

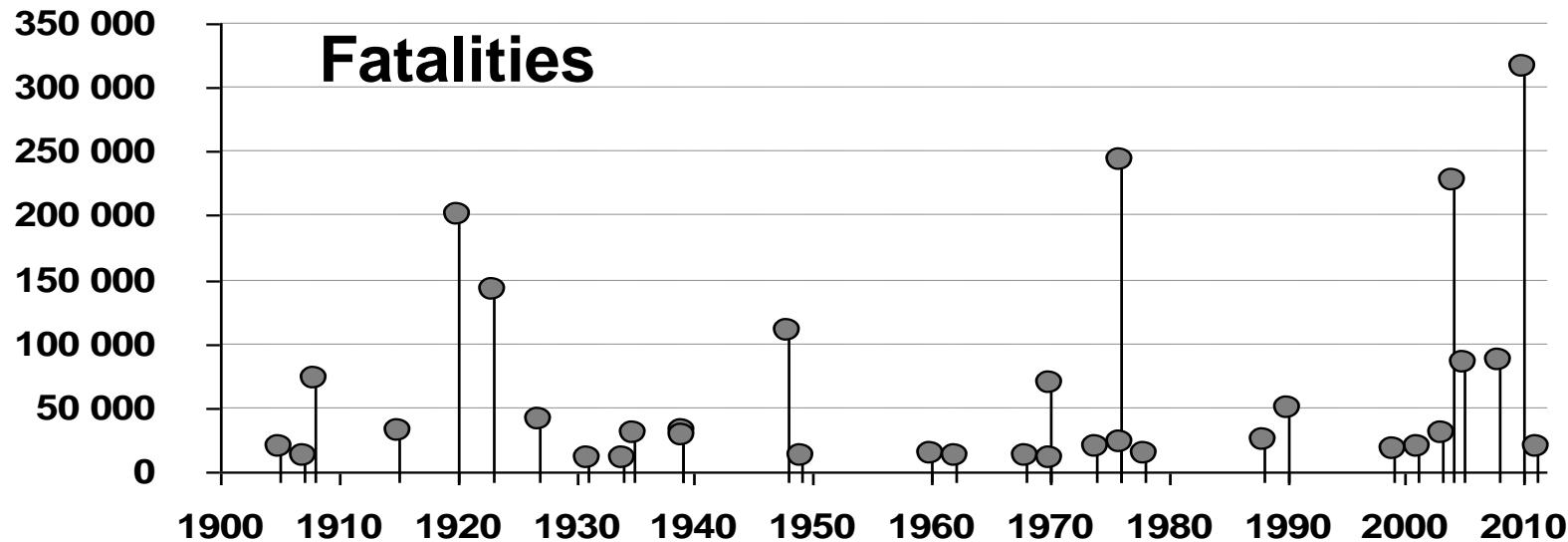
ORIGIN OF EARTHQUAKES IN AND AROUND STABLE PLATES

Christophe Vigny*



*département des Geosciences de l'ENS / UMR 8538 du CNRS
<http://www.geologie.ens.fr/~vigny>

IMPACT OF EARTHQUAKES



1. Haiti (2010): 316 000 dead Mw 7.0
2. China (1976): 243 000 dead Mw 7.5
3. Sumatra (2004) :228 000 dead Mw 9.2
4. China (1920): 200 000 dead Mw 7.8
5. Japan (1923): 143 000 dead Mw 7.9
6. Turkey (1948): 110 000 dead Mw 7.3
7. Sichuan (2008): 88 000 dead Mw 7.9
8. Pakistan (2005): 85 000 dead Mw 7.6
9. Messina (1908): 70 000 dead Mw 7.2
-
21. Japan (2011): 20 252 dead Mw 9.0

ECONOMIC IMPACT : GROWING !

Eq	Mw	Dead	Damage (billions USD)
Valdivia, Chile, 1960	9.5	5 000	~0.5
Sumatra 2004	9.2	300 000	~10
Maule, Chile 2010	8.8	800	15 – 30
Tohoku, Japan 2011	9.0	20 000	> 200



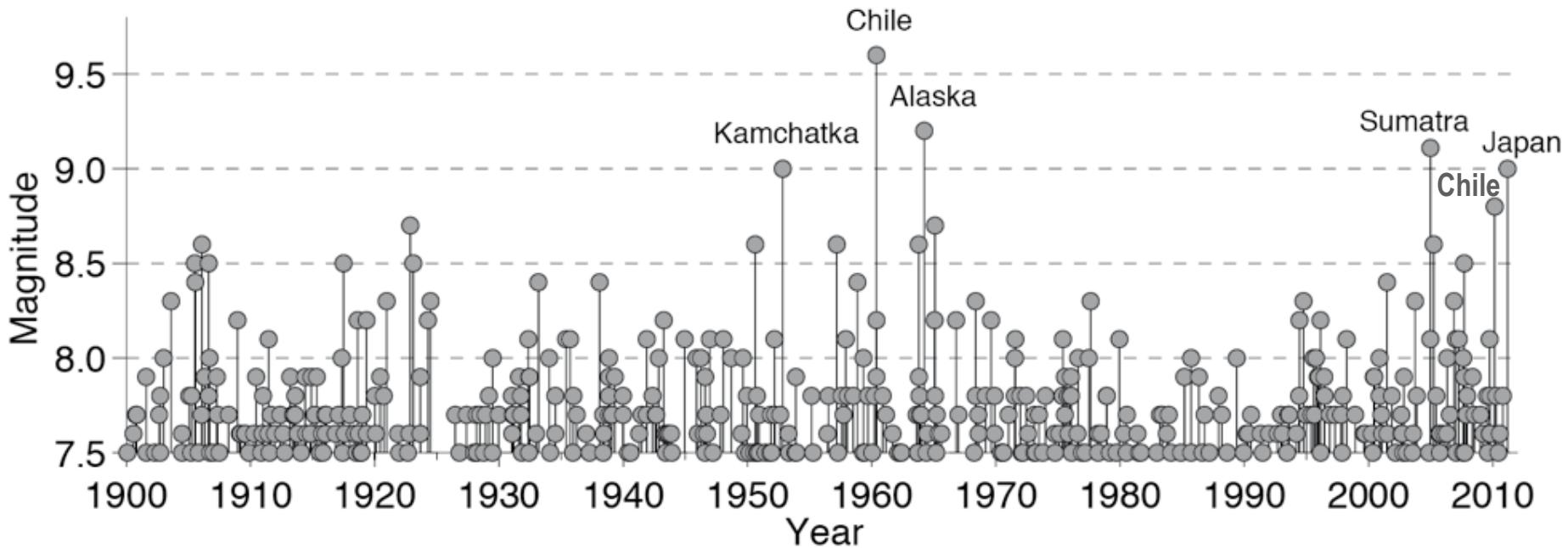
Valdivia 1960



Tohoku 2011

Q1 : HOW MANY EARTHQUAKES ?

Large Earthquakes (> Mw 7.5)



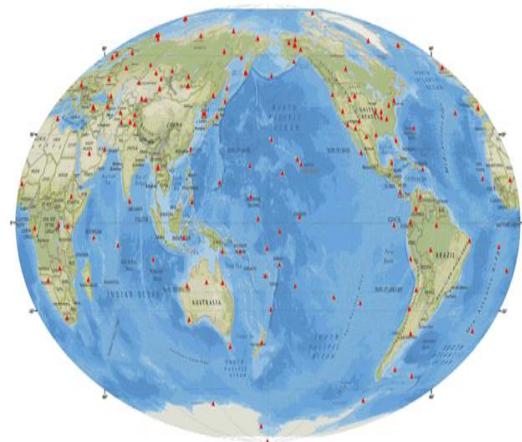
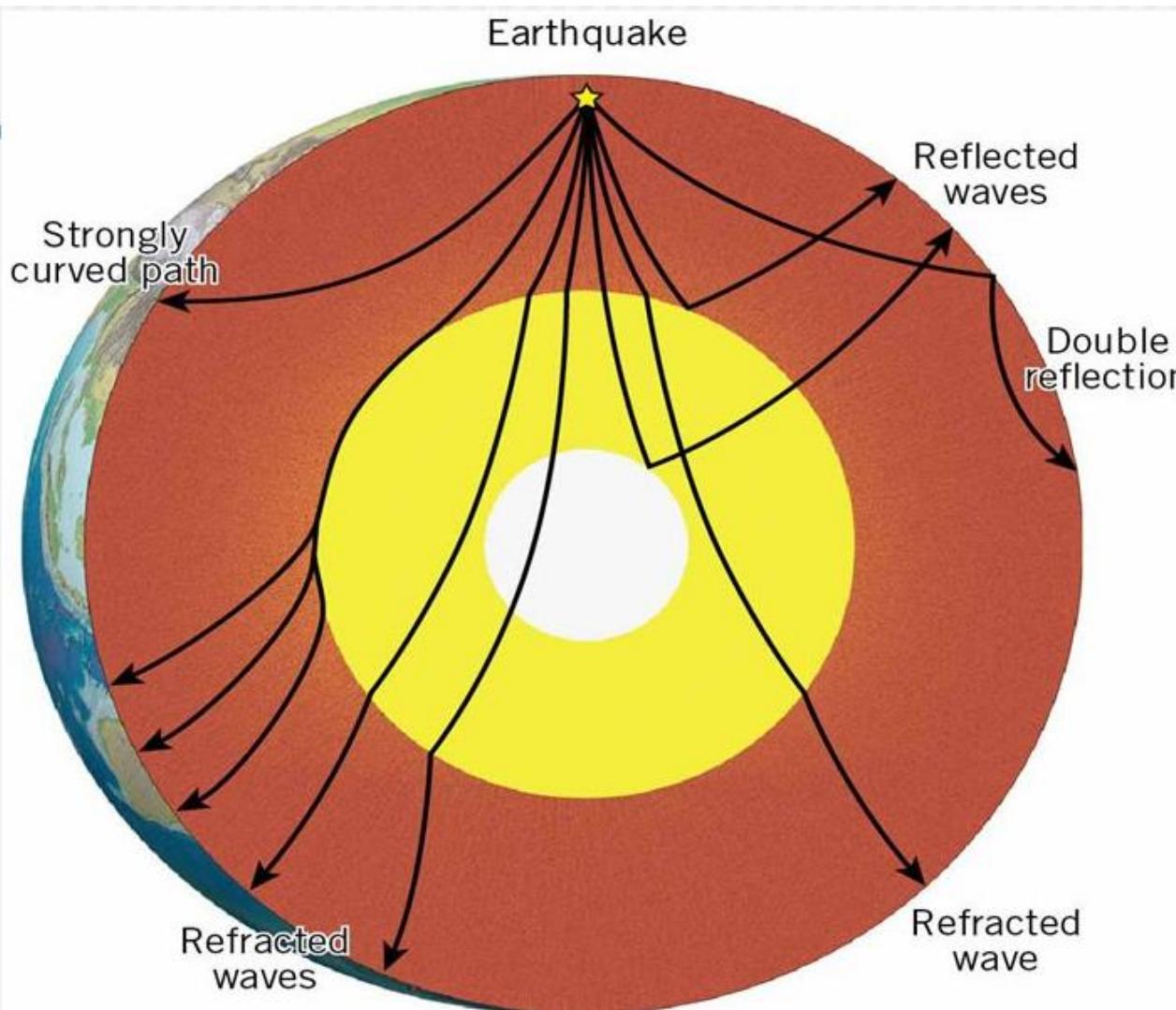
Large ones : every year !

