

# Greenhouse gas natural trends : what do we learn from ice core analyses ?



*EPICA Dome Concordia drilling trench*

*© S. Drapeau, IPEV*

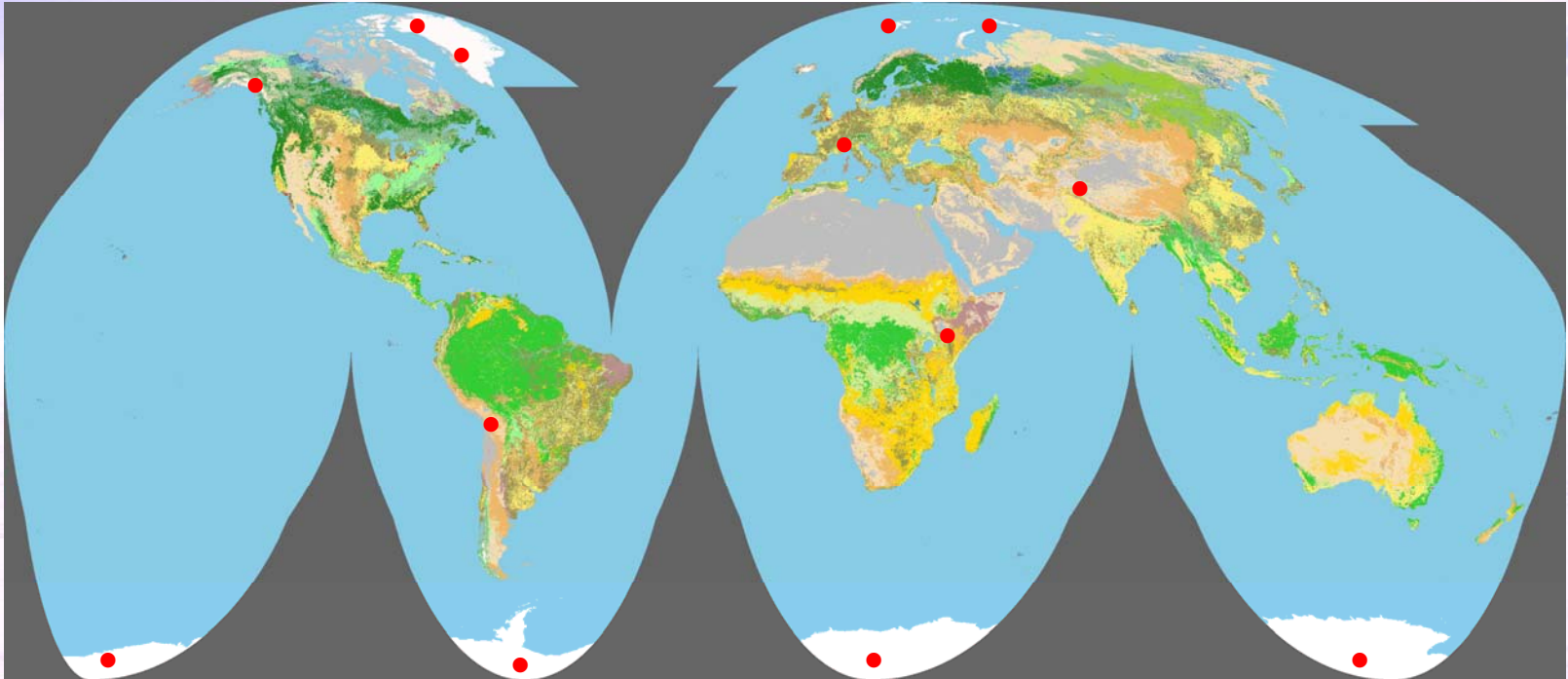
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# Outline

- Introduction to ice cores and gas trapping
- Why the future brings us to the past
- « Facts » on past greenhouse gas changes
- Where we stand regarding explanations of past greenhouse gas changes
- Where we want to go during IPY and beyond...

# Specificity of ice cores as archives



## Advantages

- Temporal resolution
- Range of parameters
- Regional to global significance
- No biology in the transfer function

## Disadvantages

- Only on poles or mountains
- A few 100,000 years coverage
- Physics and chemistry of transfer function not always straightforward

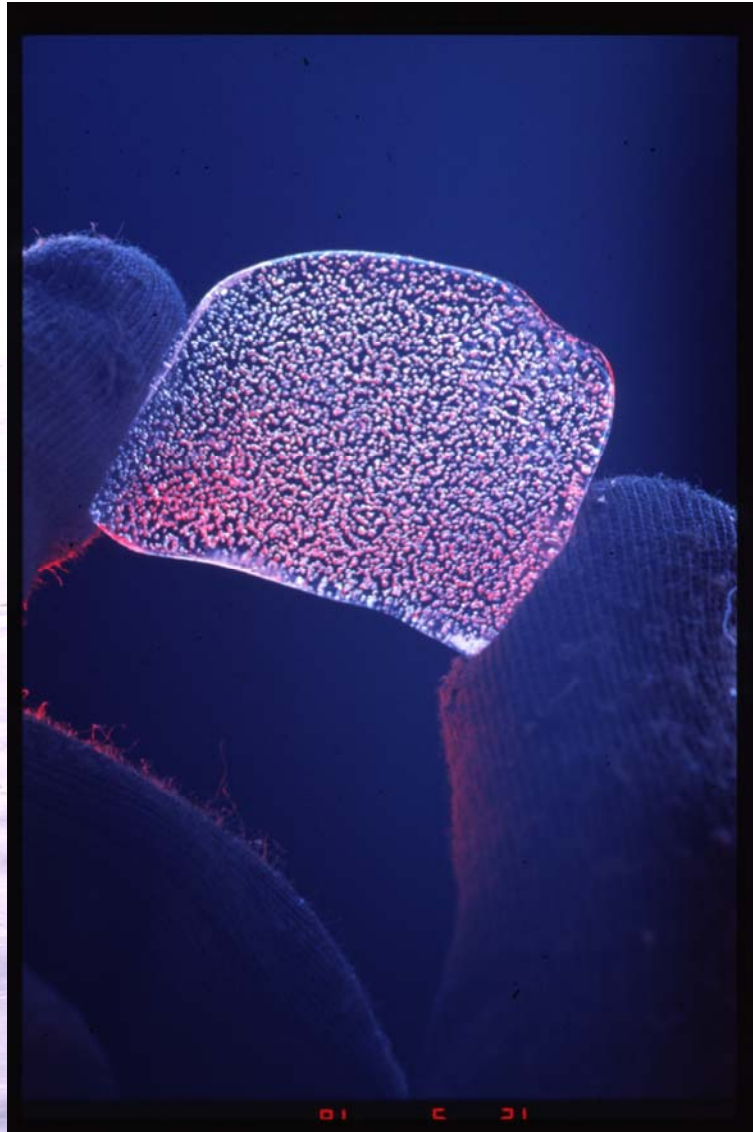
# Parameters accessible in natural ice



- Temperature (isotopes of  $\text{H}_2\text{O}$ )
- Accumulation (stratigraphy,  $^{10}\text{Be}$ )
- Aerosols of natural origin :
  - Marine ( $\text{Cl}^-$ ,  $\text{Na}^+$ ,  $\text{SO}_4^{=}$ , MSA,...)
  - Continental (dust,  $\text{Al}$ ,  $\text{Ca}^{++}$ ,  $\text{NH}_4^+$ , organic acids)
  - Volcanic ( $\text{Cl}^-$ ,  $\text{SO}_4^{=}$ )
  - Cosmic ( $\text{Ir}$ ,  $^{10}\text{Be}$ )
- Aerosols of anthropogenic origin :
  - $\text{SO}_4^{=}$ ,  $\text{NO}_3^-$ , heavy metals,  $^{137}\text{Cs}$ ,...
- Atmospheric composition ( $\text{N}_2$ ,  $\text{O}_2$ ,  $\text{CO}_2$ ,  $\text{CH}_4$ ,  $\text{N}_2\text{O}$ ,  $\text{CO}$ ,...)

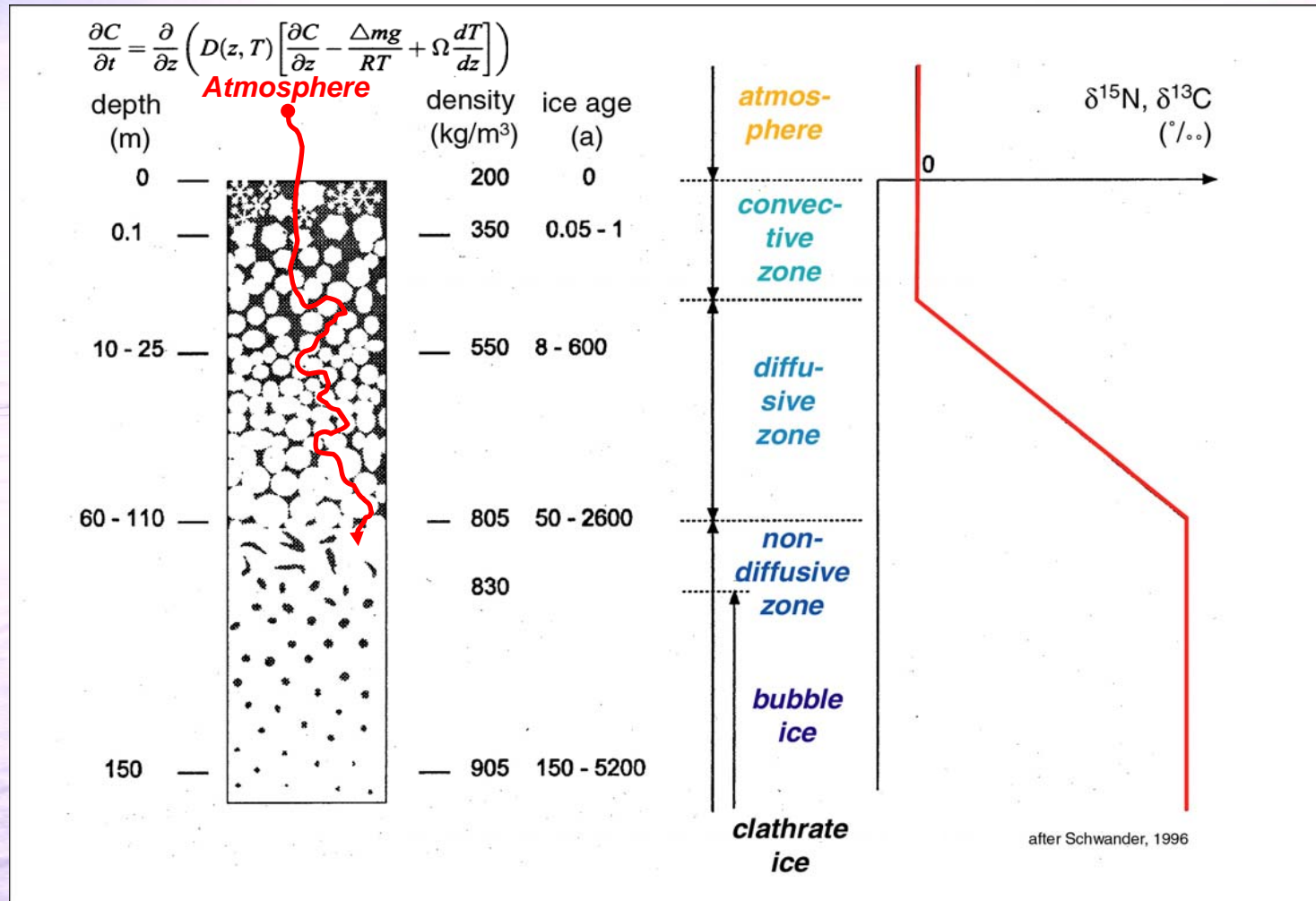
*Natural ice through polarized light  
(sample size : ~2 x 3,5 cm)  
© V. Lipenkov, LGGE-CNRS*

How does mother Nature put atmospheric samples in air bubbles trapped in natural ice ?



