental indentor and curvature



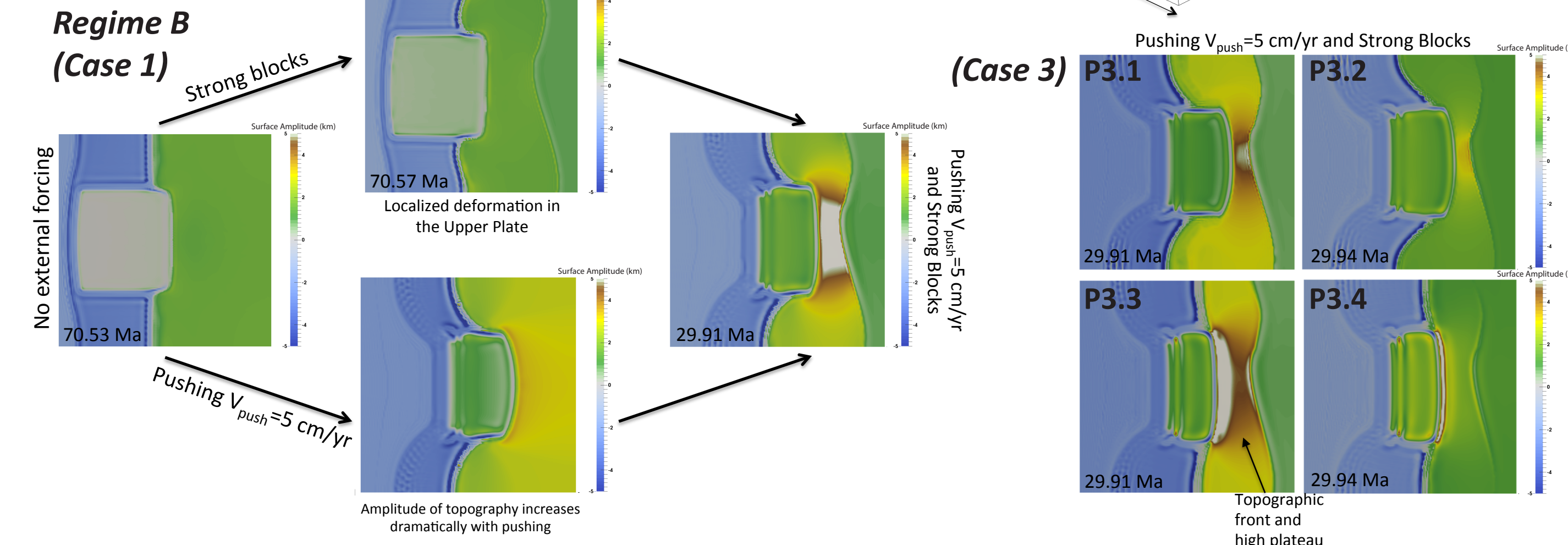
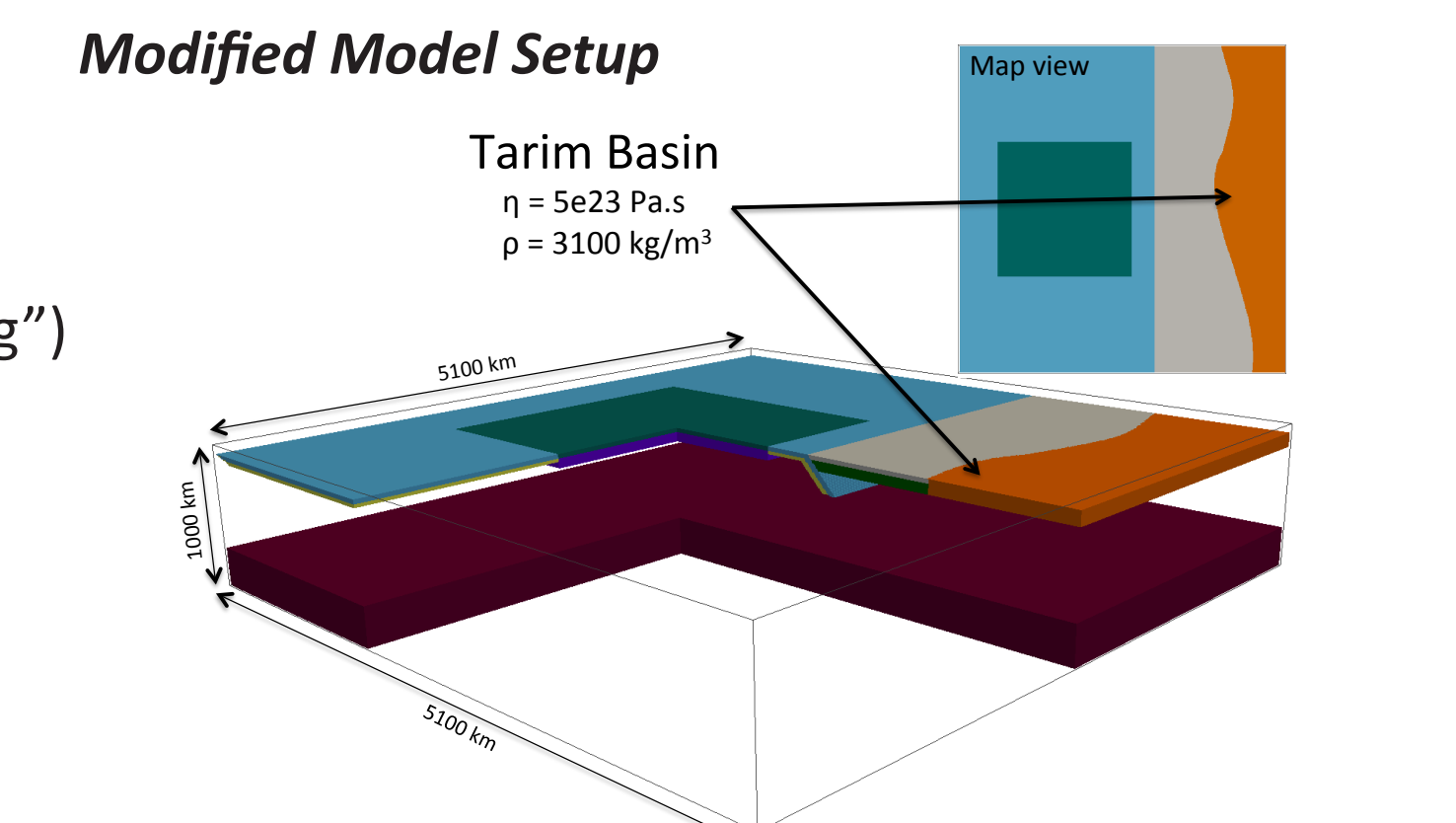
Case 3 - Upper plate and development of topographic plateaus