Giant ones : 8-10 within 100 years

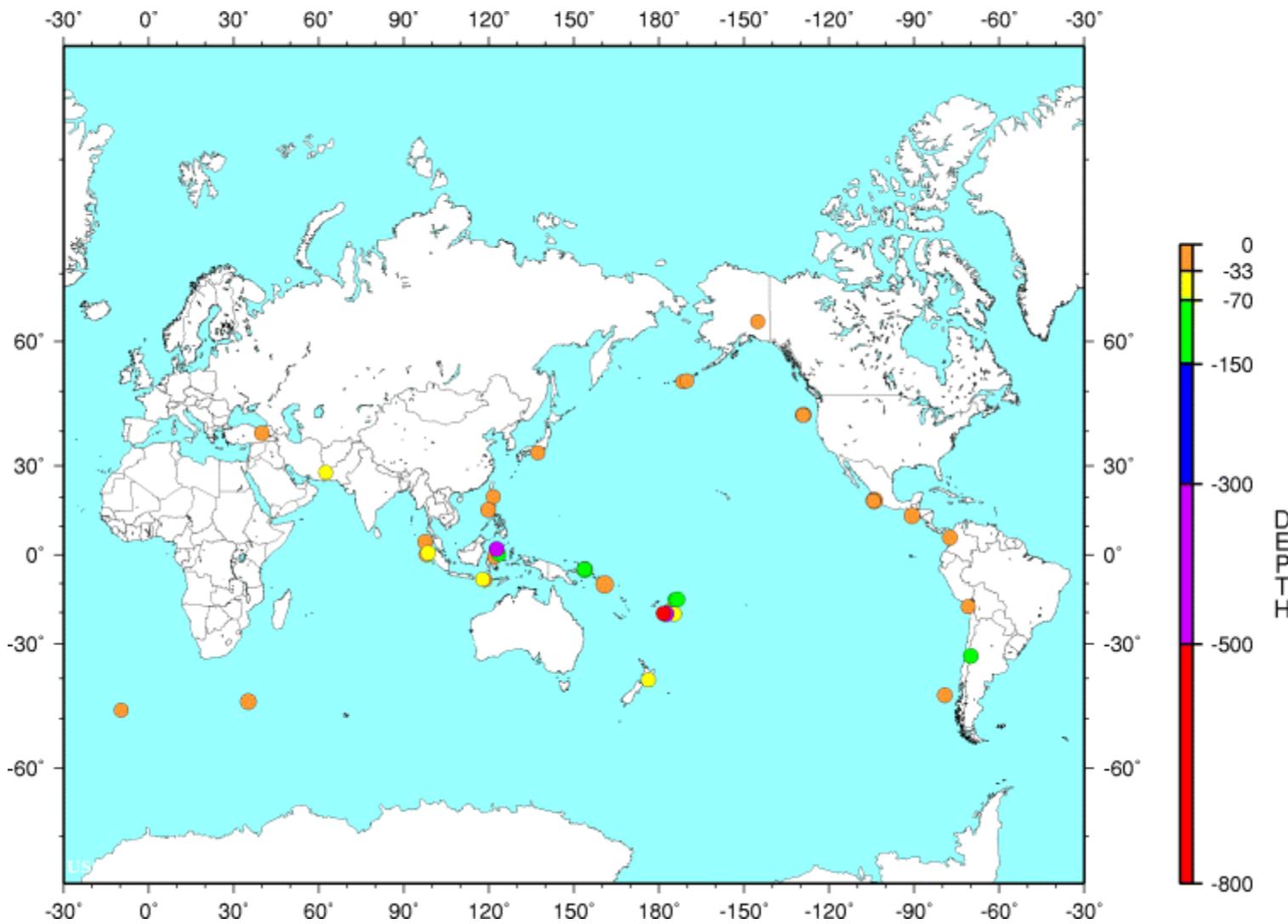
=> 1 per decade, but irregularly

40 years
quiand

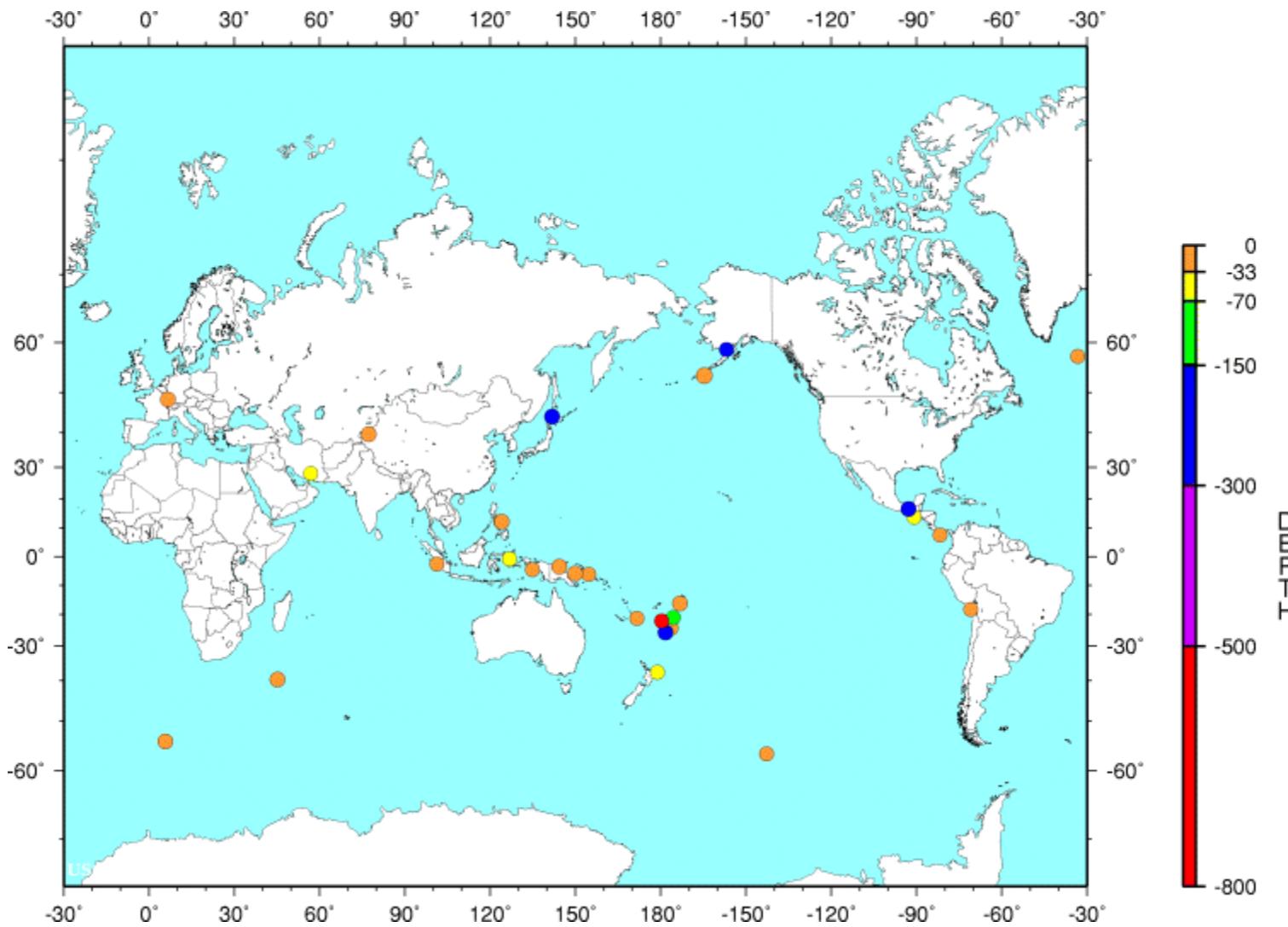
Earthquakes recorded by the global network of seismographs (> Mw 4.5)