*Bubbles in polar ice*  
© D. Etheridge, CSIRO, Australia

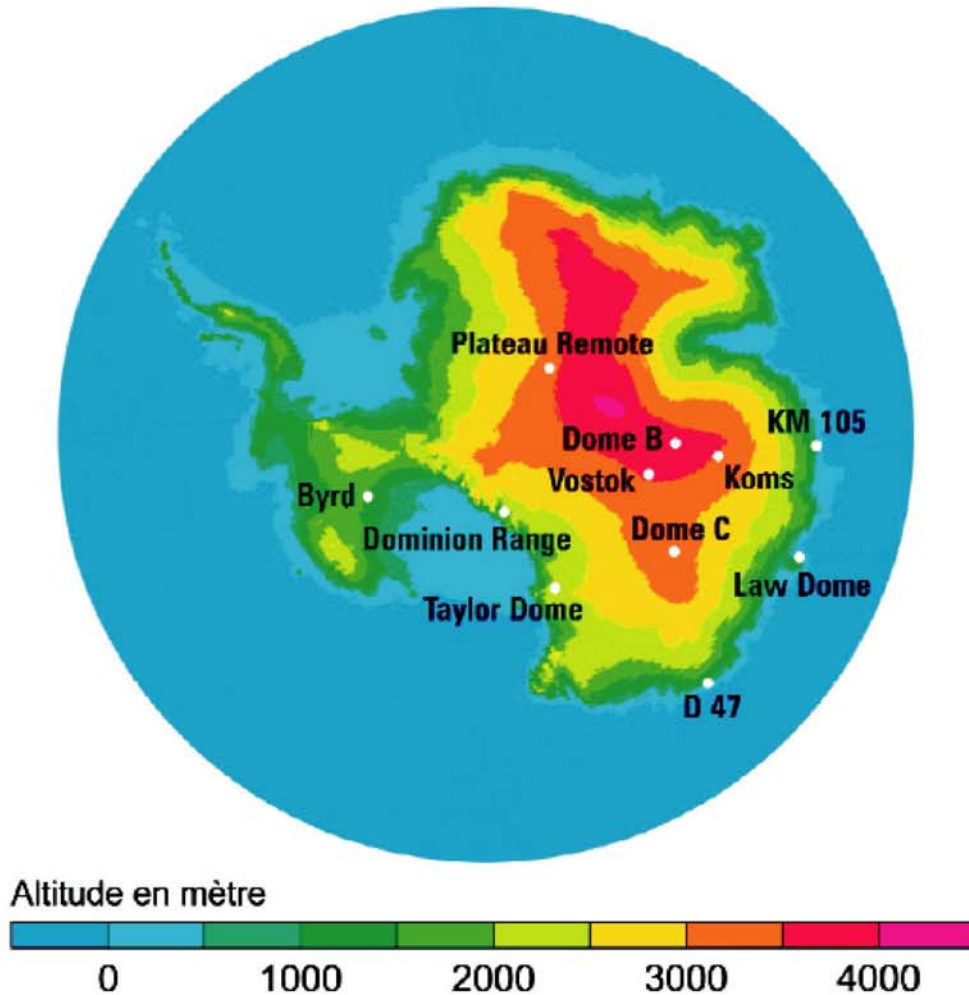
# Gas diffusion and trapping in polar firn



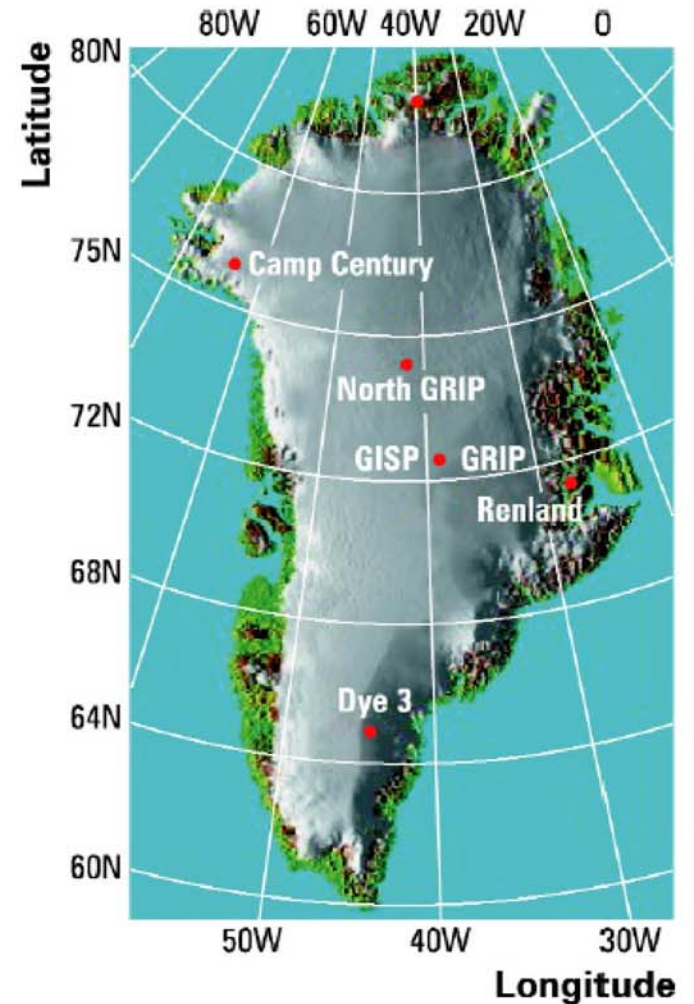
Adapted from Schwander, NATO ASI, 1996

The gas composition in polar firn and then in air bubbles results from molecular diffusion, gravitational settling and thermal diffusion

# Antarctica



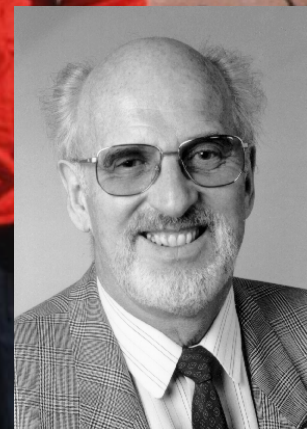
# Greenland



**1968-2006 : 38 years of international efforts to drill deep ice cores**

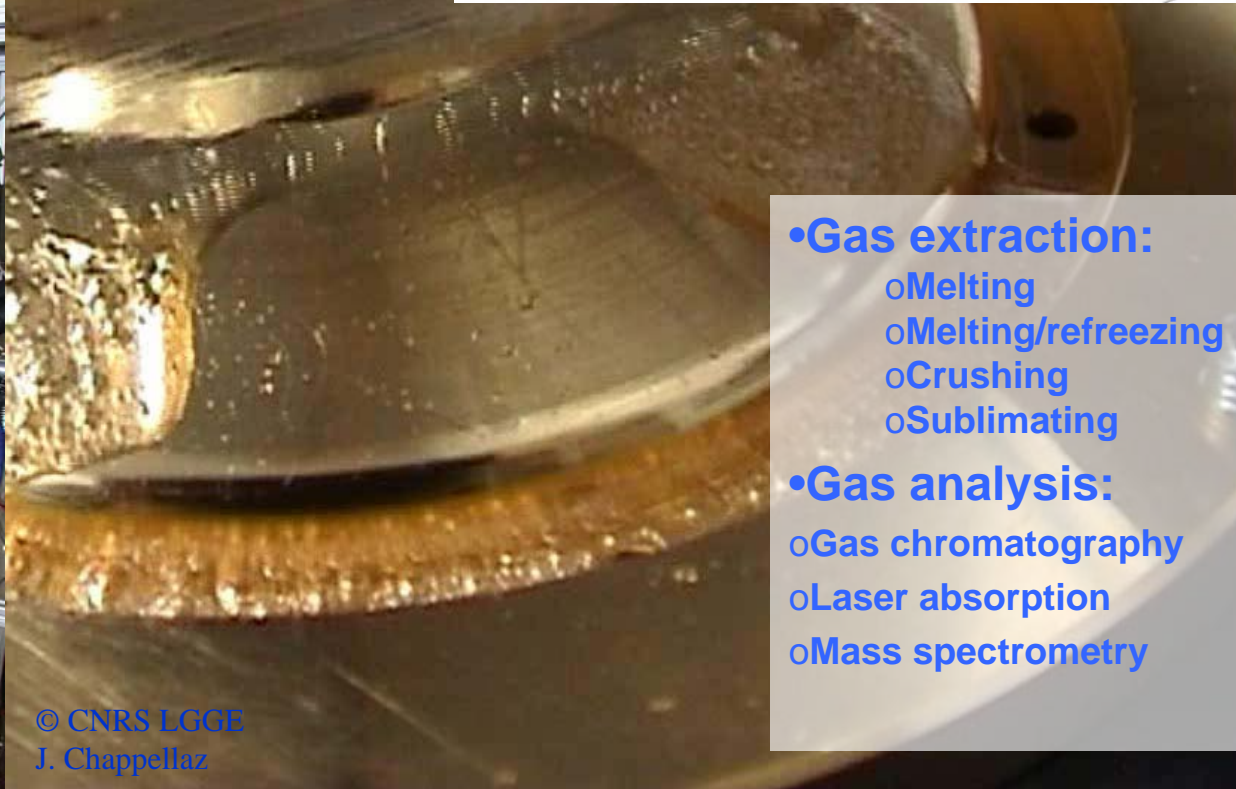
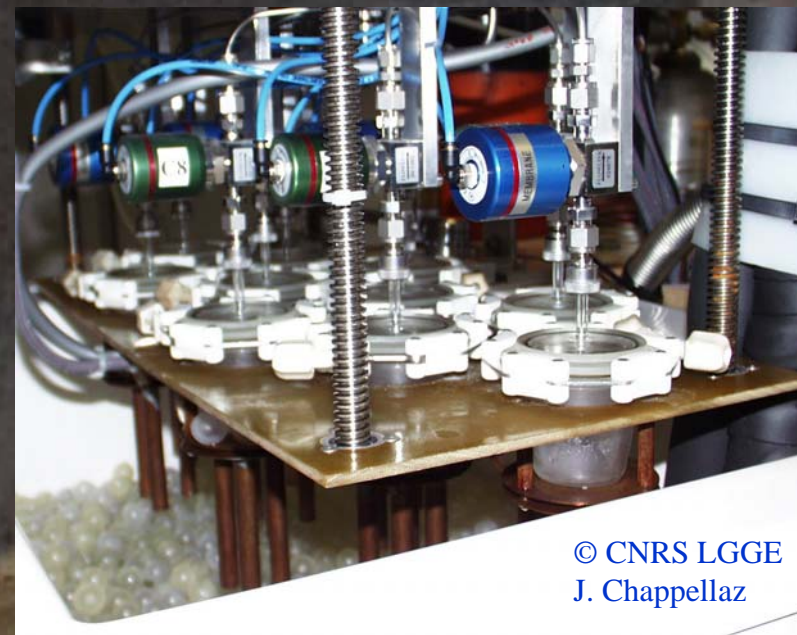
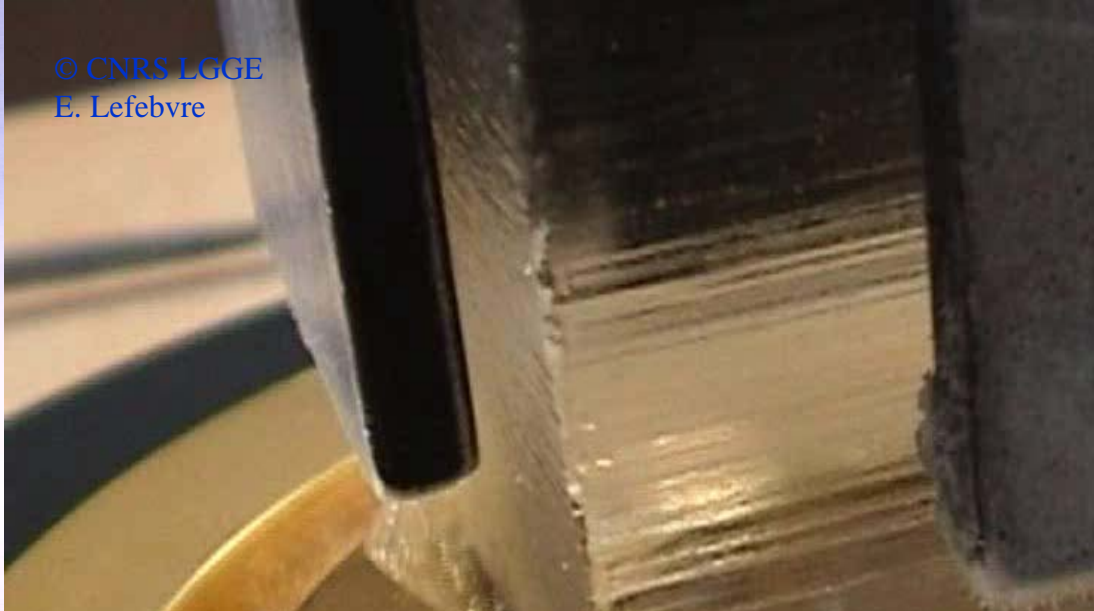