- The way slab dynamics is affected by upper plate parameters is shown in Case 1.
- Topography is strongly influenced by these parameters.
- A topographic front forms when the upper plate is strong, as a result of rapid accumulation of gravitational potential energy (GPE) locally.
- A high topographic plateau forms when the upper plate is strong and light, able to store GPE for a longer time.



Case 4 - External forcing

- Ridge Push - simulate far field forces (referred here as "pushing") (Chemenda et al., 2000, Li et al., 2008)
- Presence of strong blocks - such as the Tarim Basin N of Tibetan Plateau (Neil and Houseman, 1997)



CONCLUSIONS

- More than 200 numerical simulations and factors such as slab pull, deformable continental indentor, upper plate and external forcing have been performed here.
- Dynamics of subduction-collision 3D models is very complex with distinct behaviours beneath the continental collision and oceanic subduction at the sides.
- External pushing and the presence of strong blocks such as the Tarim Basin are necessary to create both high topographic fronts (Himalayas) and plateaus (Tibetan Plateau). Upper plate material properties also give a signature to the topographic amplitude.

REFERENCES

Ali, J.R., Aitchison, J.C., 2005. Greater India. *Earth-Science Reviews* 72, 169-188.
 Beaumont, C., Jamieson, R.A., Nguyen, M.H., Medvedev, S.E., 2004. Crustal channel flows: 1. Numerical models with applications to the tectonics of the Himalayan-Tibetan orogeny. *Journal of Geophysical Research* 109, B06406.
 Chemenda A., Burg J-P., Mattauer M., 2000. Evolutionary model of the Himalaya-Tibet system: geopoem based on new modelling, geological and geophysical data. *Earth and Planetary Science Letters*, 174:297-409.
 England, P., Houseman, G., 1986. Finite strain calculations of continental deformation. 2. Comparison with the India-Asia collision zone. *Journal of Geophysical Research - Solid Earth and Planets* 91 (B3), 3664-3676.
 Li, C., van der Hilst, R.D., Meltzer A.S., Engdahl, E.R., 2008. Subduction of the Indian lithosphere beneath the Tibetan Plateau and Burma. *Earth and Planetary Science Letters*, 274 (1-2):157-168.
 Li, Z.H., Xu, Z., Gerya, T., Burg, J.P., 2013. Collision of continental corner from 3-D numerical modeling. *Earth and Planetary Science Letters*, 380(C):98-111.
 Neil, E.A., Houseman, G.A., 1997. Geodynamics of the Tarim basin and the Tian Shan in central Asia. *Tectonics*, 16:571-584.
 Quinteros, J., Sobolev, S.V., Popov, A.A., 2010. Viscosity in transition zone and lower mantle: Implications for slab penetration. *Geophysical Research Letters*, 37(9).
 Replumaz, A., Karason, H., van der Hilst, R.D., Besse, J., Tapponnier, P., 2004. 4-D evolution of SE Asia's mantle from geological reconstructions and seismic tomography. *Earth and Planetary Science Letters*, 221(1-4):103-115.
 Royden, L.H., Burchfiel, B.C., van der Hilst, R.D., 2008. The geological evolution of the Tibetan Plateau. *Science* 321, 1054-1058.
 Tapponier, P., Molnar, P., 1976. Slip-line field-theory and large-scale continental tectonics. *Nature* 264 (5584), 319-324.

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