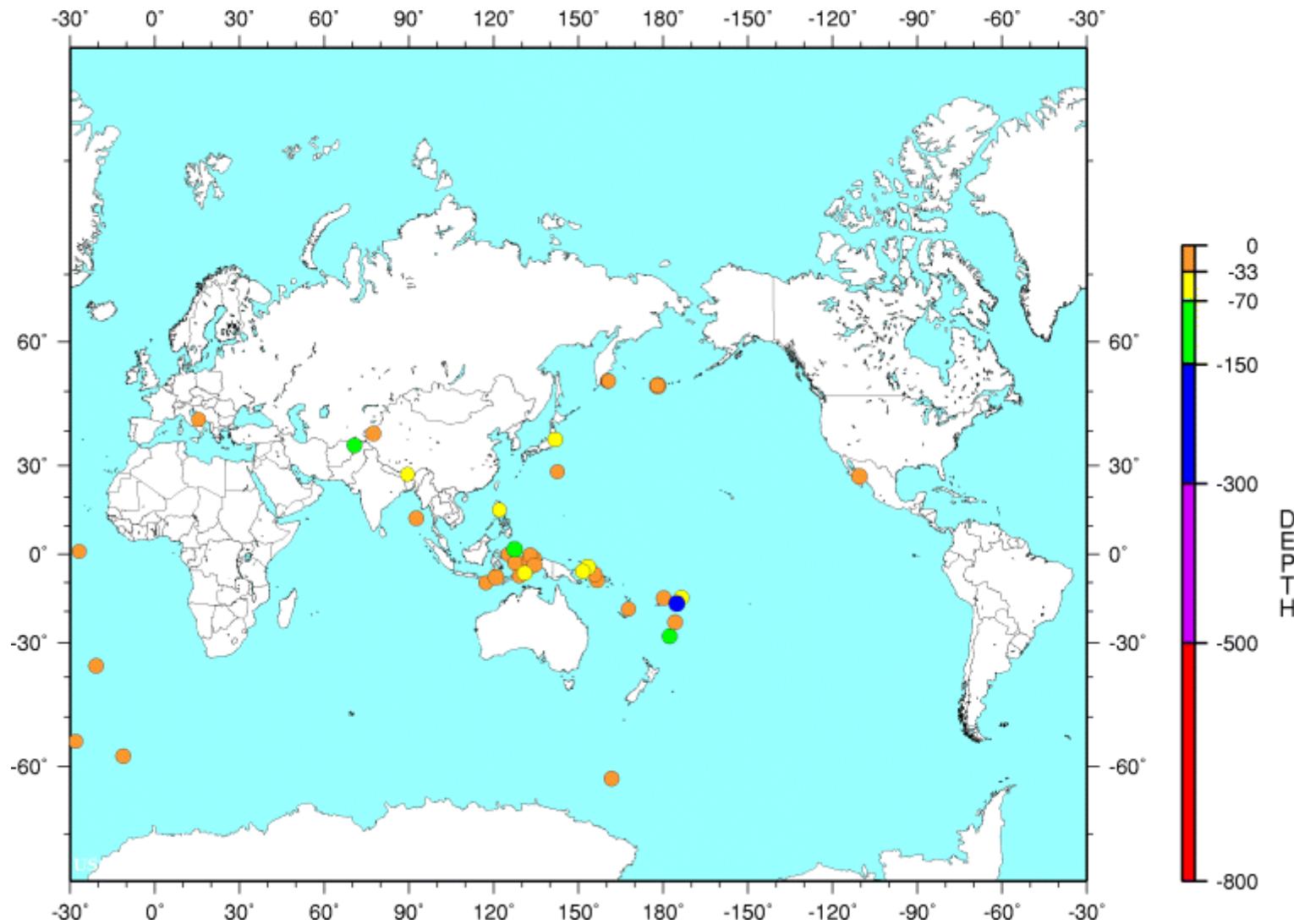
January 2003 : 40 Eq



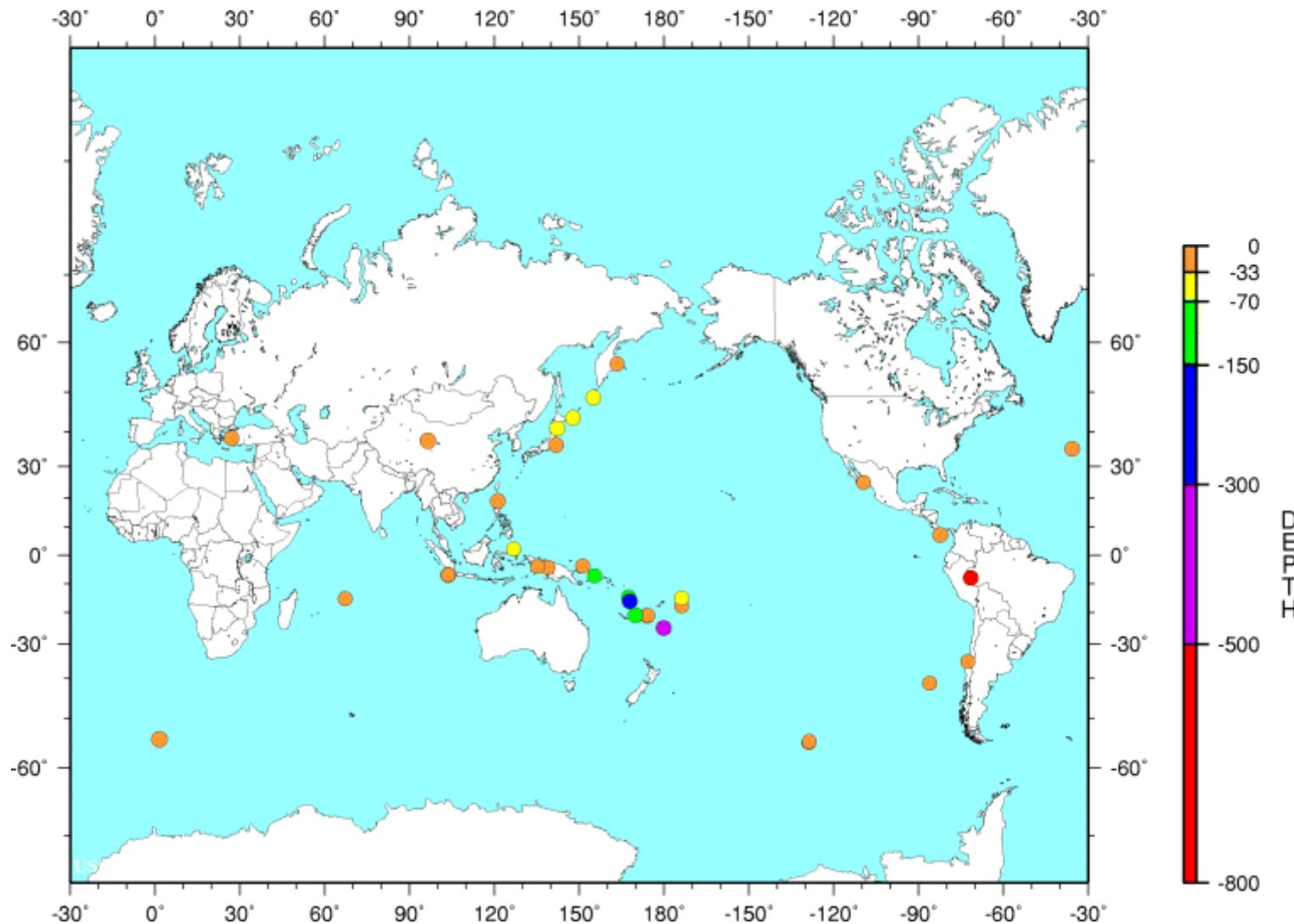
February 2003 : 29 Eq



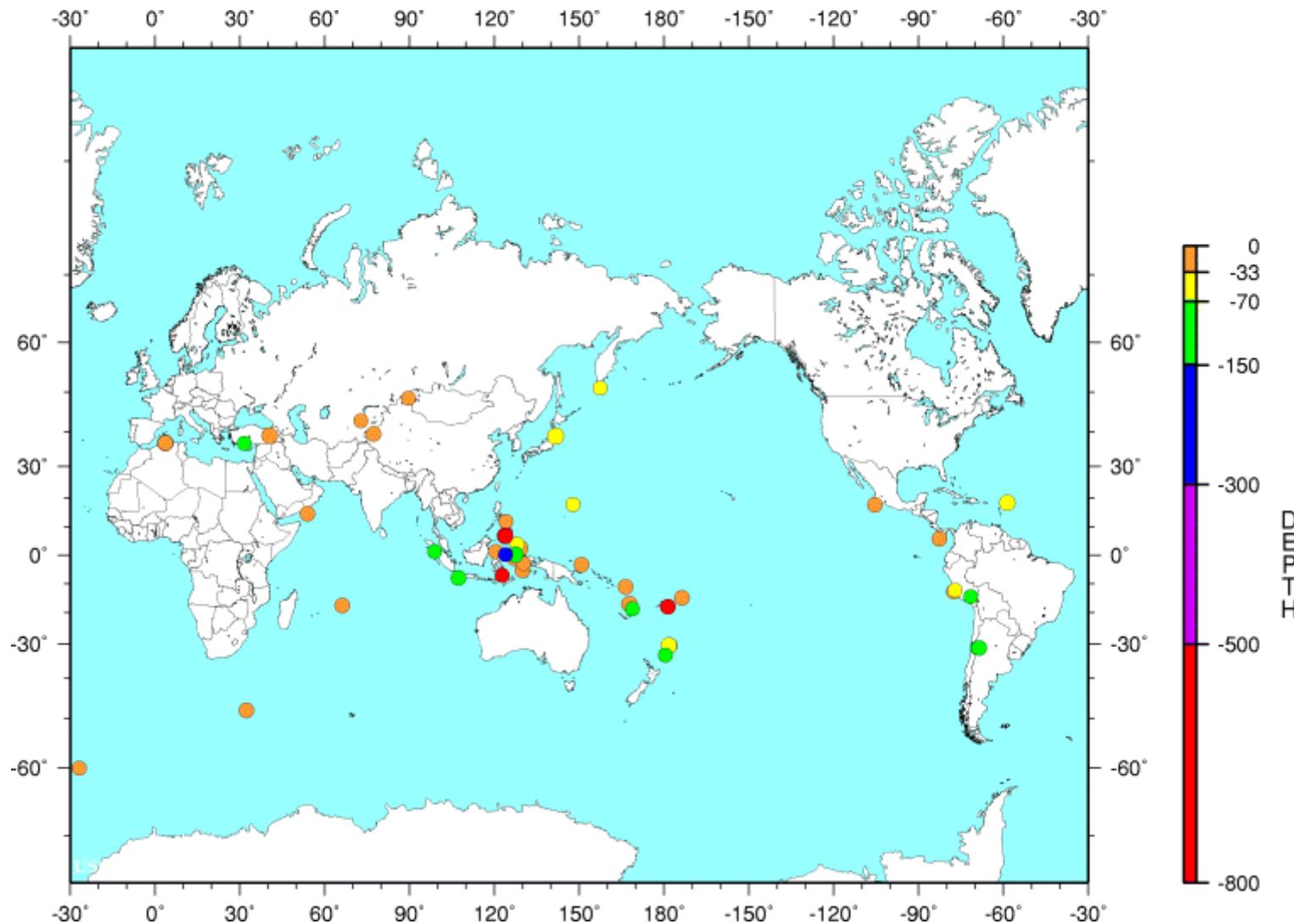
Mars 2003 : 41 Eq



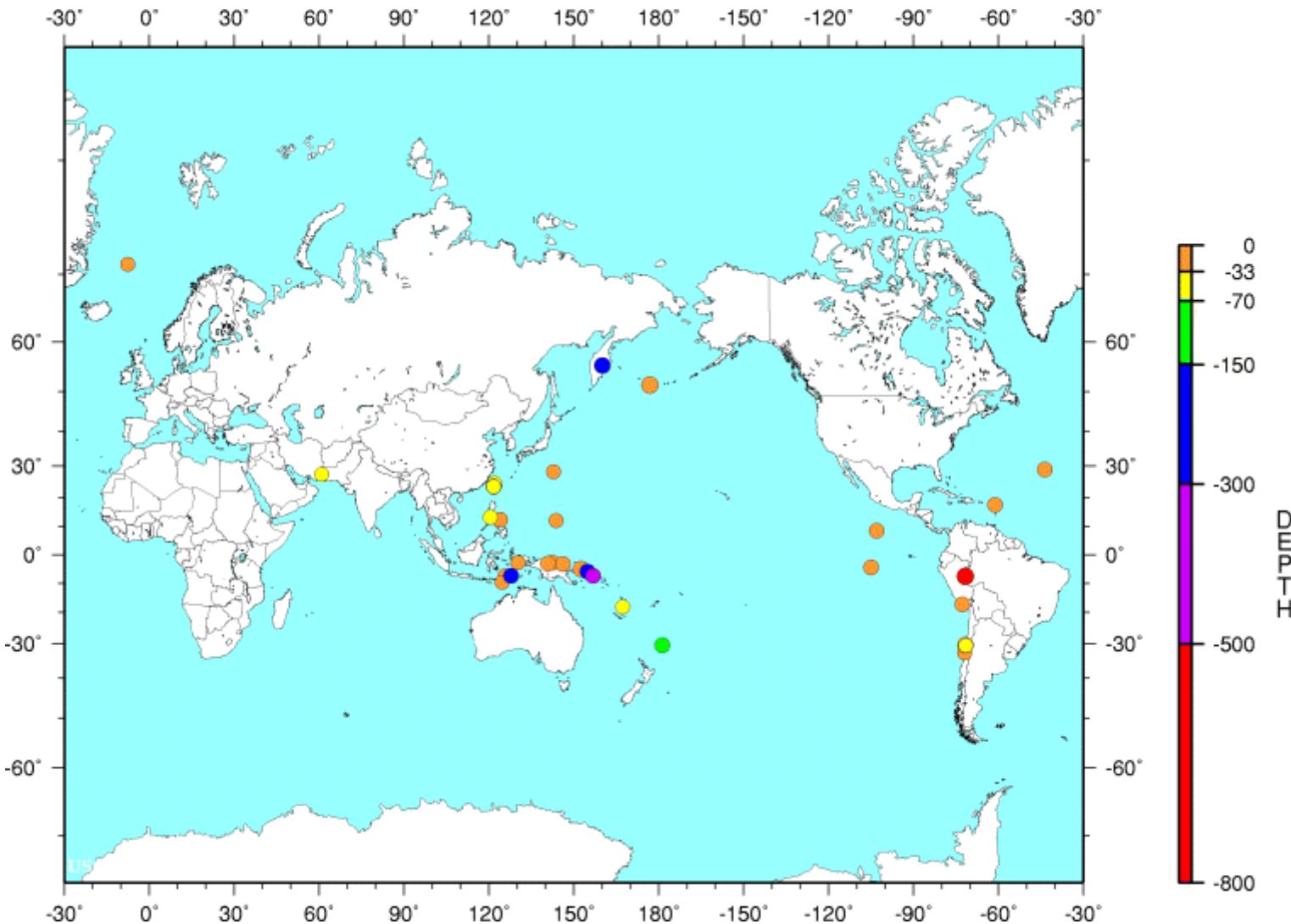
April 2003 : 33 Eq



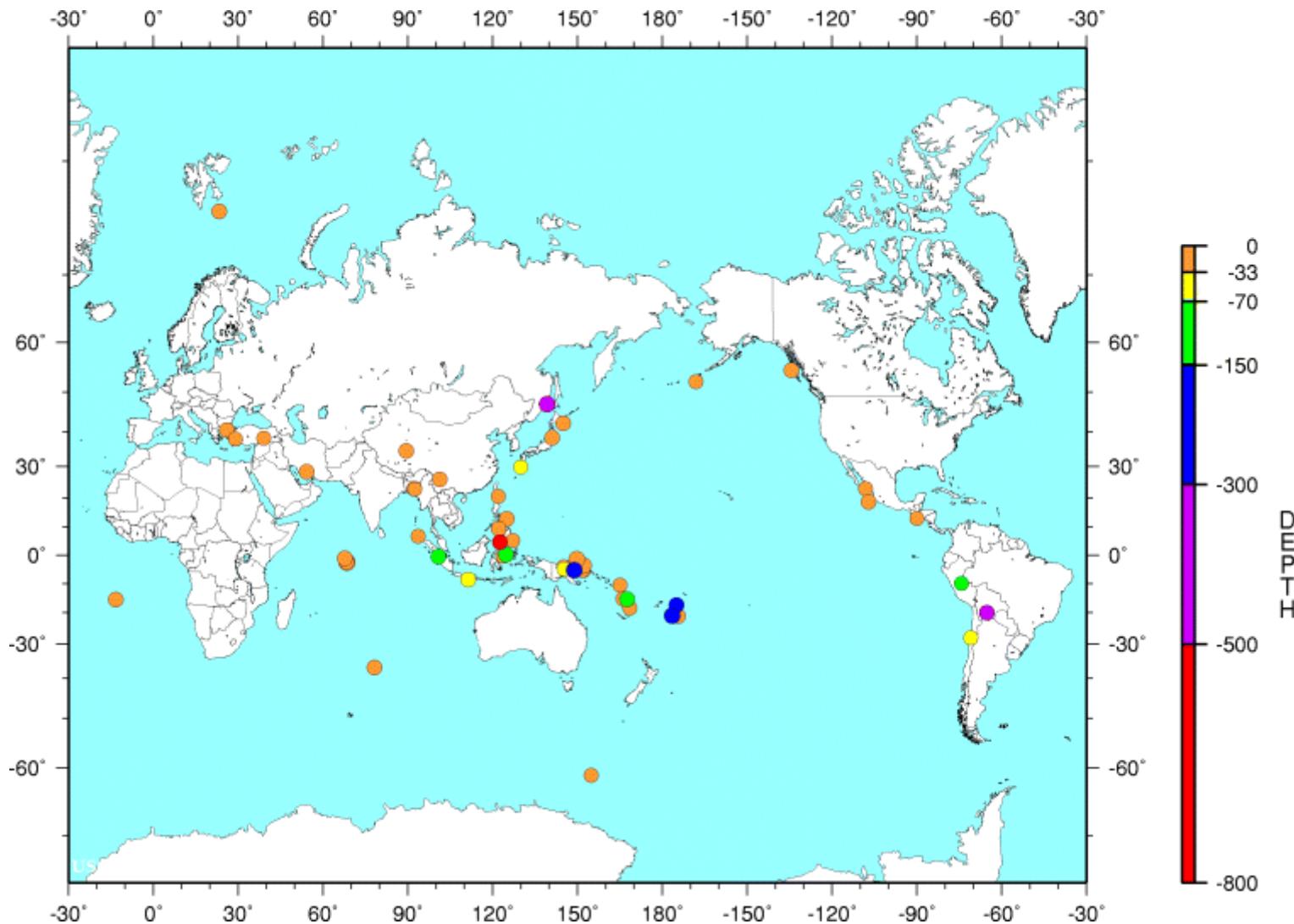
May 2003 : 46 Eq



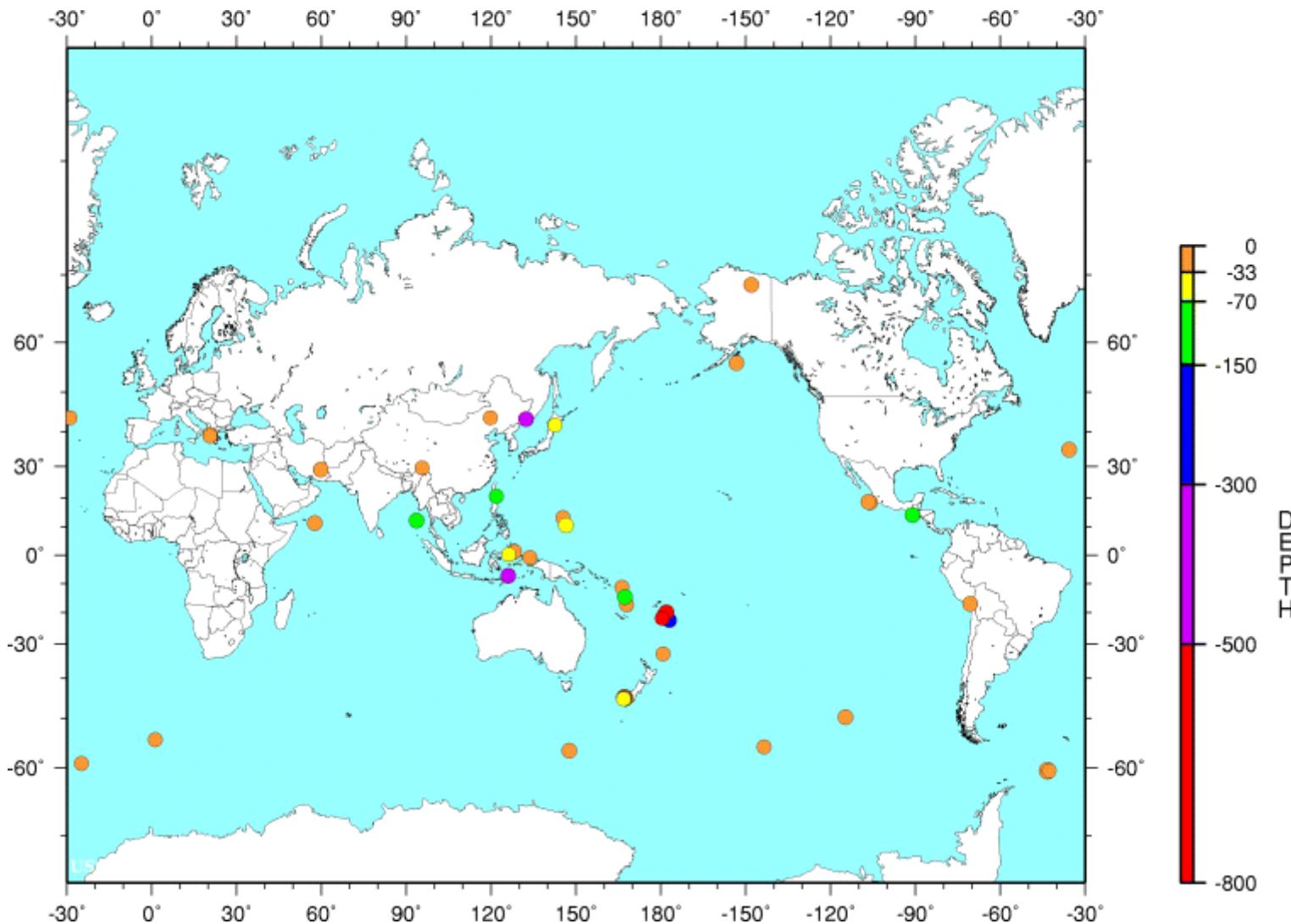
June 2003 : 33 Eq