*Willi Dansgaard*



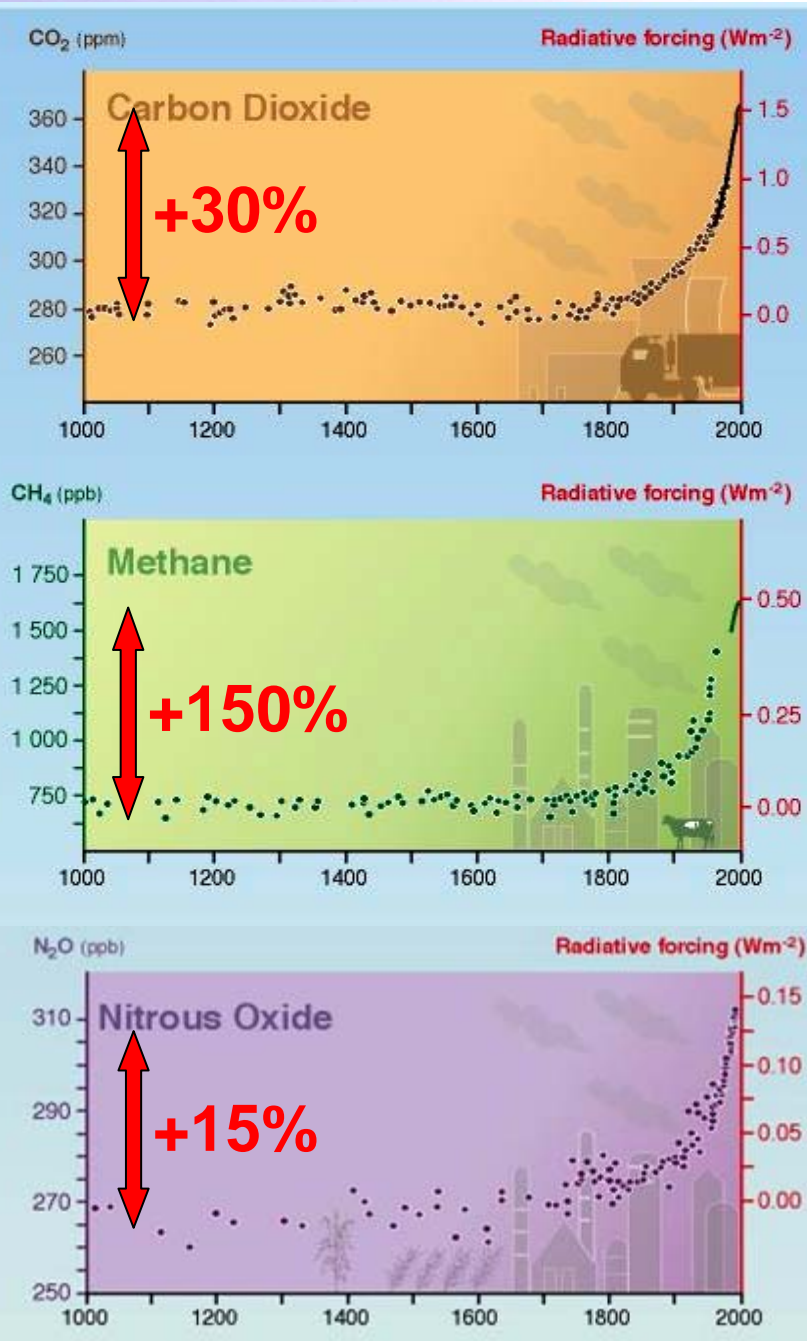
*Hans Oeschger*

**Claude Lorius**  
**Vostok station**  
**1983**



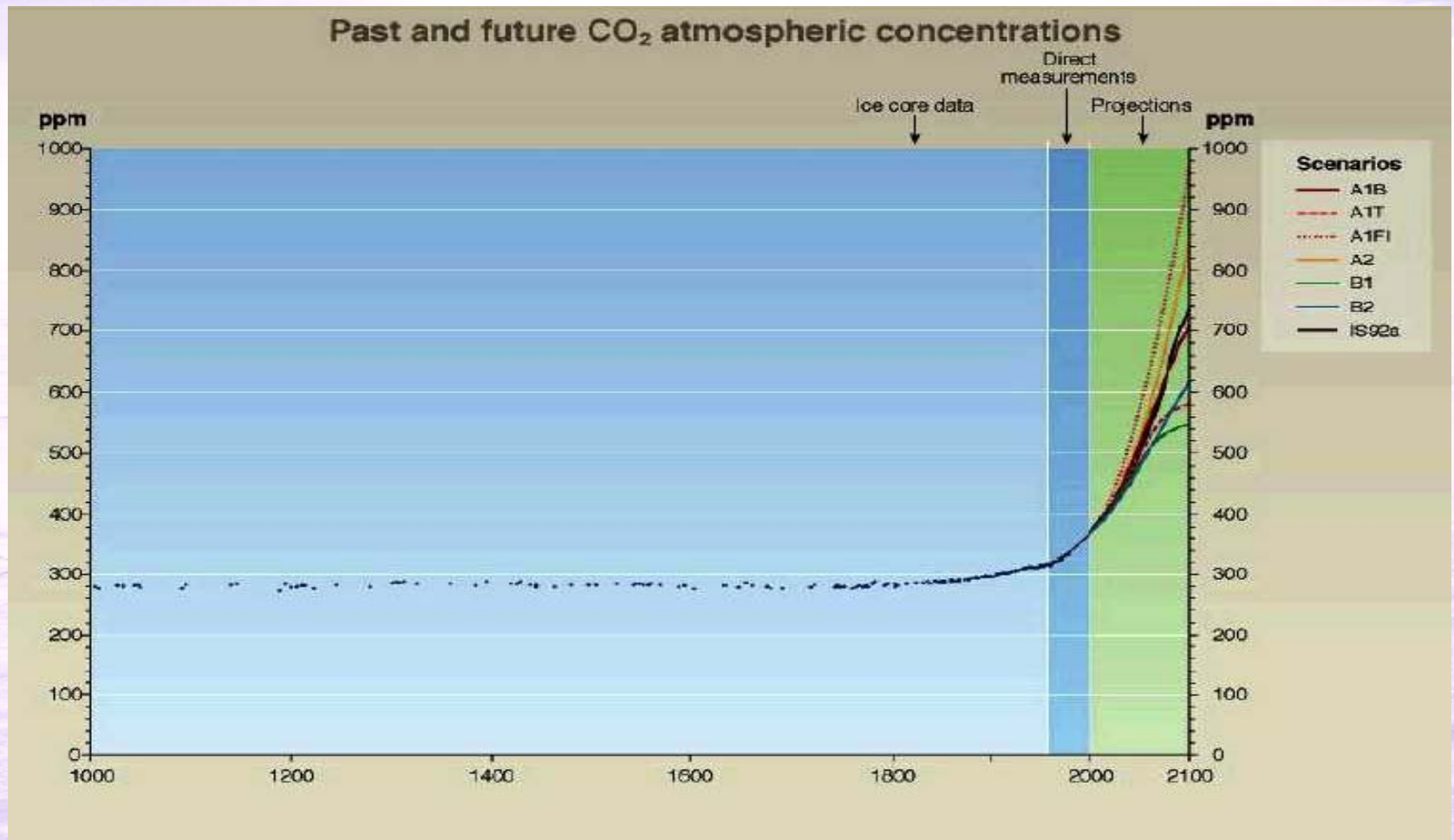
- **Gas extraction:**
  - o Melting
  - o Melting/refreezing
  - o Crushing
  - o Sublimating
- **Gas analysis:**
  - o Gas chromatography
  - o Laser absorption
  - o Mass spectrometry