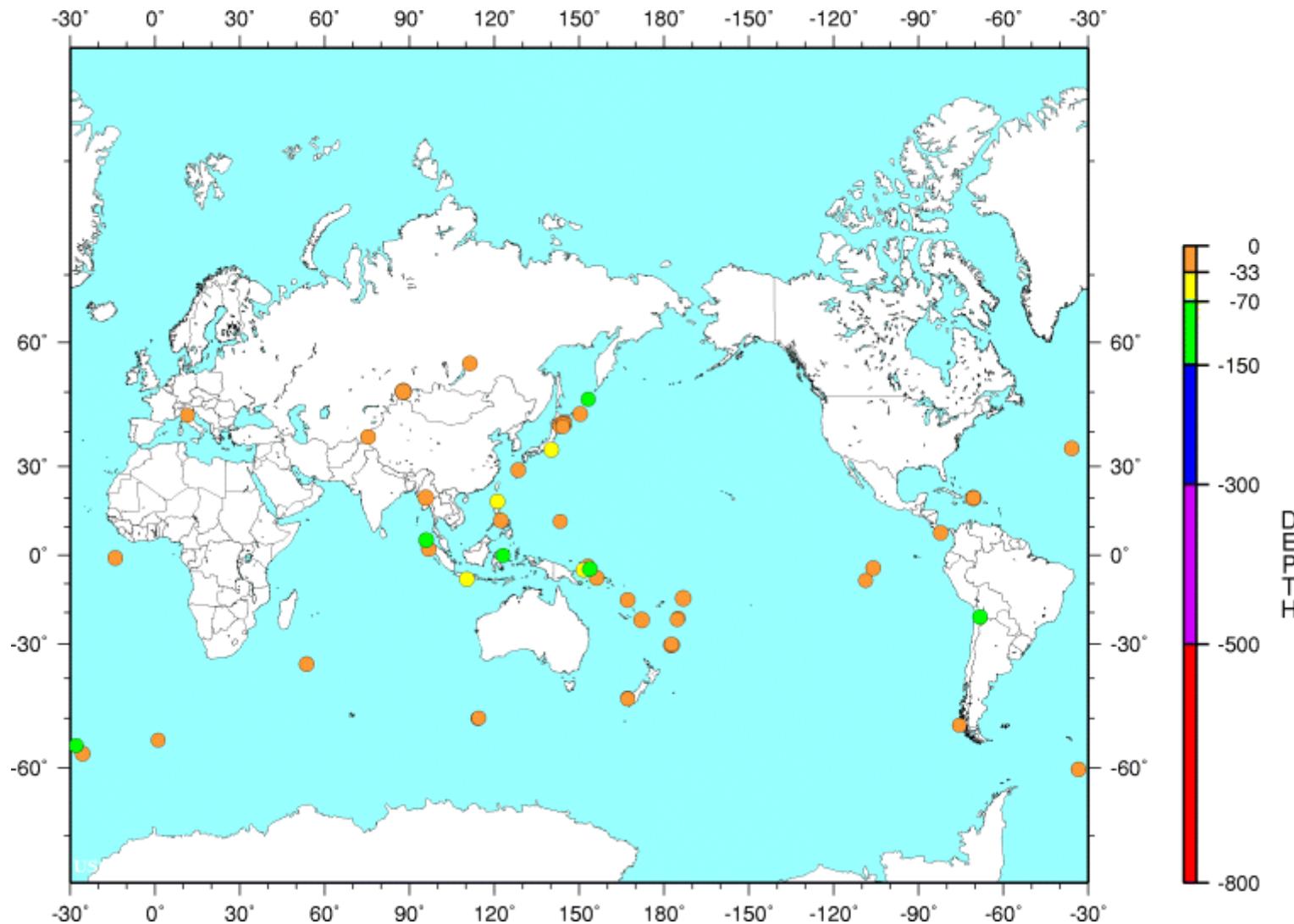
July 2003 : 53 Eq



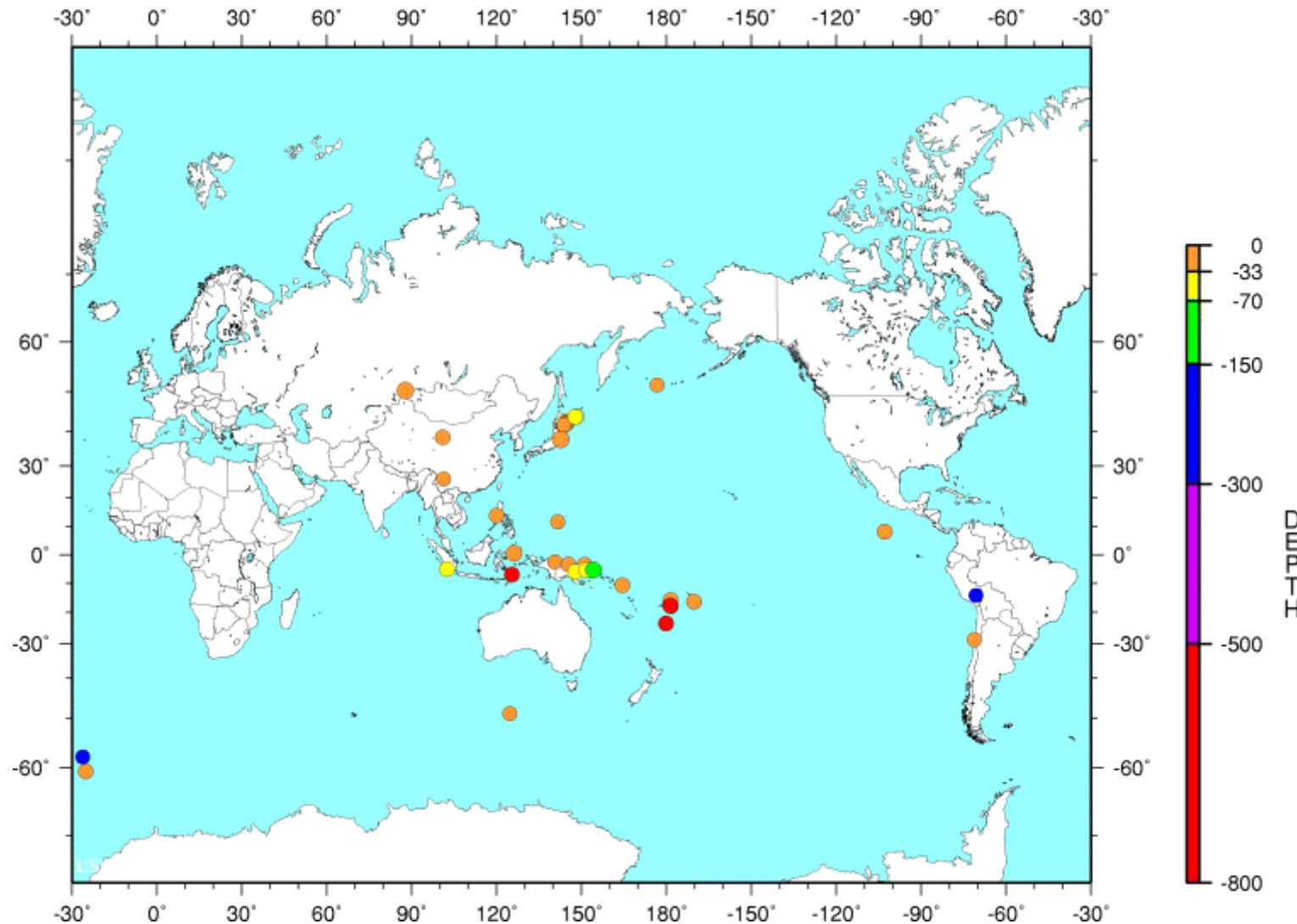
August 2003 : 47 Eq



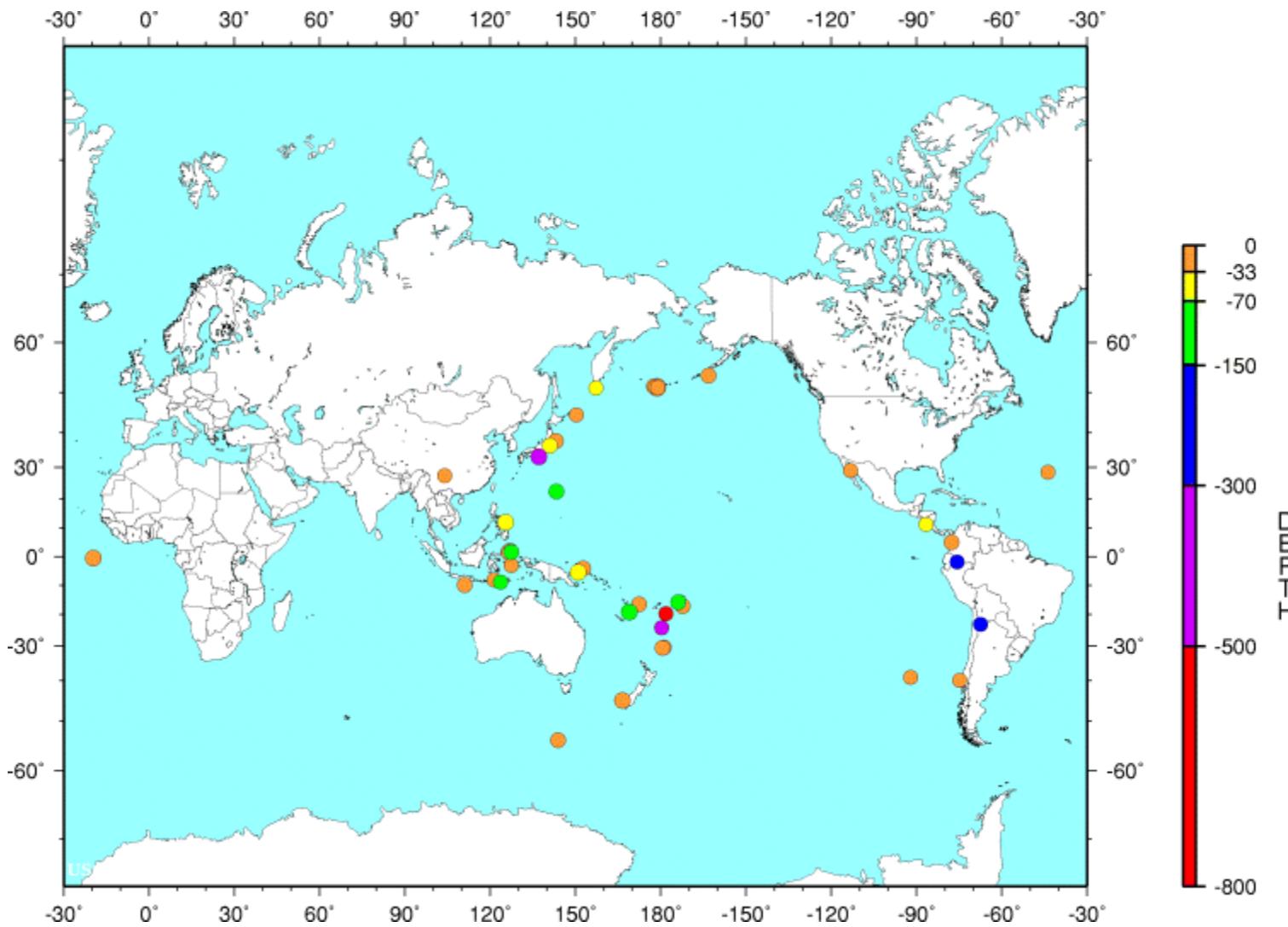
September 2003 : 63 Eq



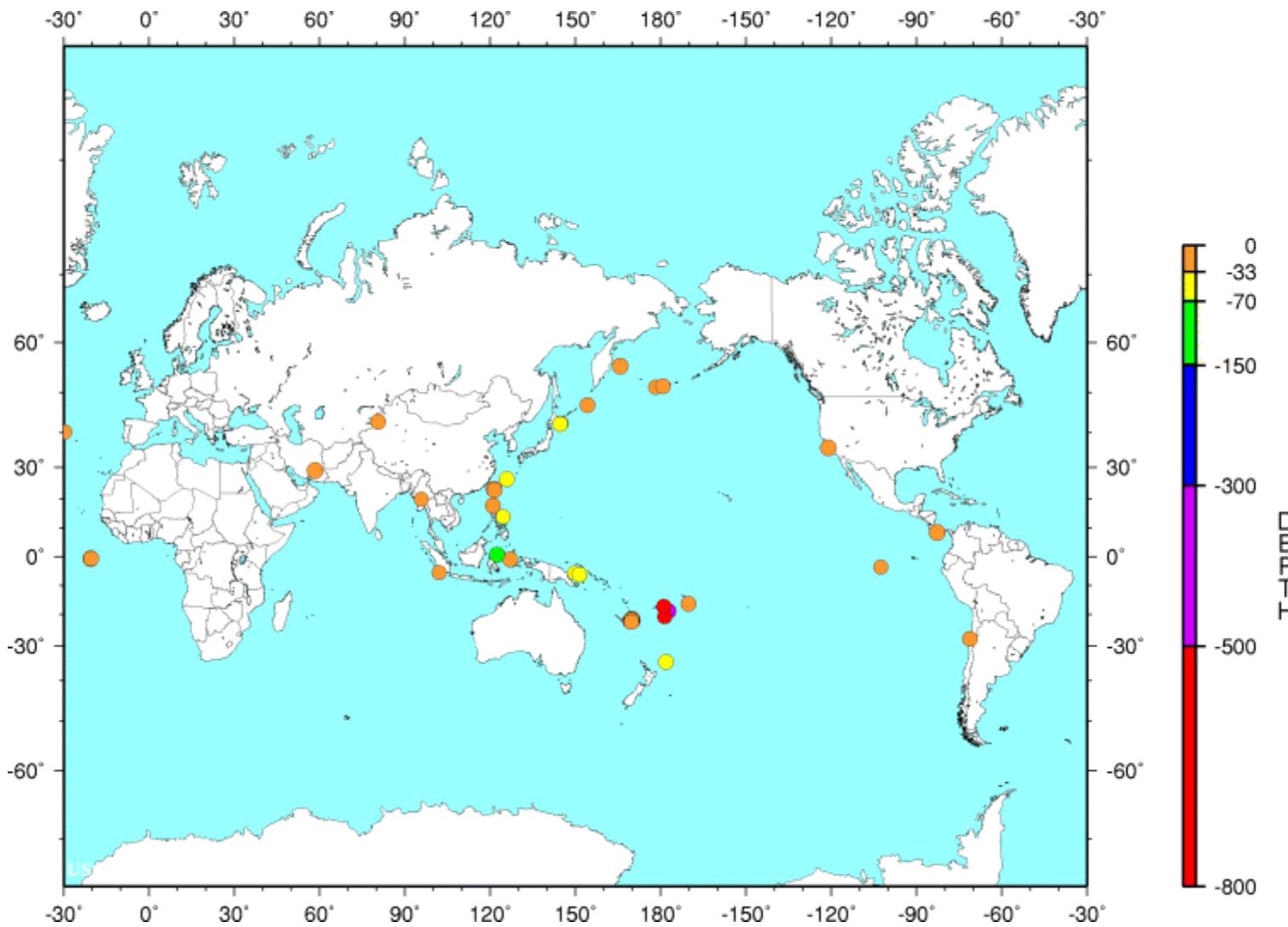
October 2003 : 38 Eq



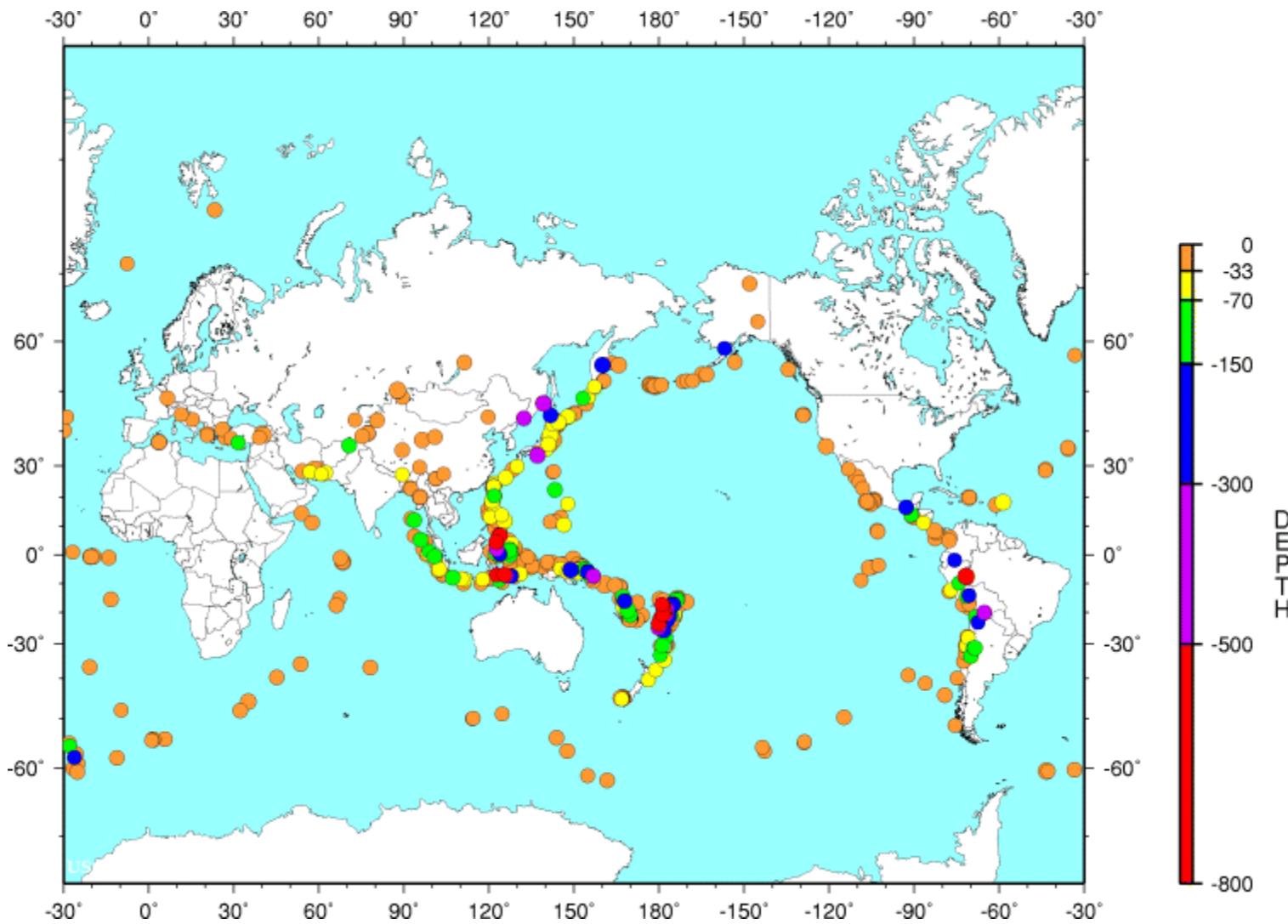
November 2003 : 43 Eq



December 2003 : 46 Eq



YEAR 2003 : 512 Eq



Q1 : HOW MANY EARTHQUAKES ?