# The ice-core warning: GHG anthropogenic increase



*2001 IPCC report*

# Ice-core lessons from the past: A key for the future

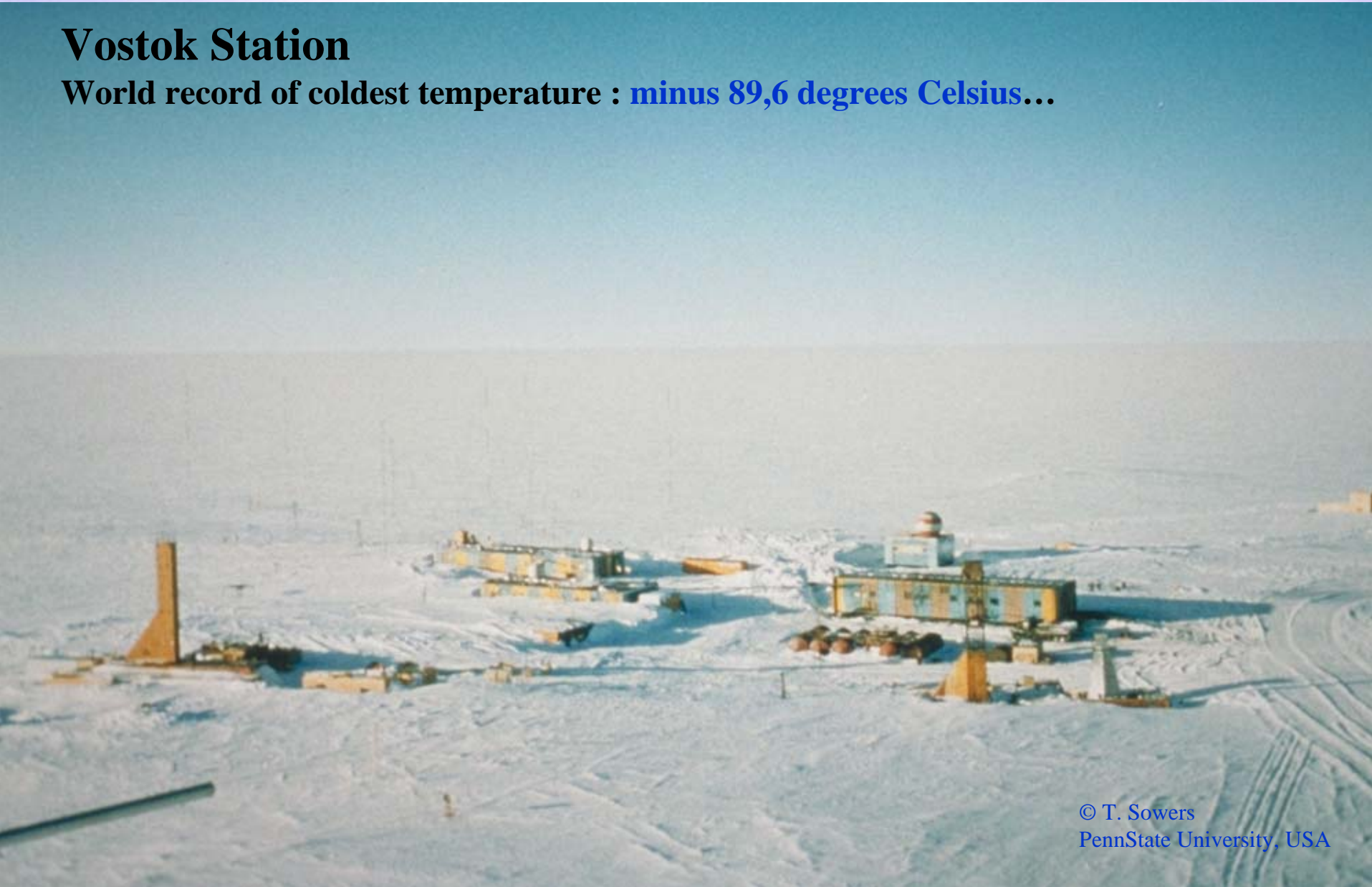


**Uncertainties on future GHG mixing ratios are linked to anthropogenic emission scenarios but also to natural feedbacks; looking into the past allows one to constraint the climate/GHG links**

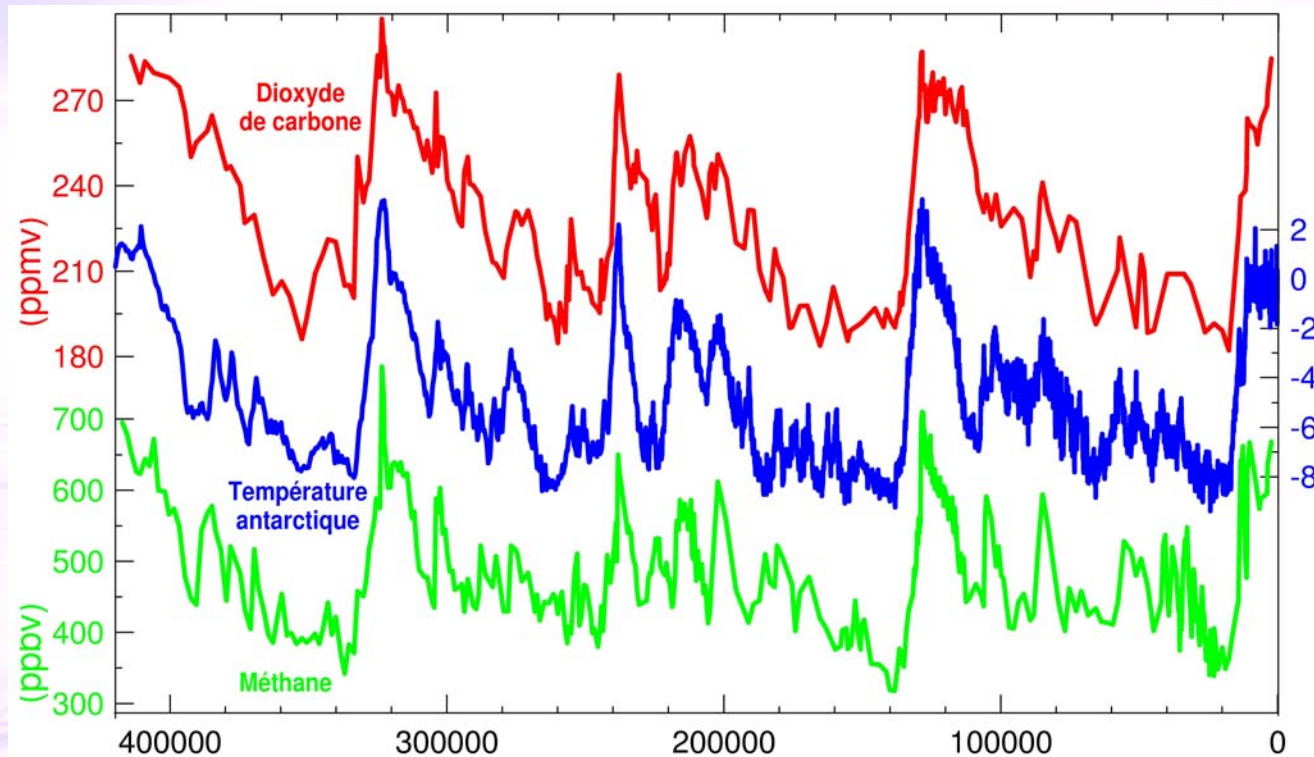
# Four Glacial-interglacial cycles: Vostok