ANSWER : it really depends on size !

small (magnitude between 4 and 5) : **1 / hour**

medium (magnitude between 5 and 6) : **1 / day**

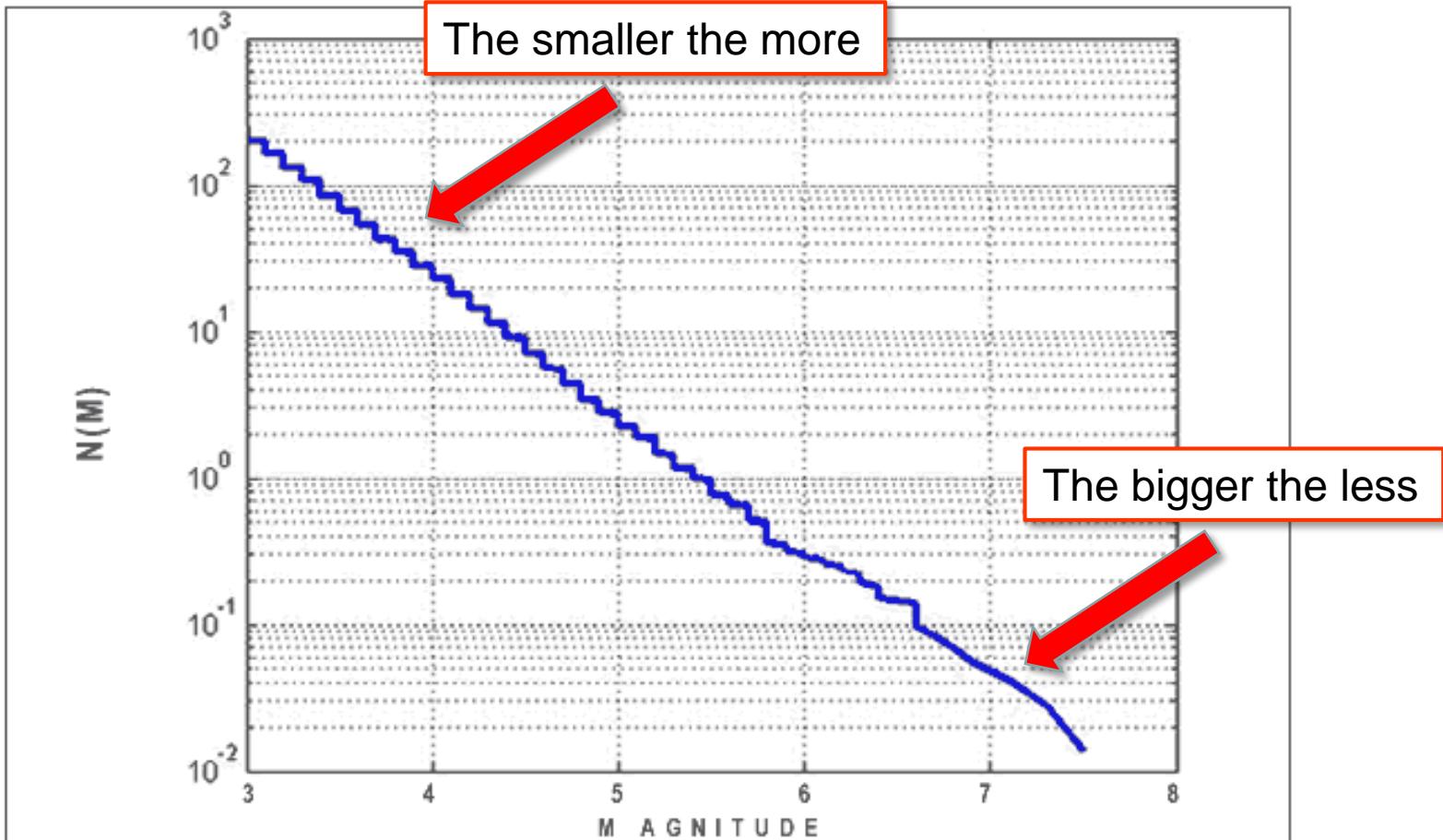
big (magnitude between 7 and 8) : **1 / month**

very big(magnitude between 8 and 9) : **1 / year**

giants (magnitude > 9) : **1 every decade**

Q1 : HOW MANY EARTHQUAKES ?

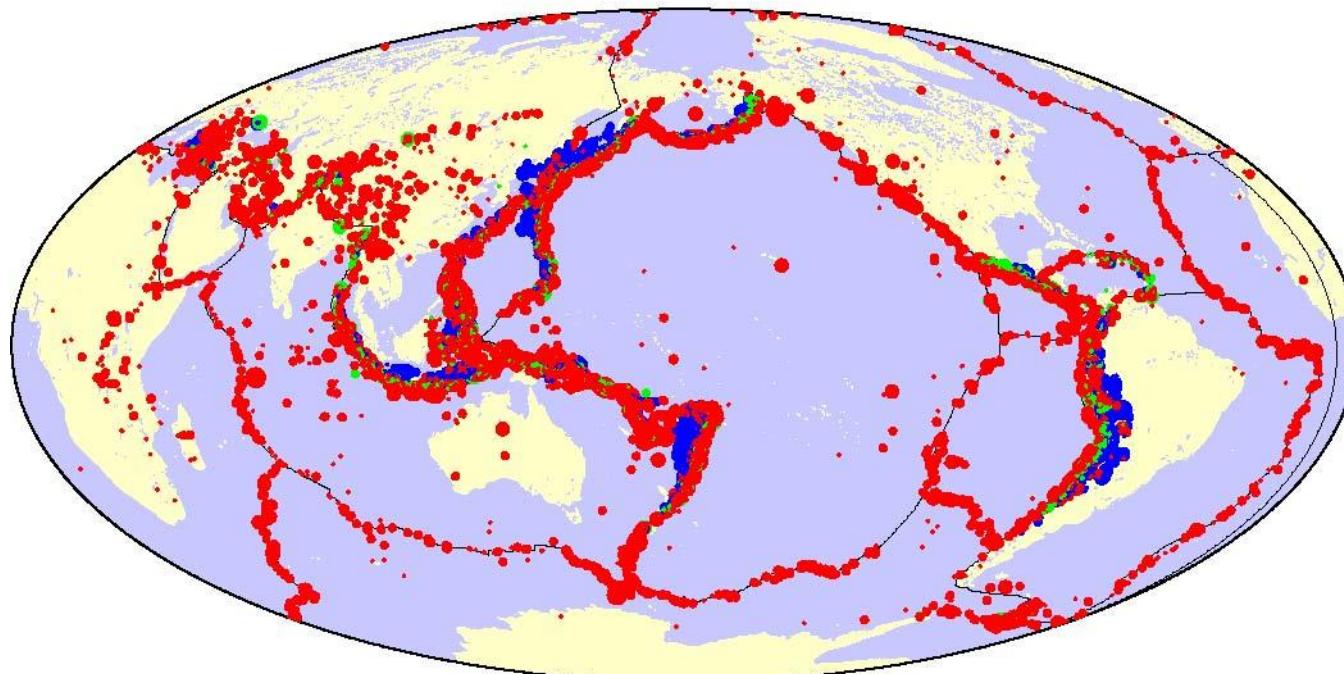
Gutenberg & Richter law



Q2 : WHERE DO EARTHQUAKES OCCUR ?

ANSWER : Not everywhere !

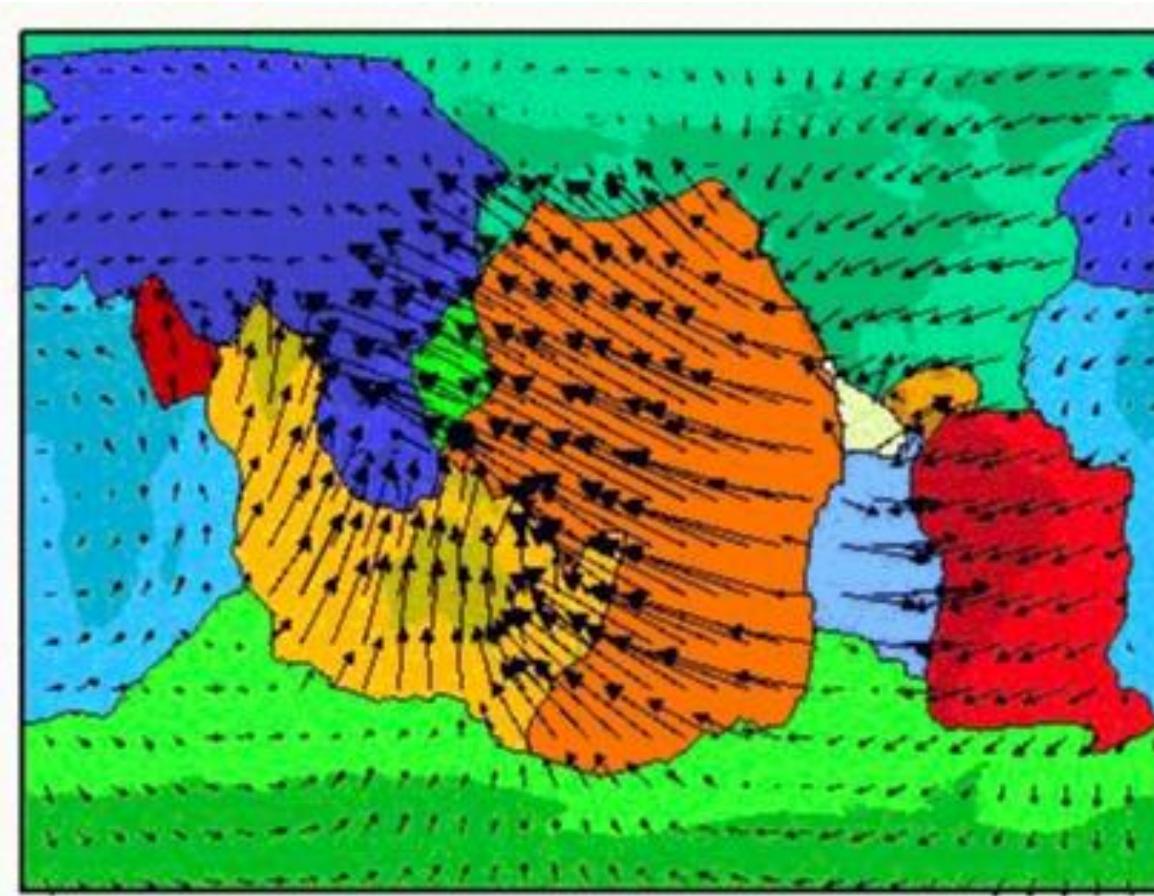
World seismicity : 30 years – 17 000 events of magnitude larger than 4.5



- PETITS : M 4.5 - 5.5 • MOYENS : M 5.5 - 6.5 • GROS : M 6.5 - 7.5 • TRES GROS : M 7.5 - 8.5
- profondeur < 50km • profondeur 50km - 100km • profondeur > 100km

Q3 : WHY DO EARTHQUAKES OCCUR ?

ANSWER : mostly(*) because of plate tectonics



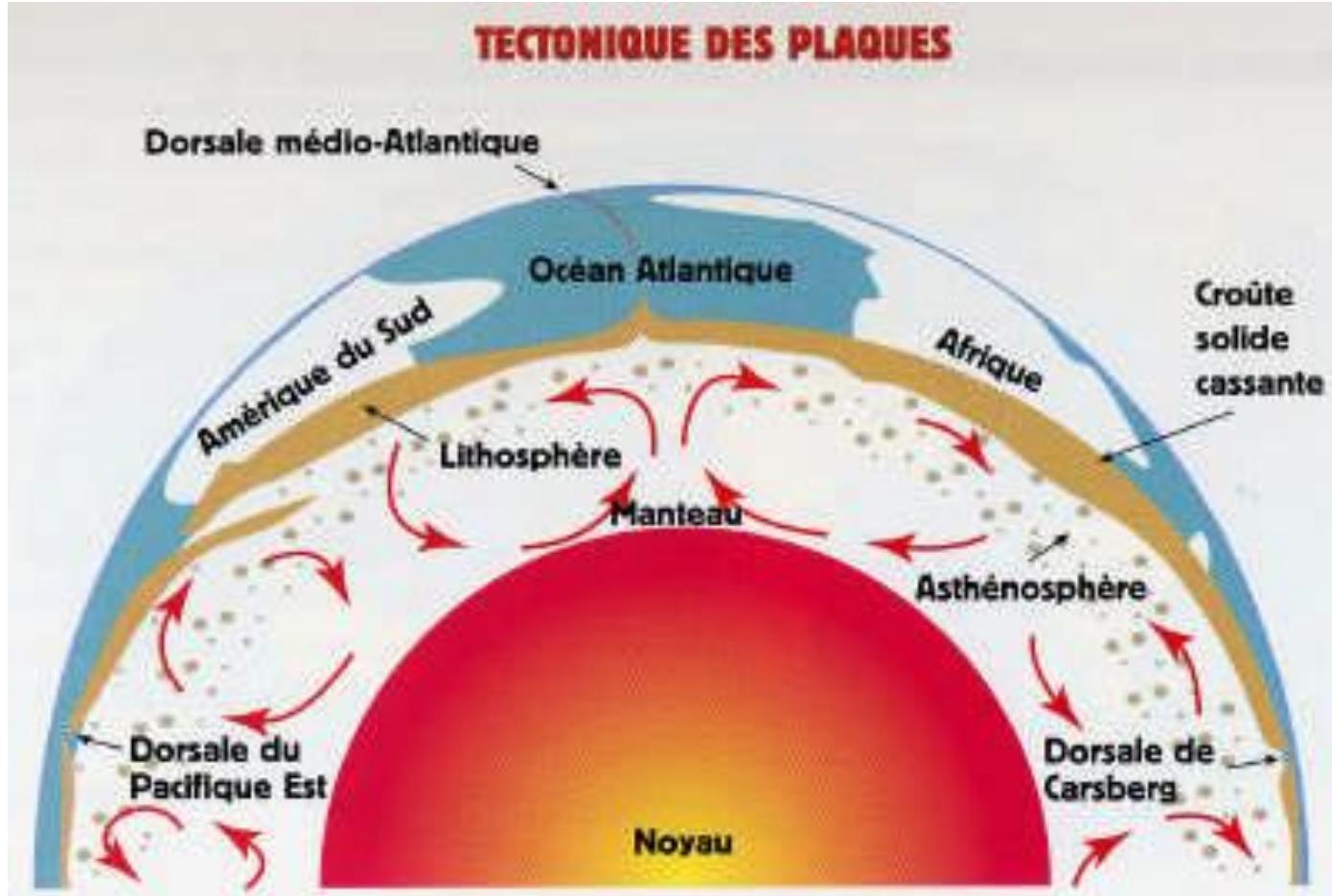
Pacifique
Eurasie
Amérique du N
Amérique du S
Afrique
Antarctique
Indie-Australie
Nazca
Philippine
Arabie
Caraibes
Coco

There are 12 major plates

...and they move ! Plate tectonics

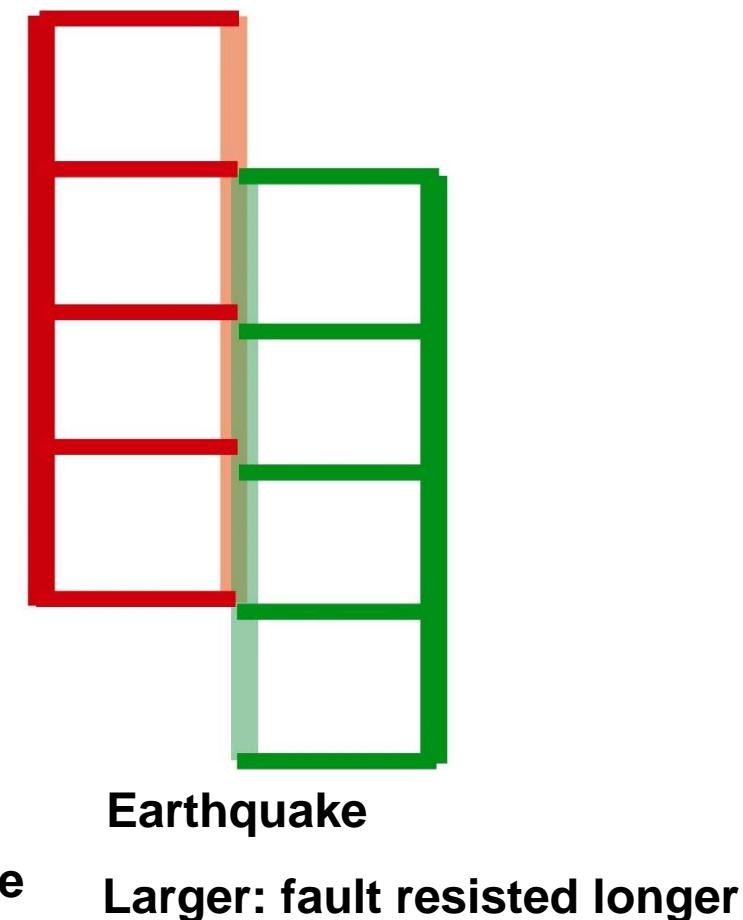
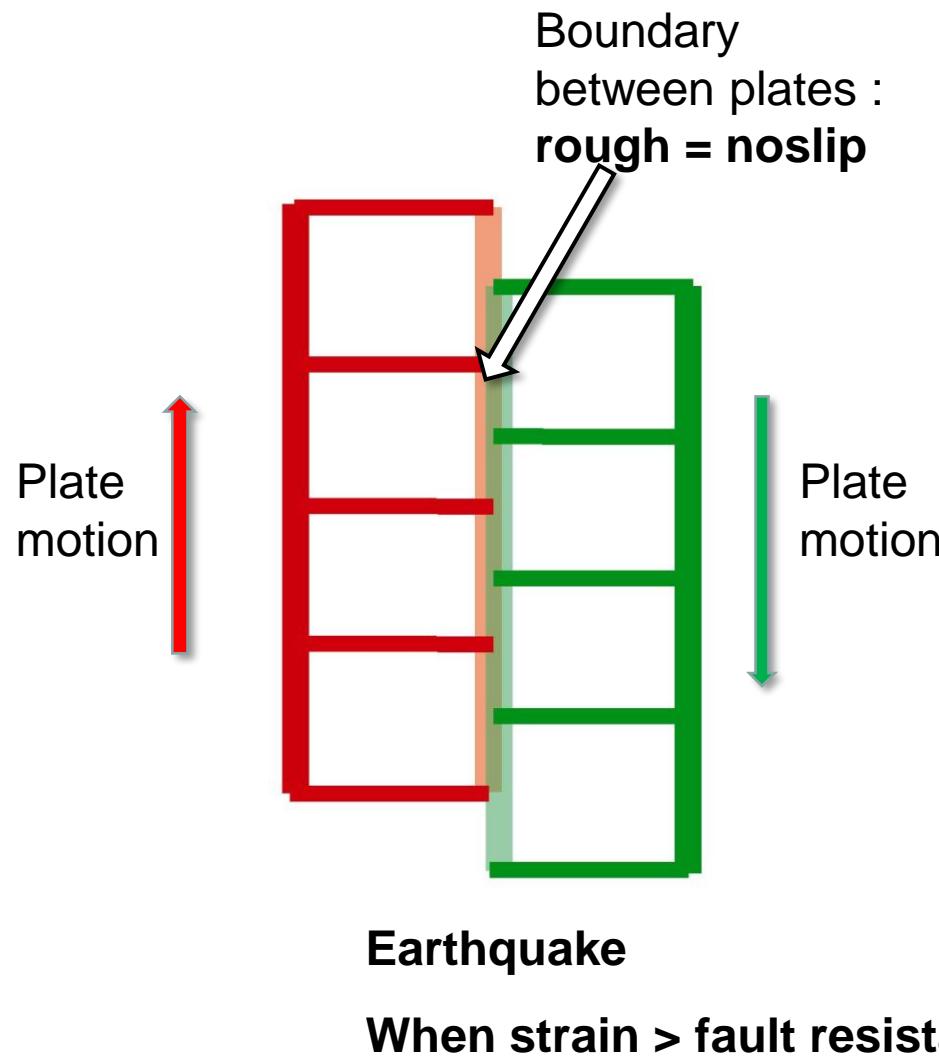
* Will get back to that later

The engine of plate tectonics is mantle convection
its fuel is rocks natural radio-activity



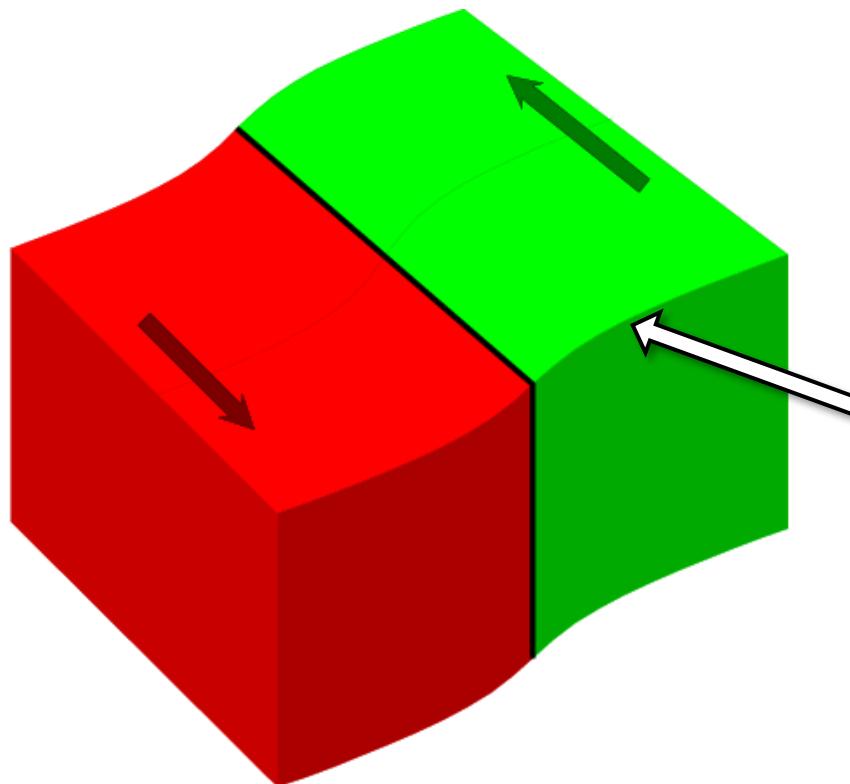
Q4 : HOW DO EARTHQUAKES OCCUR ?

Answer : because of friction, It rubs and it brakes !



Q5 : ANTICIPATE EARTHQUAKE size ?

Answer : possible, measure deformation before...

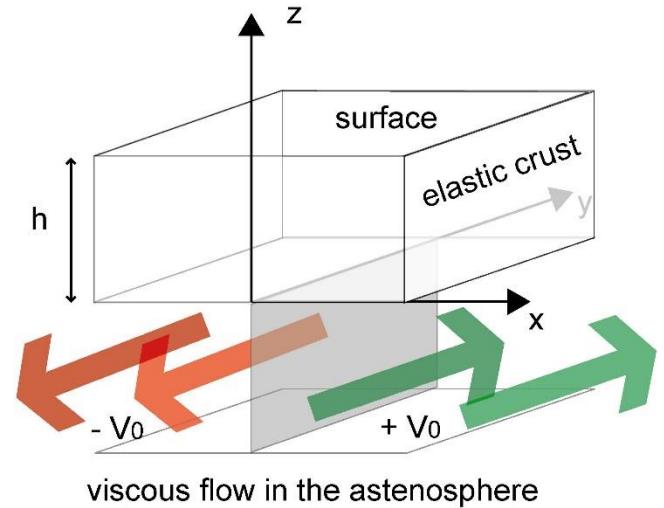


This curve
is not any
curve

Inter-seismic deformation obeys elastic medium formulation

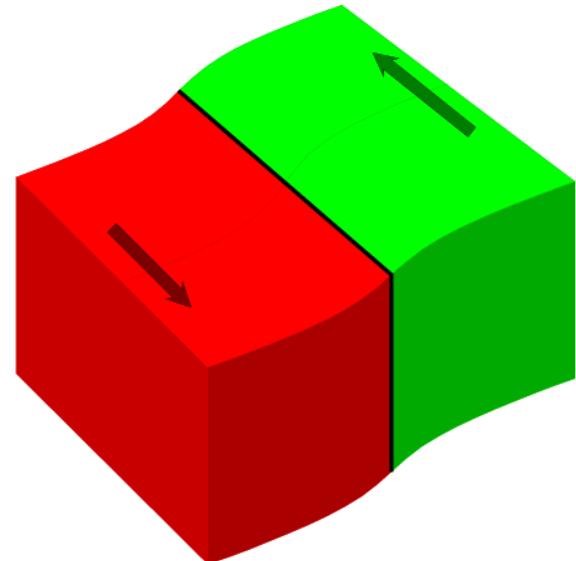
Mathematical formulation : screw dislocation model

$$U_y = K \operatorname{arctan} (x/z)$$



At the surface ($z=h$)

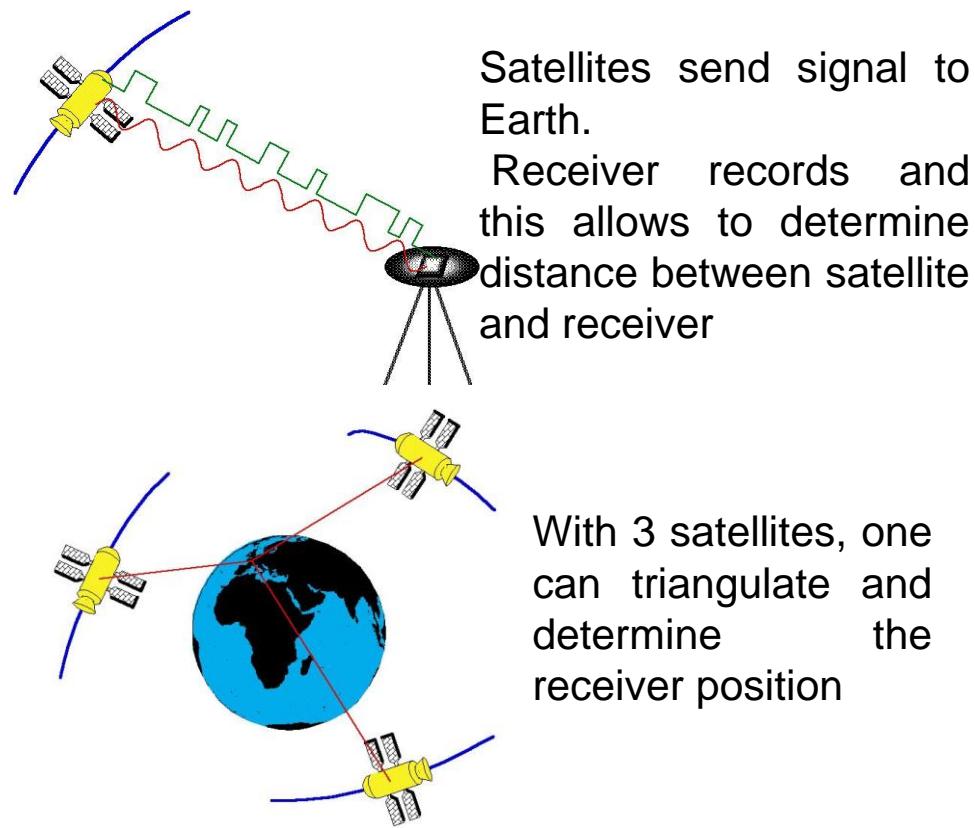
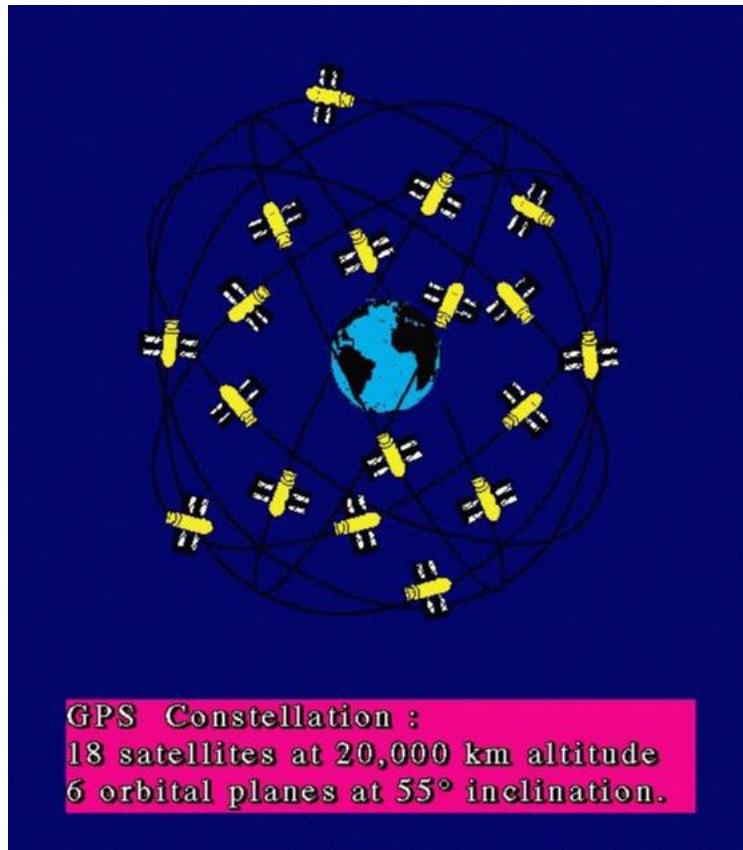
$$U_y = 2 \cdot v_0 / \pi \operatorname{arctan} (x/h)$$



Q6 : HOW MEASURE pre earthquake deformation ?

Answer : Geodesy, now space born : GPS

GPS was created in the 80s' by the US Department of Defense for military purposes. The objective was to be able to get a precise position anywhere, anytime on Earth.



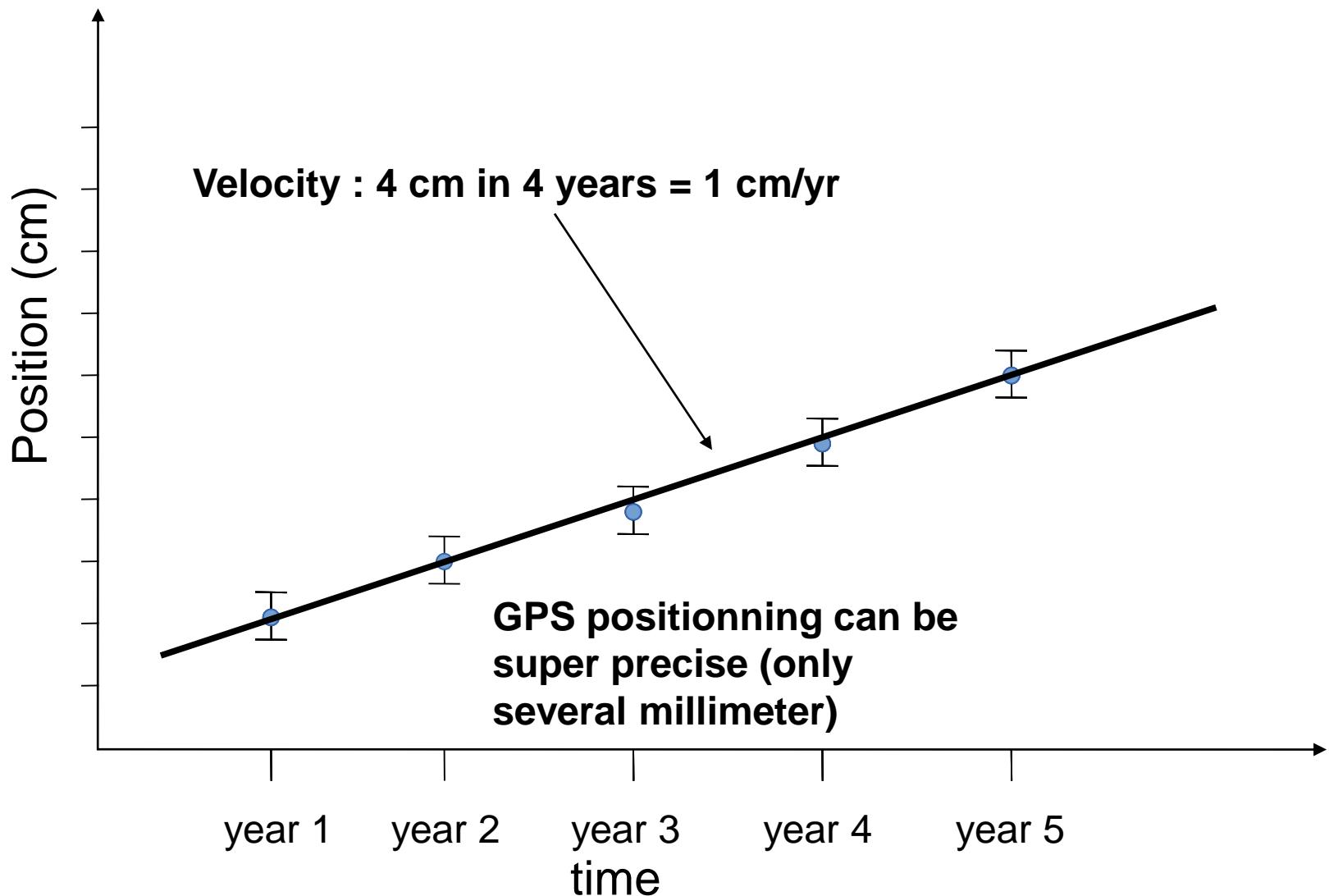
GPS (Global Positioning System)