## Vostok Station

World record of coldest temperature : **minus 89,6 degrees Celsius...**



# Four Glacial-interglacial cycles: Vostok



TIME →

*Petit et al., Nature 1999*

**High co-variance of temperature, carbon dioxide and methane**

**Maximum range of natural changes :**

**CO<sub>2</sub> : 185-300 ppmv (~20 ppmv / °C)**

**CH<sub>4</sub> : 350-800 ppbv (~75 ppbv / °C)**

# Eight Glacial-interglacial cycles: EPICA/Dome C



DC2  
1999/2000 : casing 112m →

DC1  
← 1996/1997 : casing 108m

← 1997/1998 : 364m

← 1998/1999 : 781m

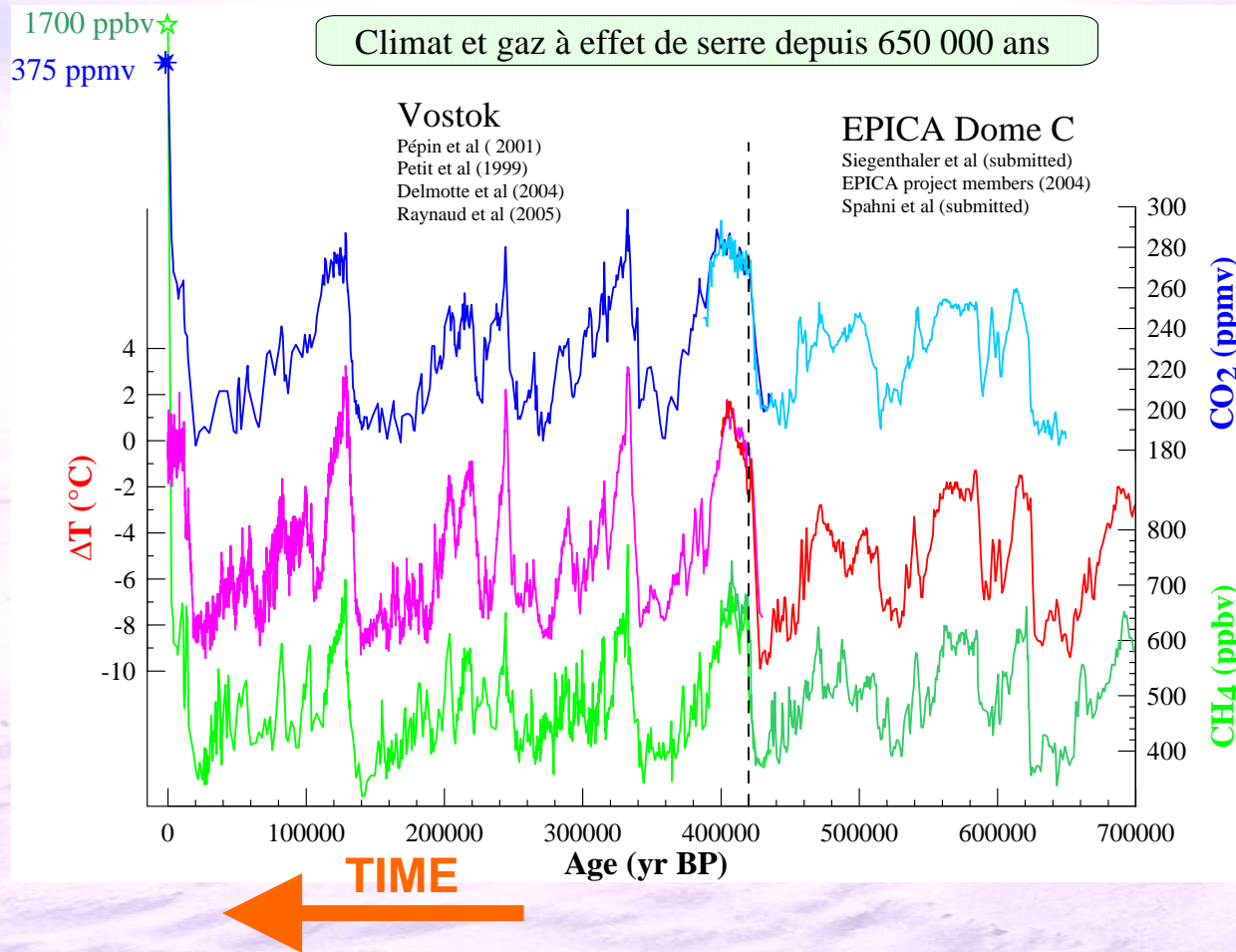
2000/2001 : 1459m →

2001/2002 : 2864m →

2002/2003 : 3200m →

2004/2005 : 3270m →

# Six Glacial-interglacial cycles: EPICA/Dome C

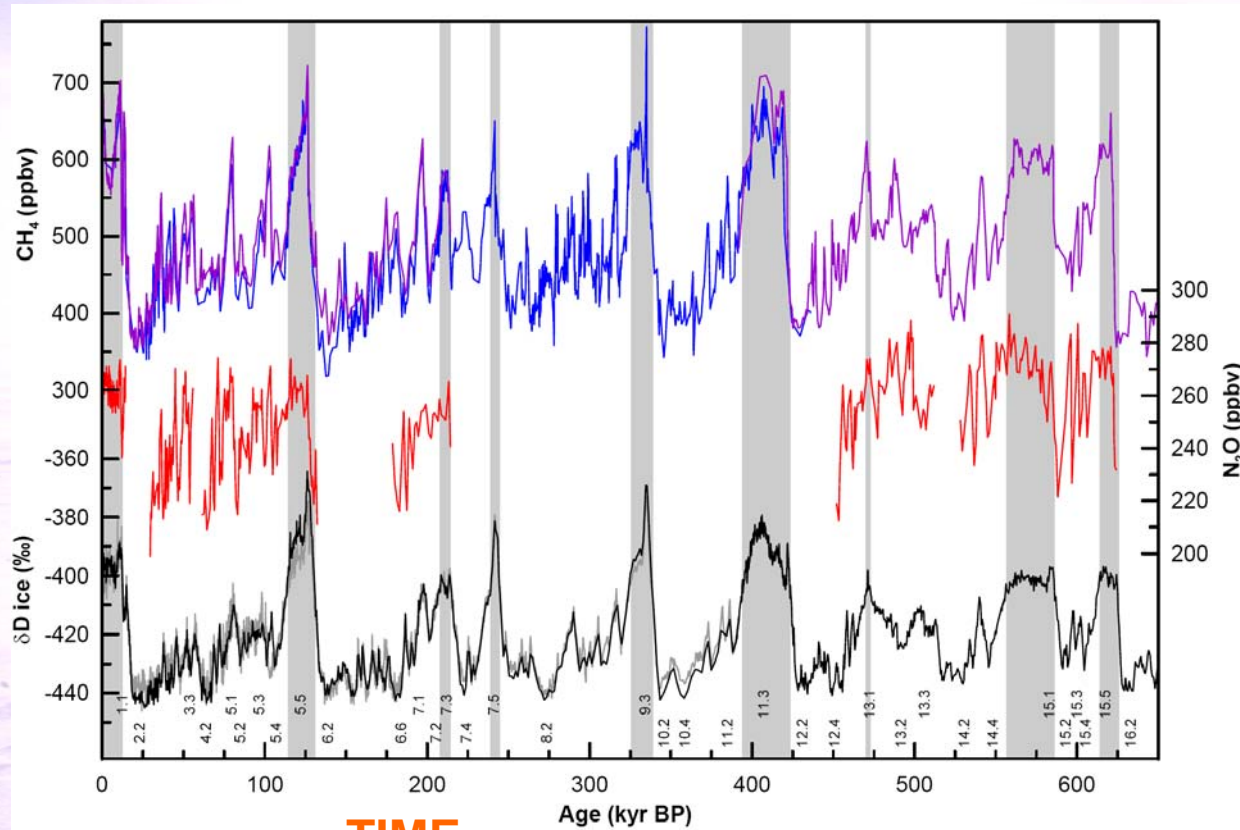


Maximum range of natural changes :

$\text{CO}_2$  : 185-300 ppmv ( $\sim 20$  ppmv /  $^{\circ}\text{C}$ )

$\text{CH}_4$  : 350-800 ppbv ( $\sim 75$  ppbv /  $^{\circ}\text{C}$ )

# Six Glacial-interglacial cycles: N<sub>2</sub>O



Spahni et al. Science, 2005

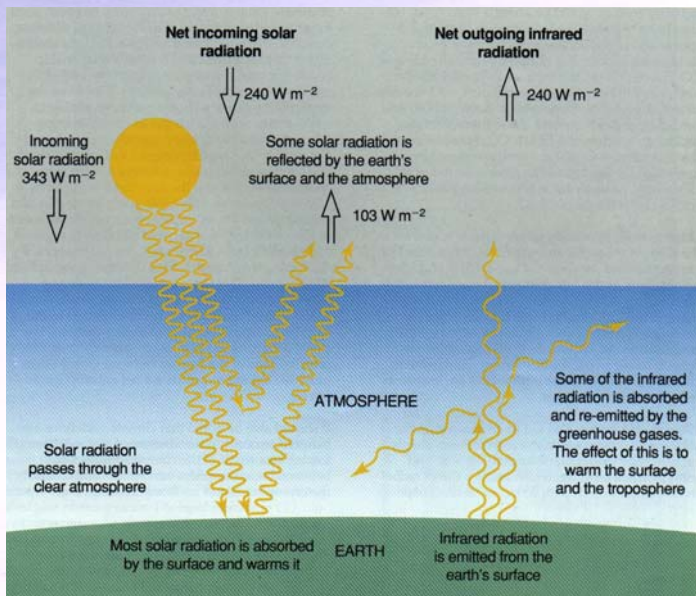
Maximum range of natural changes :

CO<sub>2</sub> : 185-300 ppmv (~20 ppmv / °C)

CH<sub>4</sub> : 350-800 ppbv (~75 ppbv / °C)

N<sub>2</sub>O : 200-275 ppbv (~15 ppbv / °C)

# Learning from the past: GHG as climatic feedbacks

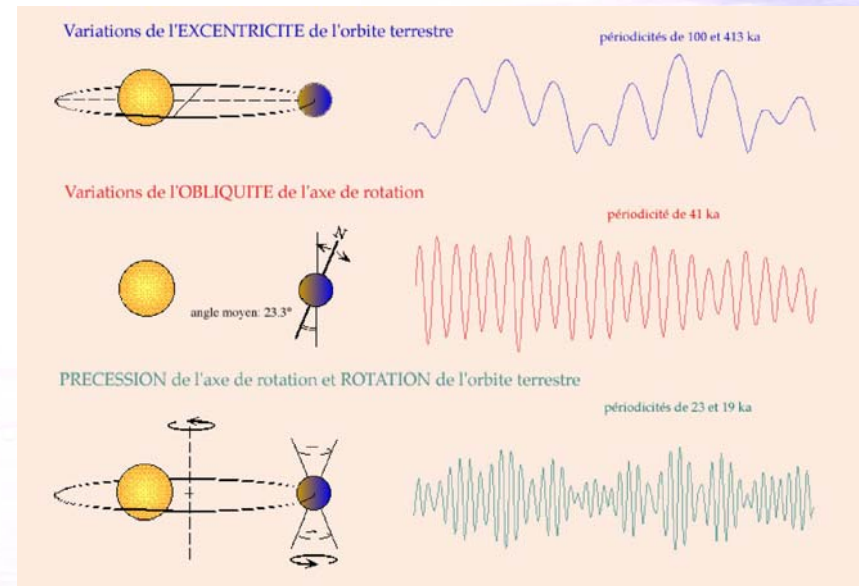


## Energy balance from glacial to interglacial conditions :

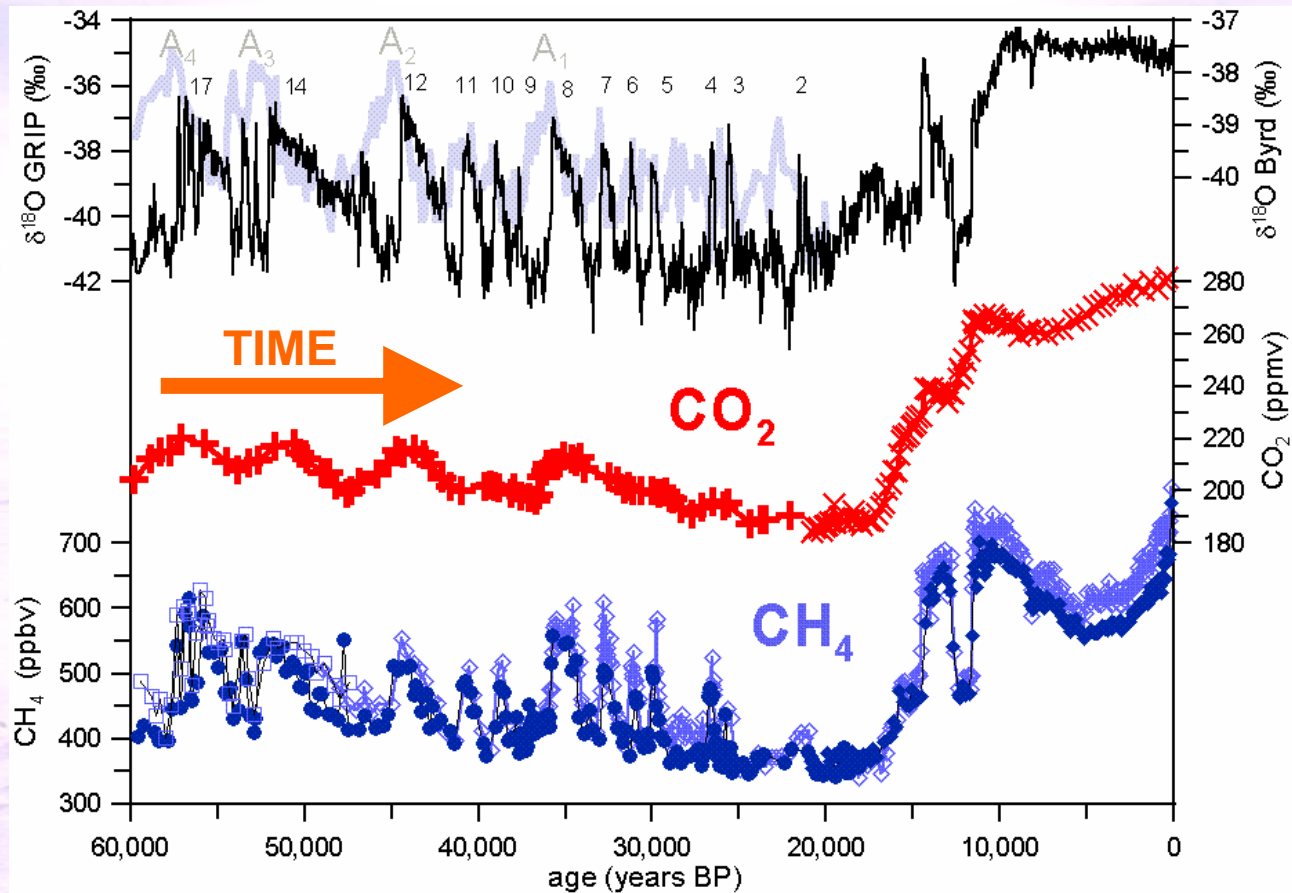
- + $2.6 \pm 0.5 \text{ W/m}^2$  from the combined effect of  $\text{CO}_2$ ,  $\text{CH}_4$  and  $\text{N}_2\text{O}$
- + $3.5 \pm 1 \text{ W/m}^2$  from the albedo effect (snow, ice and vegetation)
- + $0.5 \pm 1 \text{ W/m}^2$  from dust and aerosols

## Sequence of events from ice core data suggests :

- early warming in the South, forced by insolation
- GHG amplification
- albedo, dust and aerosol amplification



# Learning from the past: the timing of events



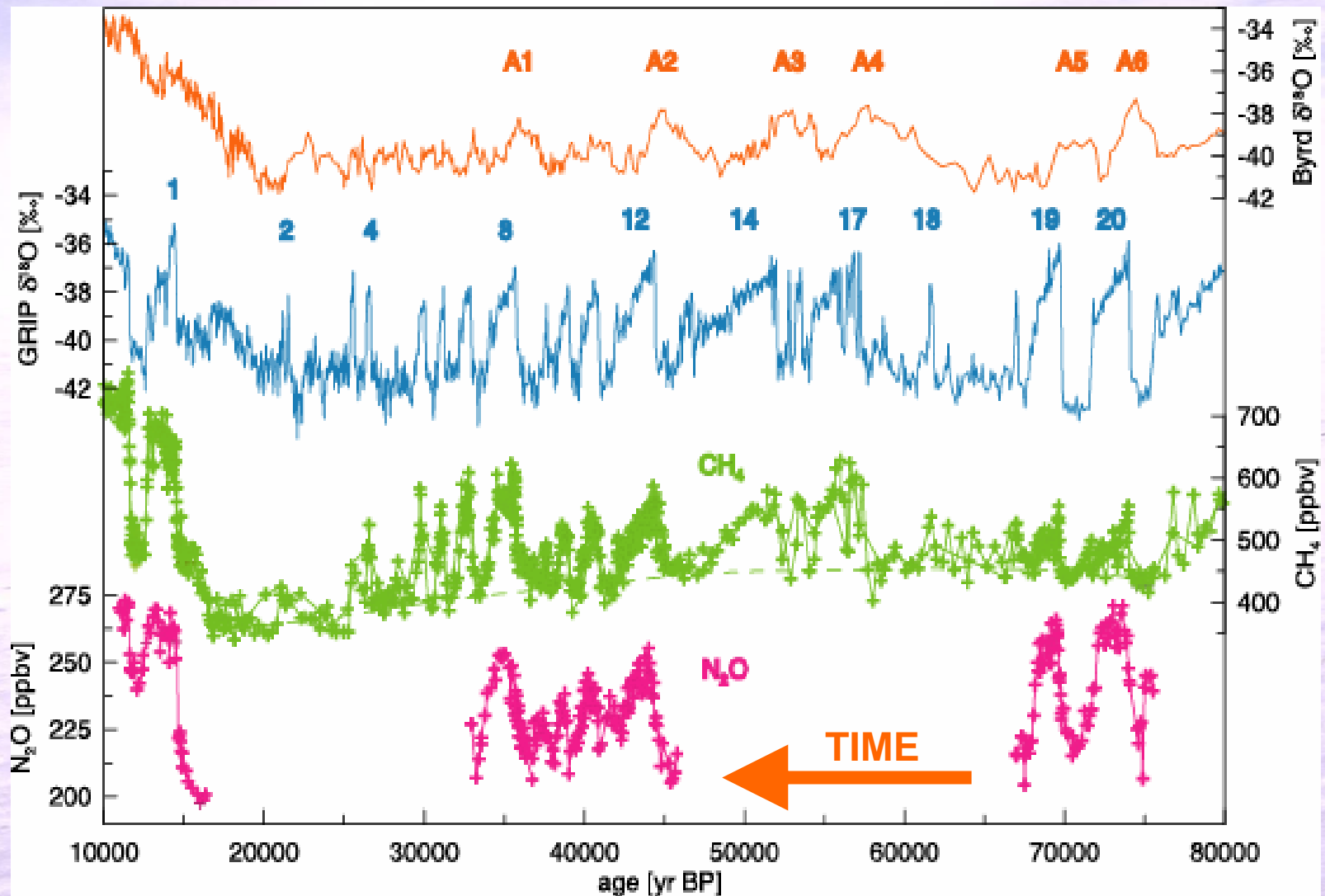
*Stauffer et al., 2002 and ref. therein*

**CO<sub>2</sub> :** 20 ppmv variability correlated with Antarctic temperature

**CH<sub>4</sub> :**

- 100-200 ppbv variability associated with changes in North Atlantic and Greenland climate
- synchronous with  $t^\circ$  or lags by a few decades
- increases over 50 to 150 yr

# Learning from the past: the timing of events



*Flückiger et al., GBC 2004*

**$\text{N}_2\text{O}$  : relatively large (40-50 ppbv) variability associated with D-O events, but for long-lasting events  $\text{N}_2\text{O}$  increases earlier than  $\text{CH}_4$  and Greenland T**

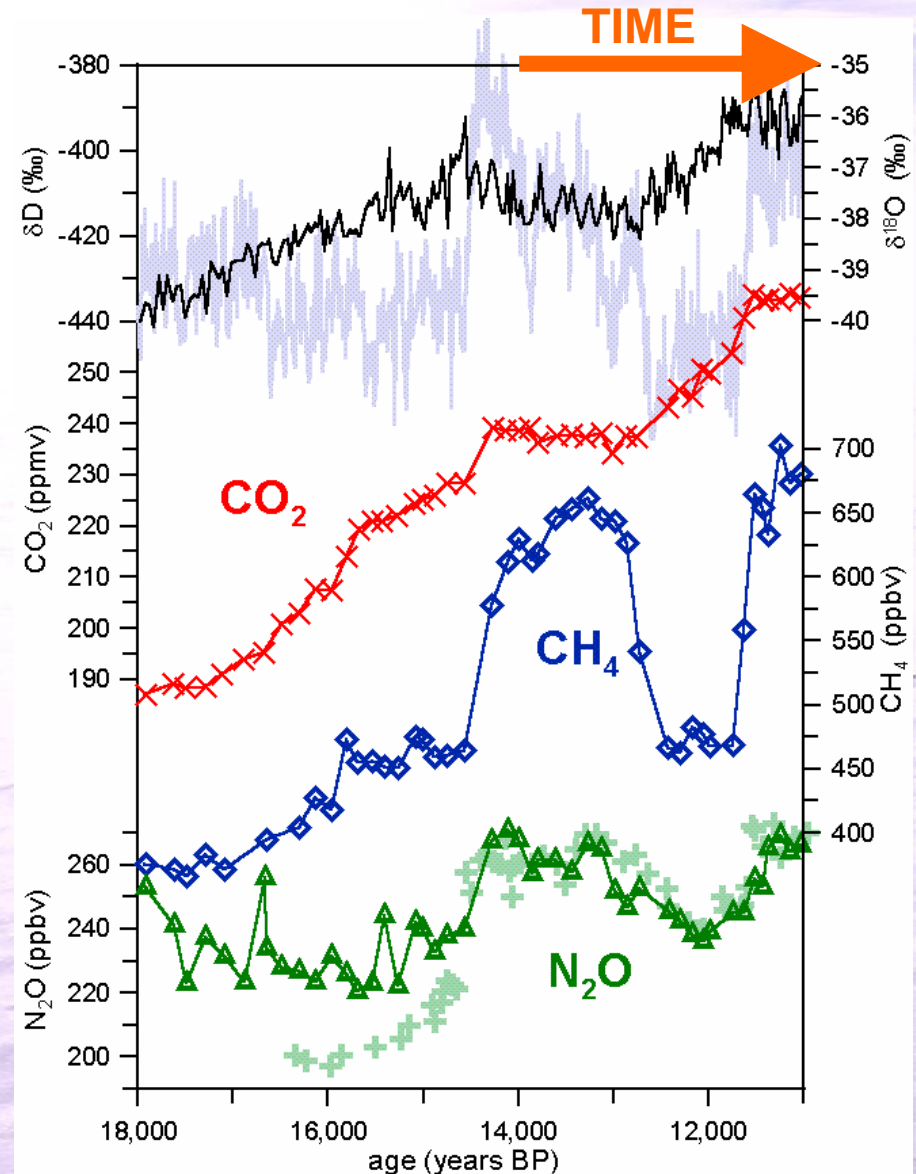
# Learning from the past: the timing of events

## The last Glacial-interglacial transition

$\text{CO}_2$  : parallels Antarctic warming

$\text{CH}_4$  : parallels N. Atlantic warming

$\text{N}_2\text{O}$  : parallels N. Atlantic warming  
but with slower response than  $\text{CH}_4$