GPS antenna on tripod



**Geodanic
marker**



GPS (Global Positioning System)

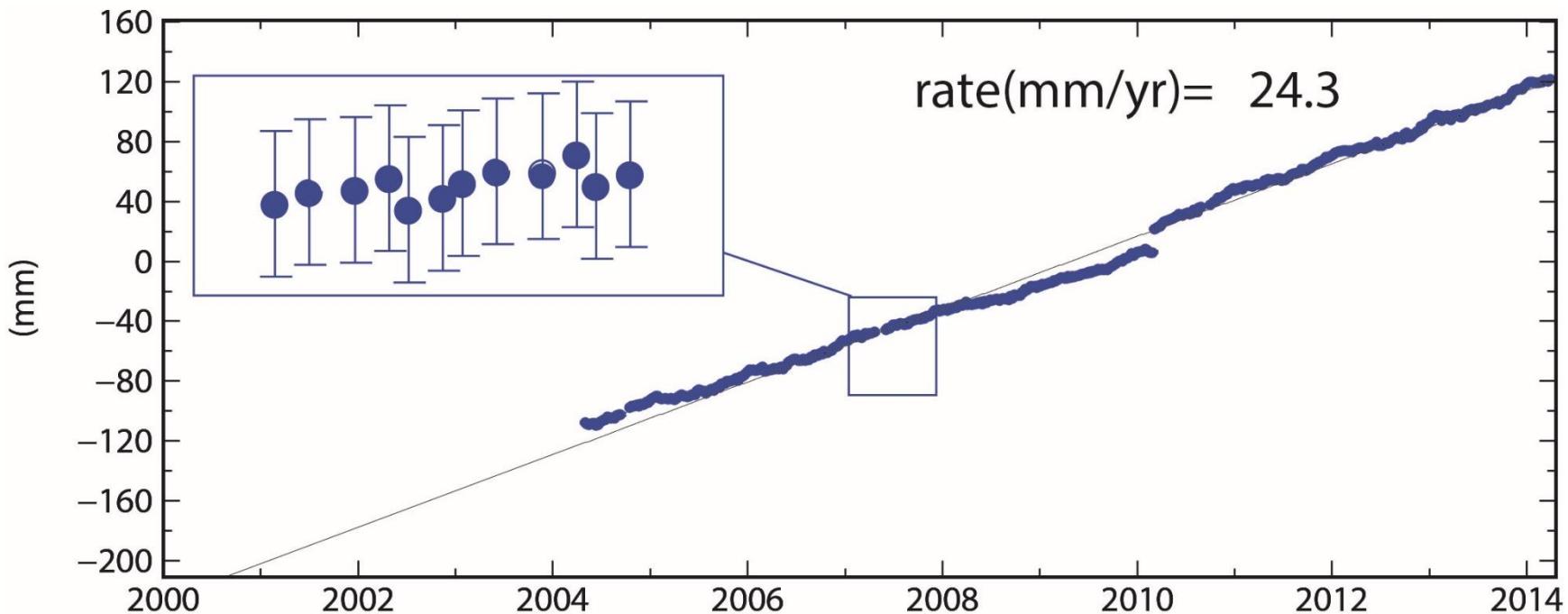


Permanent stations

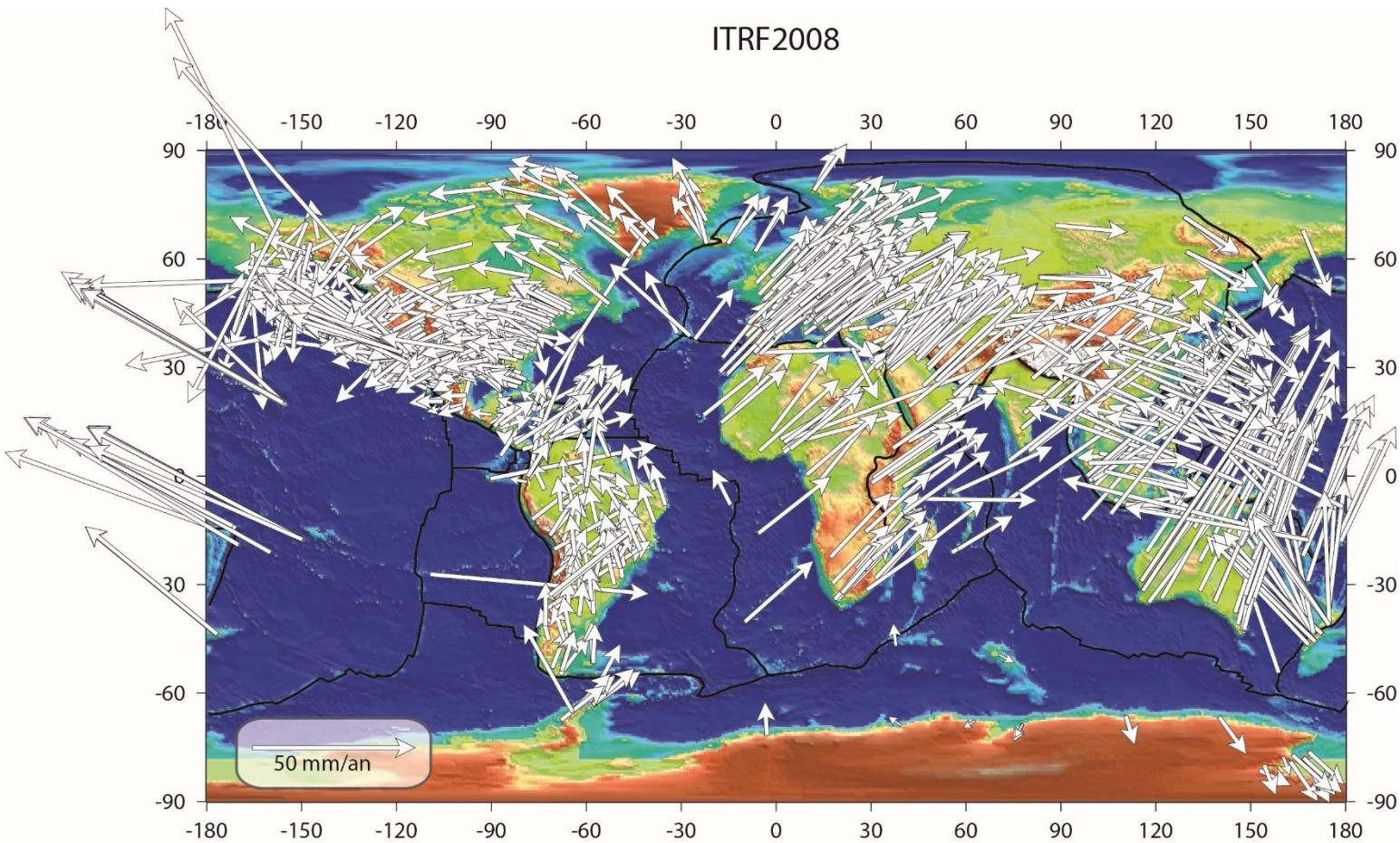


10

Continuous GPS (cGPS) : time series

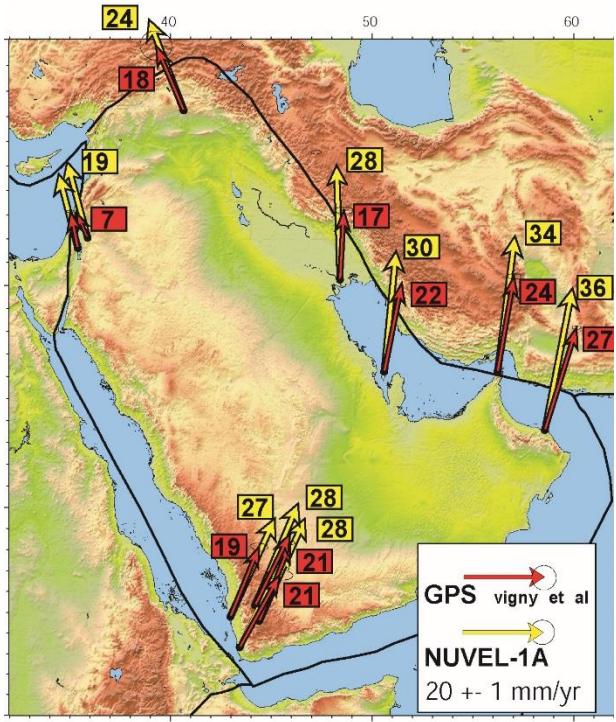


A spectacular application of GPS: Measurement and quantification of plate tectonics

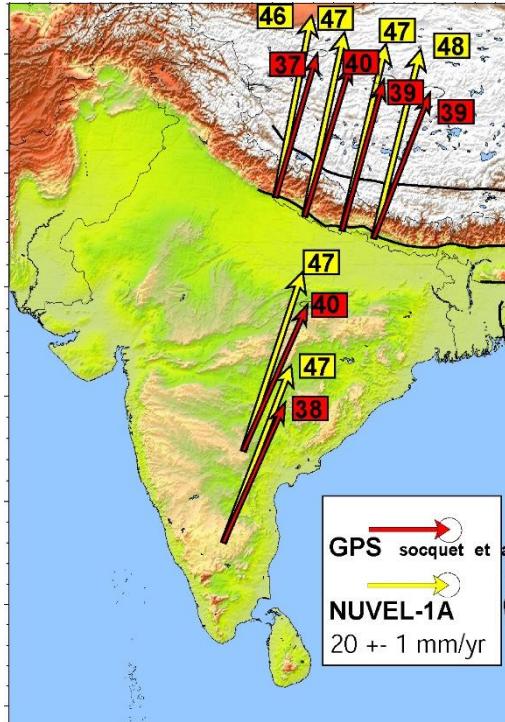


**GPS velocities compare well with “geological” velocities: ITRF = Nuvel
⇒ plate motions are stable : 10 km in 1 Myr = 1 cm/yr**

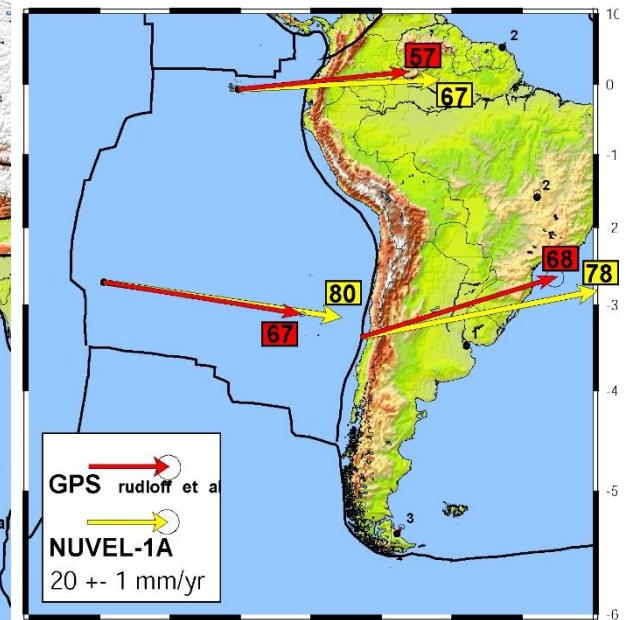
Confirmation of plate tectonics : direct measurement from space



ARABIA

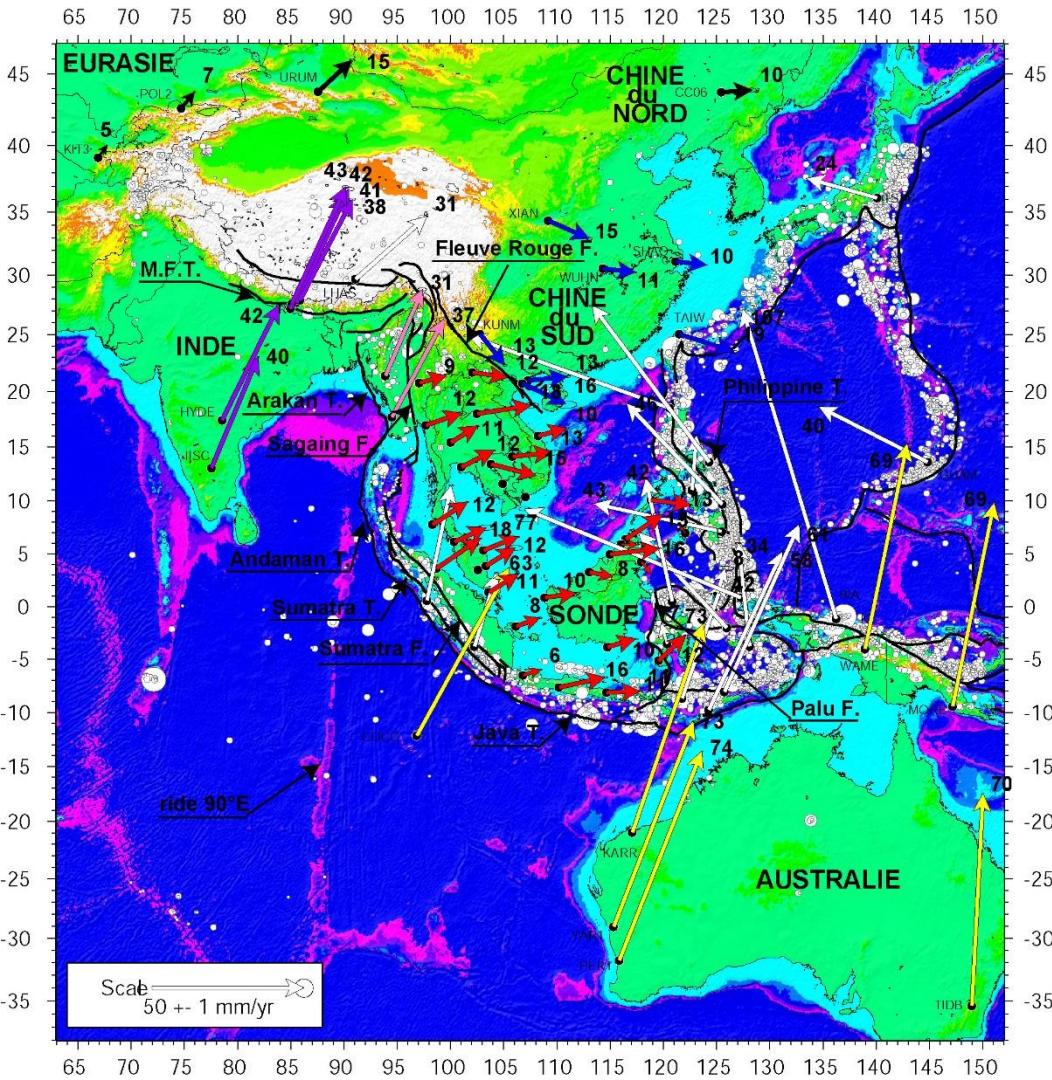


INDIA



SOUTH AMERICA

Discovery of platelets : Sundaland (1997)