*Stauffer et al., 2002 and ref. therein*

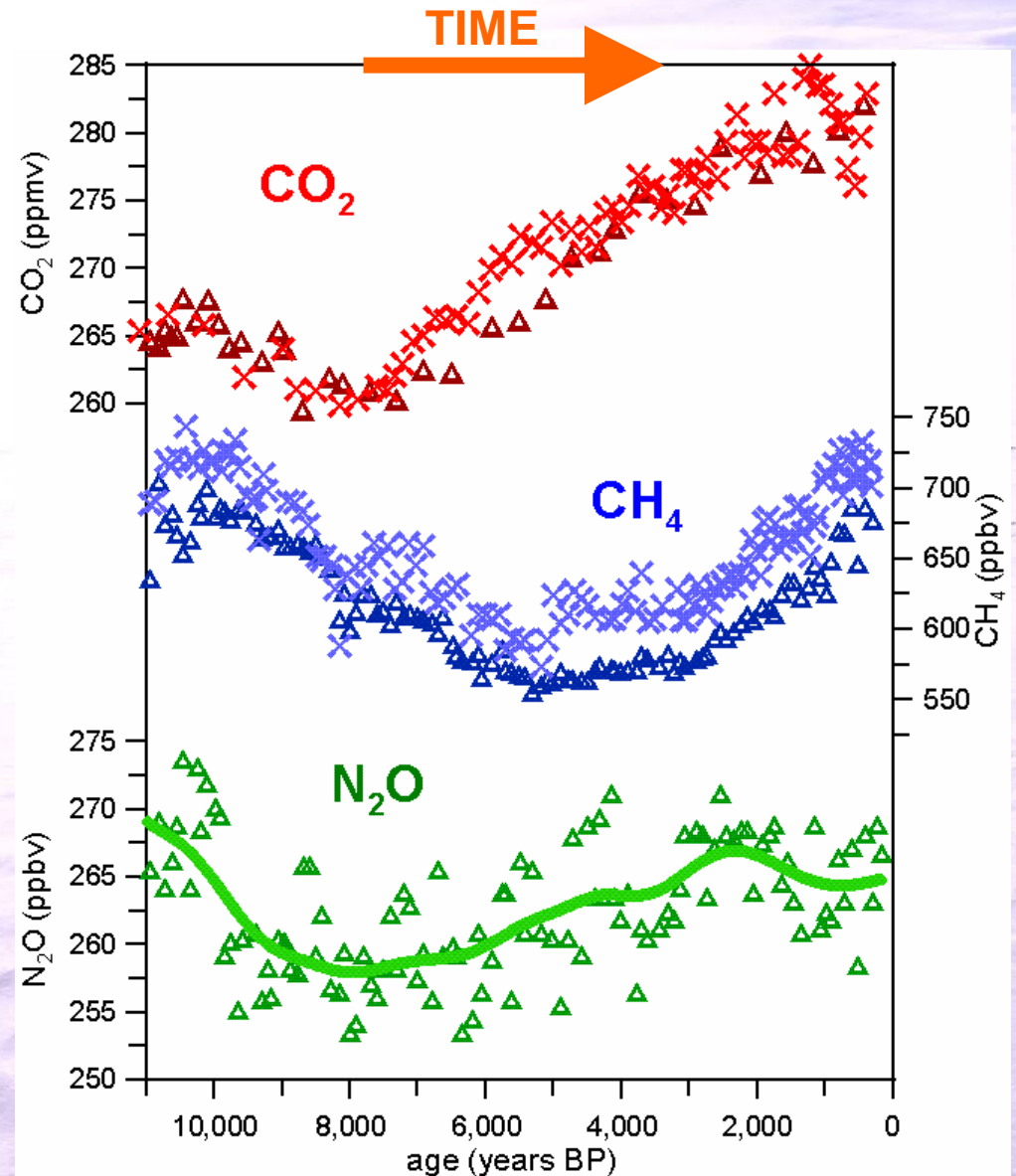
# Learning from the past: the timing of events

## The Holocene

$\text{CO}_2$  : 20 ppmv change with minimum around 8000 yr BP

$\text{CH}_4$  : 150 ppbv change with minimum around 5000 yr BP

$\text{N}_2\text{O}$  : 15 ppbv change with minimum around 8000 yr BP



*Stauffer et al., 2002 and ref. therein*

# Summary of ice-core observations relevant to the climate/GHG relationship

All GHG : range of natural variability appears remarkably narrowed and associated either with northern or southern climate records

CO<sub>2</sub> : slow (millennial) evolution mostly correlated with Southern latitude climate, except during the Holocene and the glacial inceptions

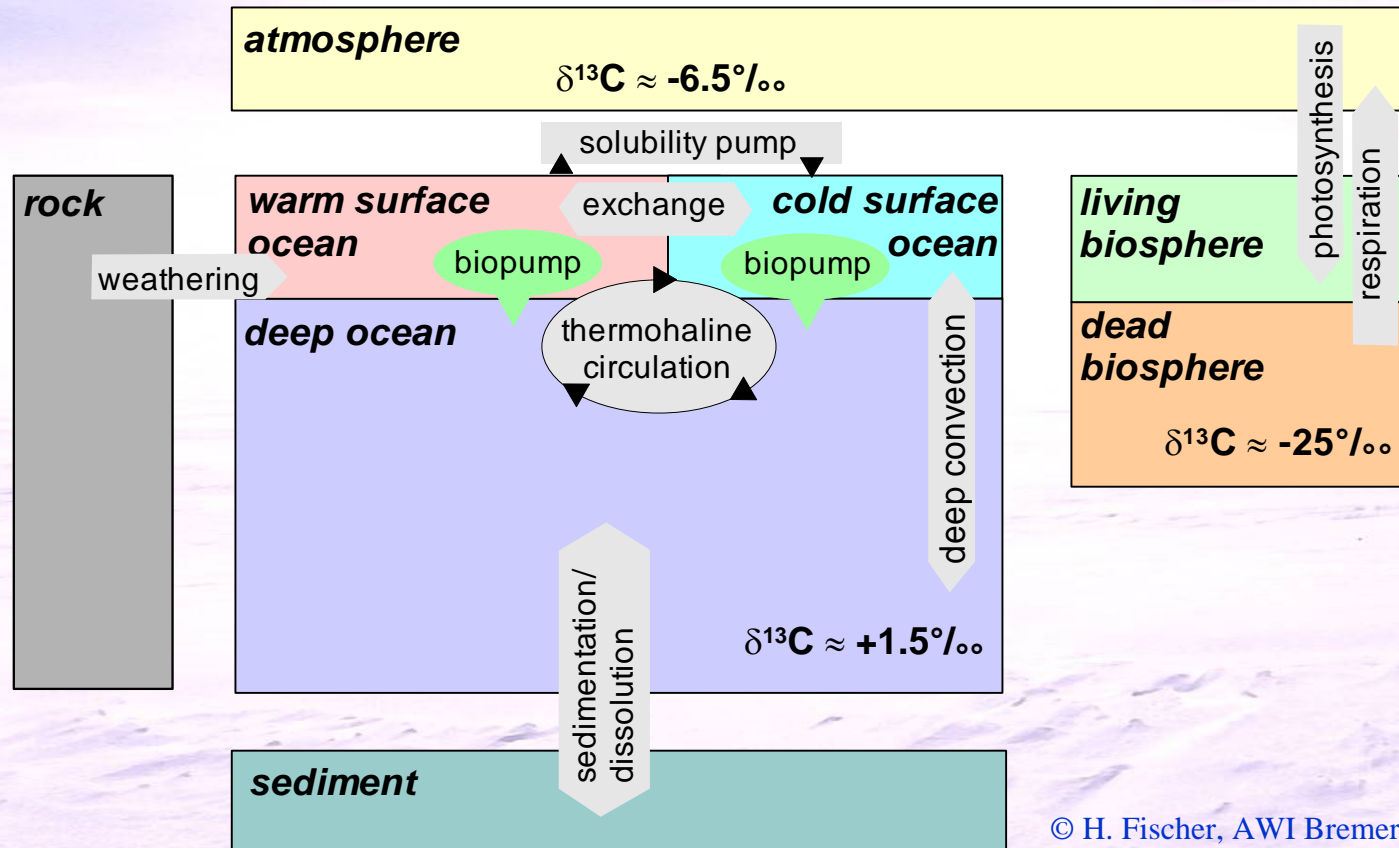
CH<sub>4</sub> : rapid (centennial) evolution mostly correlated with Northern Atlantic climate, except during the Holocene

N<sub>2</sub>O : depending on the time period, mimics CO<sub>2</sub> (Holocene) or CH<sub>4</sub> (D-O events, deglaciation)

# What explanations ?

- Ice core measurements cannot answer alone
- A combination of other observations (marine and continental realms) and modelling is required
- The search for the holy grail continues...
- But... Here is where we stand...

# Causes of CO<sub>2</sub> natural variability



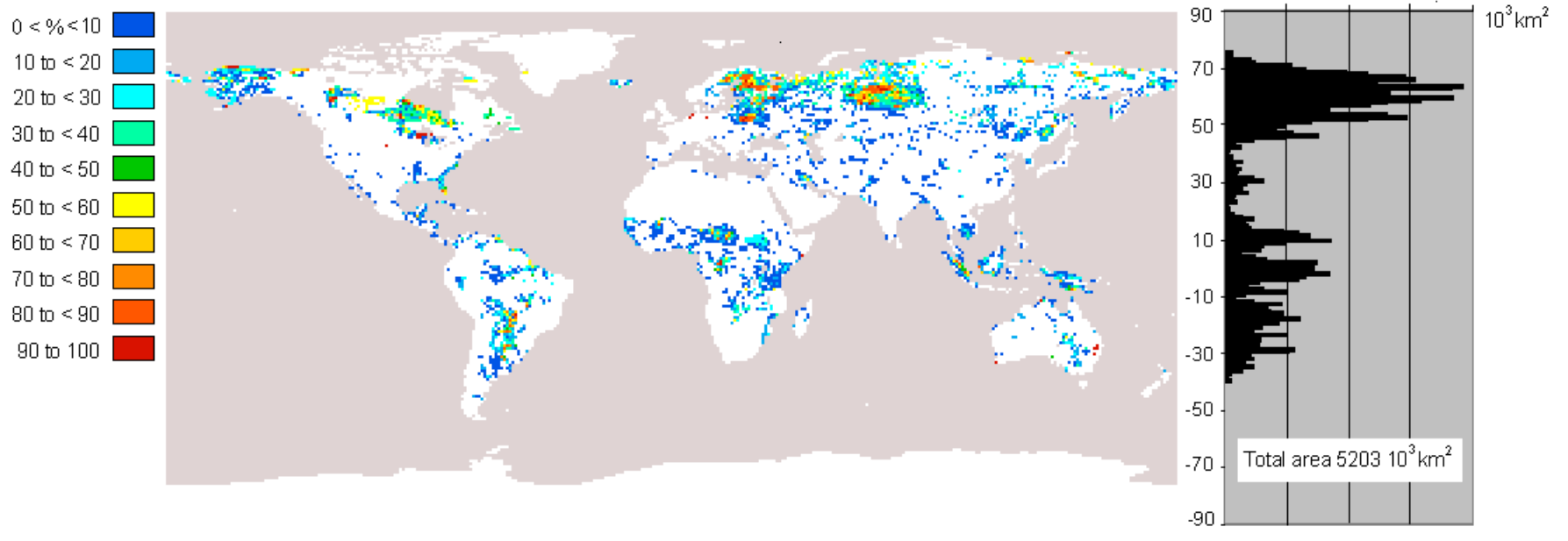
© H. Fischer, AWI Bremerhaven, Germany

**Probable dominant role of processes within the Southern ocean (surface temperature, sea-ice extent, iron fertilisation, surface/deep ocean exchange)**

**Increase since 8000 yr BP (and maybe « high » CO<sub>2</sub> during glacial inceptions) possibly related with continental biomass reduction**

# Causes of CH<sub>4</sub> natural variability

- Wetlands : 3/4 of natural emissions



*Matthews and Fung, GBC, 1987*

**Boreal versus Tropical ?**

# Causes of CH<sub>4</sub> natural variability

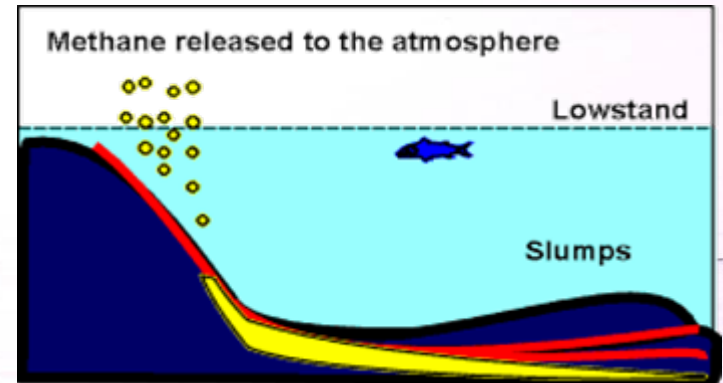
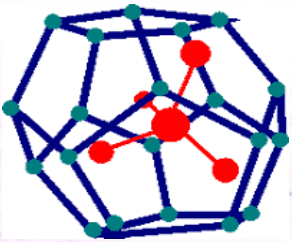
- Oxidative capacity of the atmosphere

Type of model (reference)	Difference between OH at Last Glacial Maximum and OH today (in %) <i>A positive number means more OH at the Last Glacial Maximum</i>
Bi-dimensional photochemical (Valentin, thesis U. Mainz, 1990)	+30-40%
Uni-dimensional photochemical (Pinto and Khalil, Tellus 1991)	+20%
Uni-dimensional photochemical (Lu and Khalil, Tellus 1992)	+40%
Uni-dimensional photochemical multi- box (Thompson et al., Tellus 1993)	+32%
Uni-dimensional photochemical convective with two boxes (Crutzen and Bruhl, Geophys. Res. Lett. 1993)	-5%
Bi-dimensional coupled climate- chemistry (Martinerie et al., J. Geophys. Res. 1995)	+13%
Uni-dimensional coupled climate- chemistry (Karol et al., J. Geophys. Res. 1995)	-63% à +5%
Tri-dimensional forced with climate simulations (Valdes et al., Geophys. Res. Lett. 2005)	+25%*
Tri-dimensional coupled climate- chemistry (Kaplan et al., Global Biogeochem. Cycles 2006)	+28%*

Amplifying role but requires a source forcing

# Causes of CH<sub>4</sub> natural variability

- Hydrates



Where decadal CH<sub>4</sub> changes have been measured in ice cores, they do not support a catastrophic hydrate degassing scenario

## Sources of Nitrous Oxides



CG Figure 39

## Causes of N<sub>2</sub>O natural variability

The natural N<sub>2</sub>O budget involves nitrification and denitrification in tropical soils (~2/3) and in the open & coastal oceans (~1/3)

N<sub>2</sub>O changes during the last glaciation are associated with observed enhanced primary production and denitrification in upwelling zones of the global ocean. Could explain part of the amplitude and timing of N<sub>2</sub>O

But preliminary isotopic measurements on the N<sub>2</sub>O molecule in ice cores suggest a ~constant continent/ocean N<sub>2</sub>O source ratio...

Role of troposphere/stratosphere exchange on N<sub>2</sub>O residence time ?  
Role of solar modulation on N<sub>2</sub>O photodissociation (sink) ?

# Conclusion: what do we learn from GHG measurements in ice cores ?

- The range of GHG natural changes is remarkably narrow in the course of the last six glacial-interglacial cycles ; GHG played a major role in amplifying insolation changes driving the natural climate changes
- The strong coupling between the natural evolution of CO<sub>2</sub>, CH<sub>4</sub> and N<sub>2</sub>O mixing ratios and climate implies that future natural feedbacks with climate warming must be expected
- Although the potential mechanisms are known, the exact nature of this coupling still remains much unclear and therefore the causes and amplitude of future feedbacks. At least for carbon dioxide, the southern ocean should be scrutinized.
- Ice cores have not yet given their last word...

# Perspectives on ice-core measurements of greenhouse gases and other related trace gases

Development of the isotopic fingerprint

But small sample size and high accuracy are challenging...

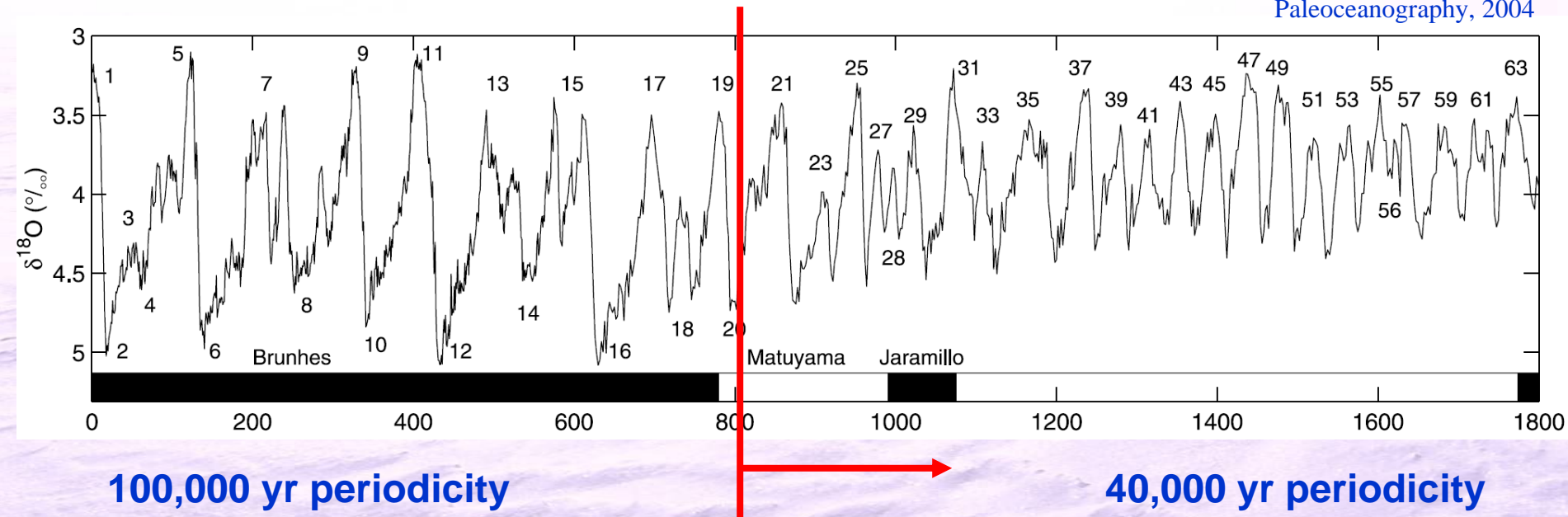
Isotopologue	Abundance ( $10^{-15}$ g per g of ice)	Constraint
$^{13}\text{CH}_4$	15 - 30	Biomass burning
$\text{CH}_3\text{D}$	0,2 - 0,5	Hydrates + OH at transient
$^{13}\text{CO}$	2 - 4	$\text{CH}_4$ oxidation
$\text{C}^{18}\text{O}$	0,4 - 0,9	Combustion sources
$\text{C}^{17}\text{O}$	0,1 - 0,2	OH change
$^{13}\text{CO}_2$	6000 - 12000	Continent/Ocean Biosphere
$^{15}\text{N}_2\text{O}$	3 - 5	N Cycle soil/ocean microbiology in ice
$\text{N}_2^{18}\text{O}$	1,5 - 3	
$\text{N}^{15}\text{NO}$	0,7 - 1,5	Tropo/strato exchange + solar cycle

# Perspectives on ice-core measurements of greenhouse gases and other related trace gases

International Partnership in Ice Core Sciences (IPICS)

International Polar Year : in search for the oldest ice in Antarctica

Lisiecki and Raymo,  
Paleoceanography, 2004



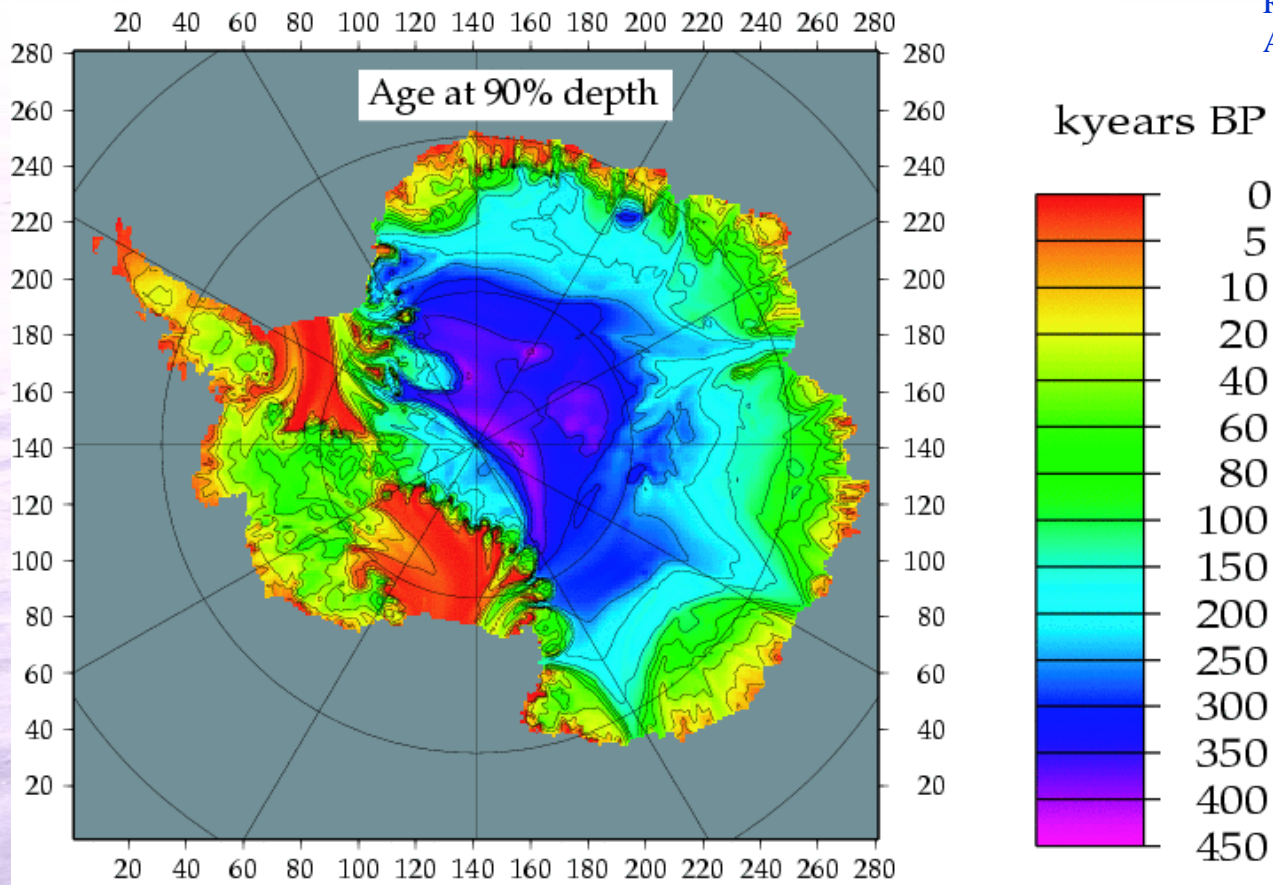
Higher  $\text{CO}_2$  atmospheric levels before 1 M yr ago ?

# Perspectives on ice-core measurements of greenhouse gases and other related trace gases

International Partnership in Ice Core Sciences (IPICS)

International Polar Year : in search for the oldest ice in Antarctica

Rybak and Huybrechts,  
Ann. Glaciology, 2003



*Thank you for your attention...*



*The EPICA / Dome C European drilling and science team*