SUNDA plate:
60 GPS sites
12 years of measurements

→ independent from Eurasia
and moves 1 cm/yr Eastward

INDIA plate:
6 GPS sites
10 years of measurements

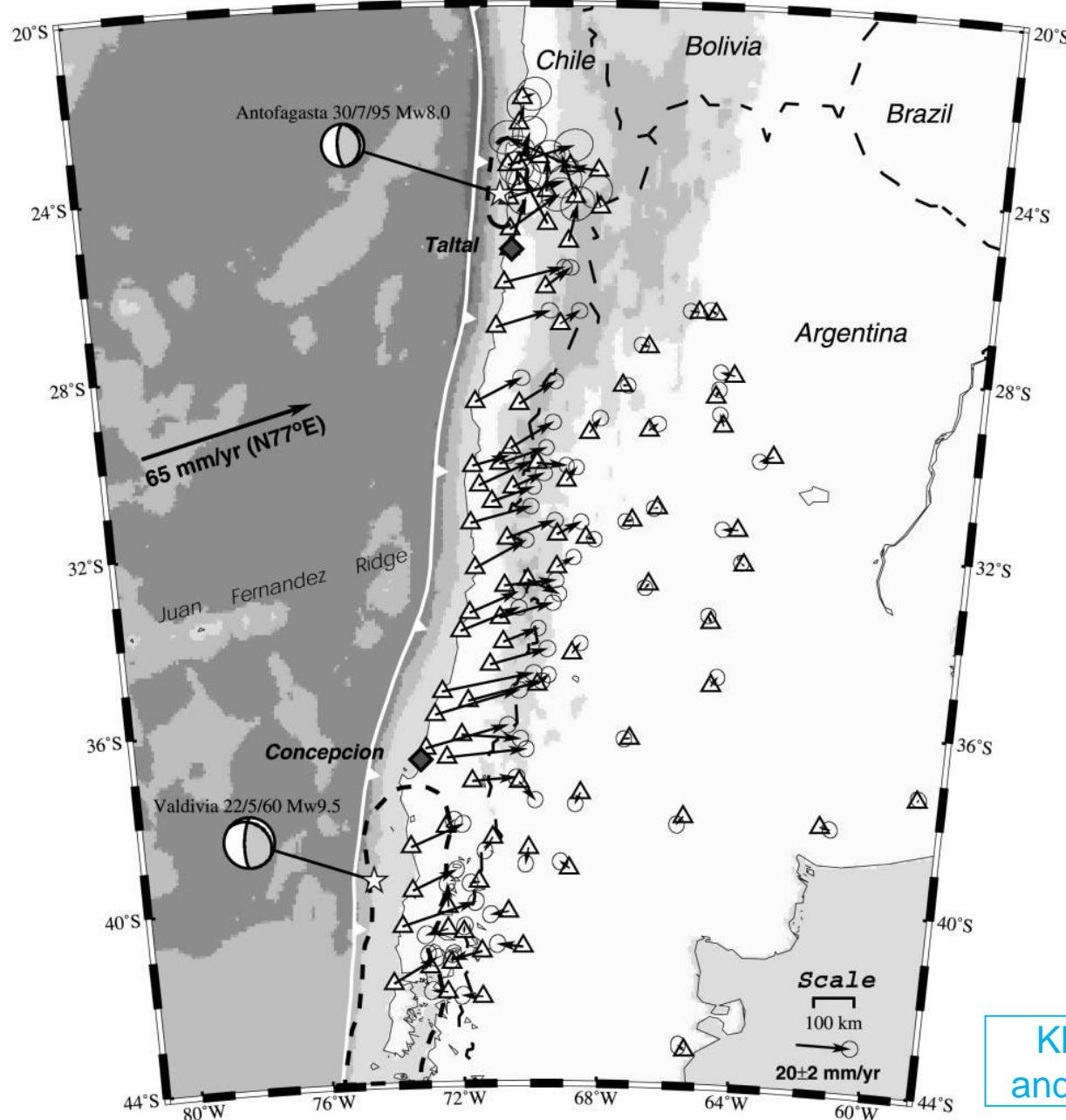
→ only 4 cm/yr wrt GPS Eurasia
even less (3.5 cm/yr) wrt Sunda

BURMA platelet (or sliver):
22 GPS sites
2 years of measurements

→ Nor India, nor Eurasia, not even
Sunda

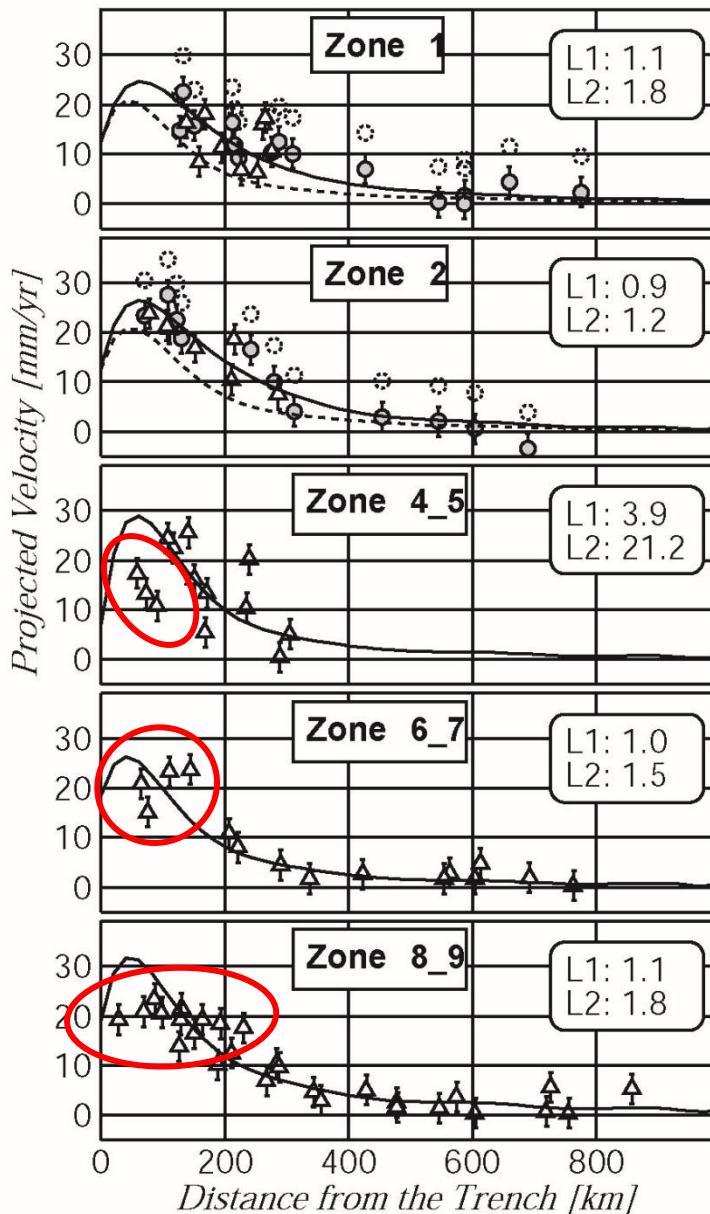
South China ????

SAGA network (GFZ-Potsdam) since ~1990



Klotz, Khazadze
and co-authors, 2002

Klotz, Khazaradze and co-authors

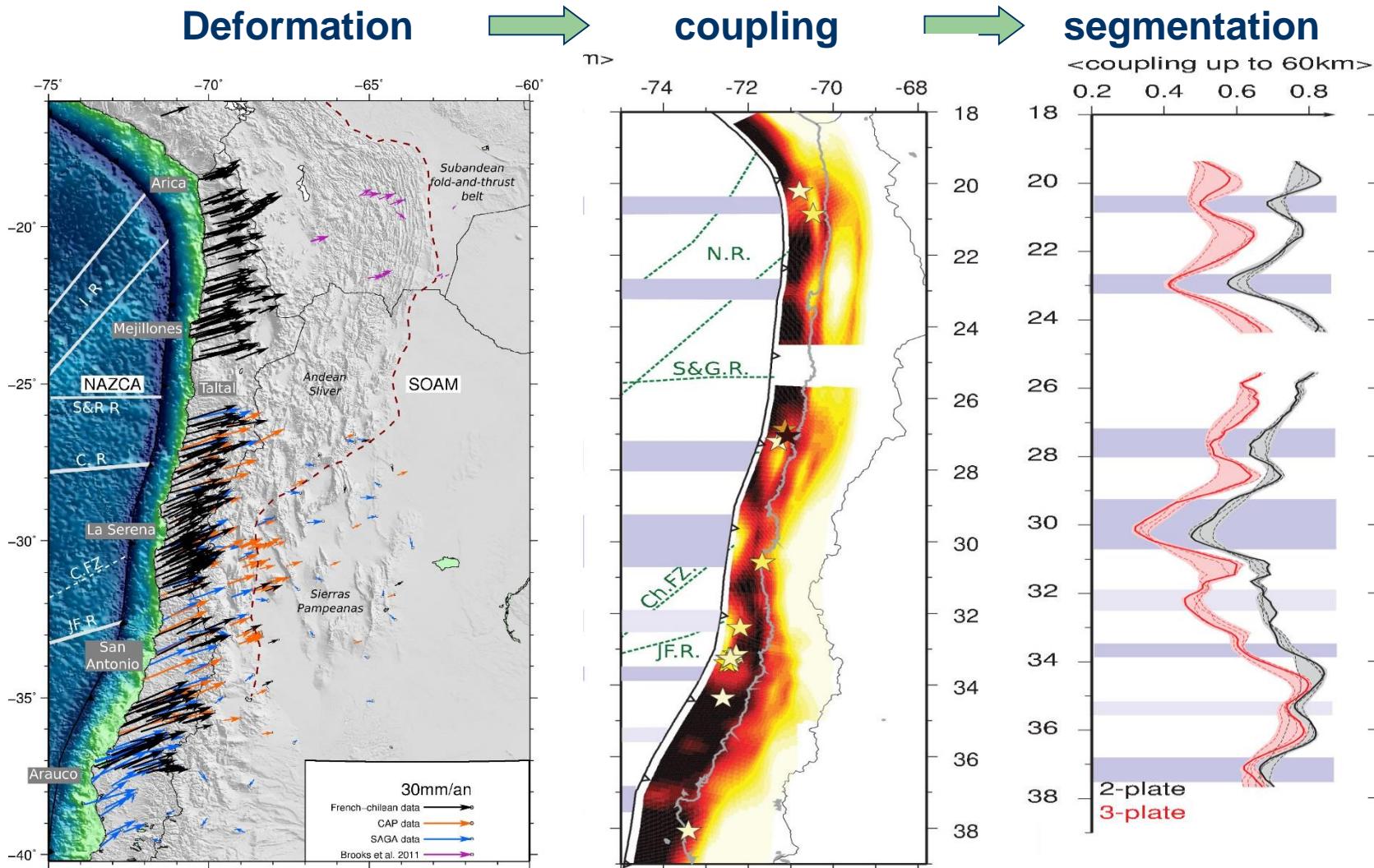


Curves show elastic deformation prediction according to standard model (Full Coupling, ...)

But

Intra-plate deformation during inter-seismic loading...

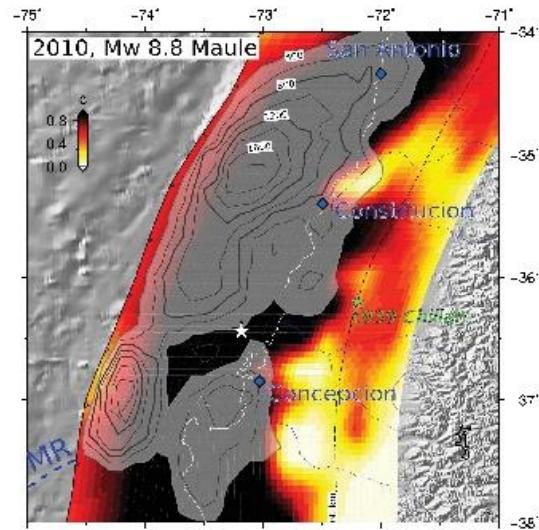
25 years of GPS measurements in Chile => coupling maps



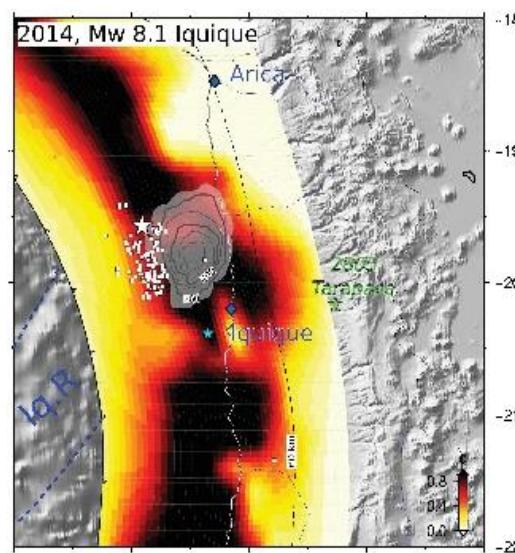
FINDING1 : Earthquakes occur in high coupling areas

interesting : epicenters seem located in lower coupling areas....

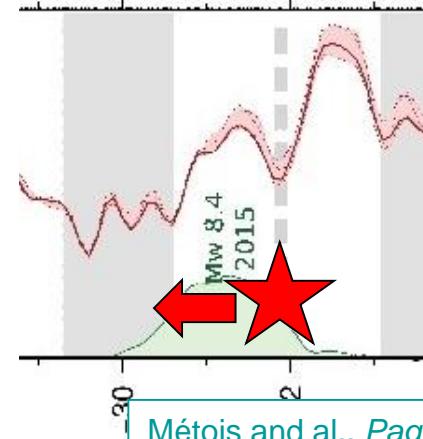
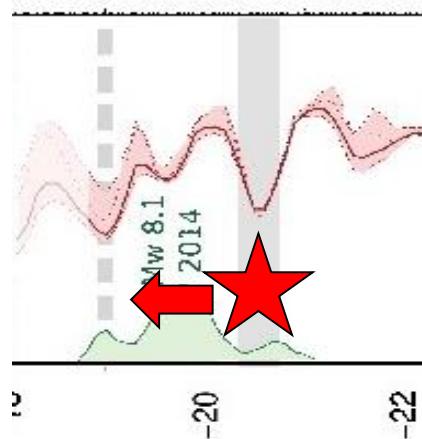
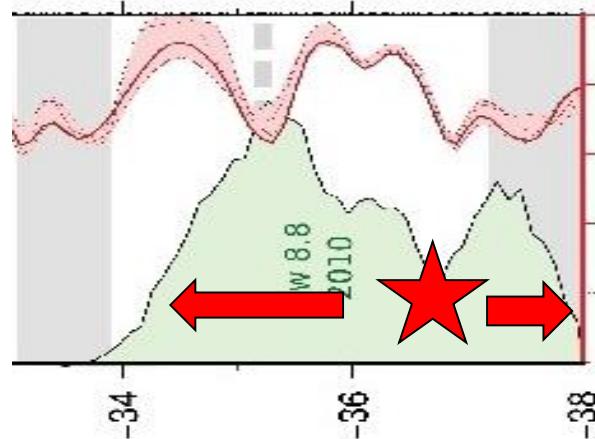
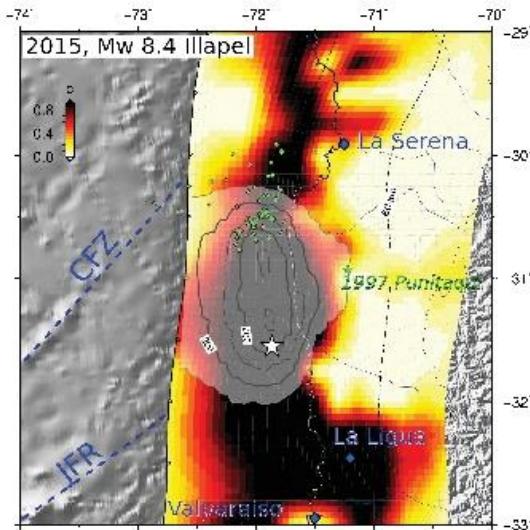
Maule 2010 8.8



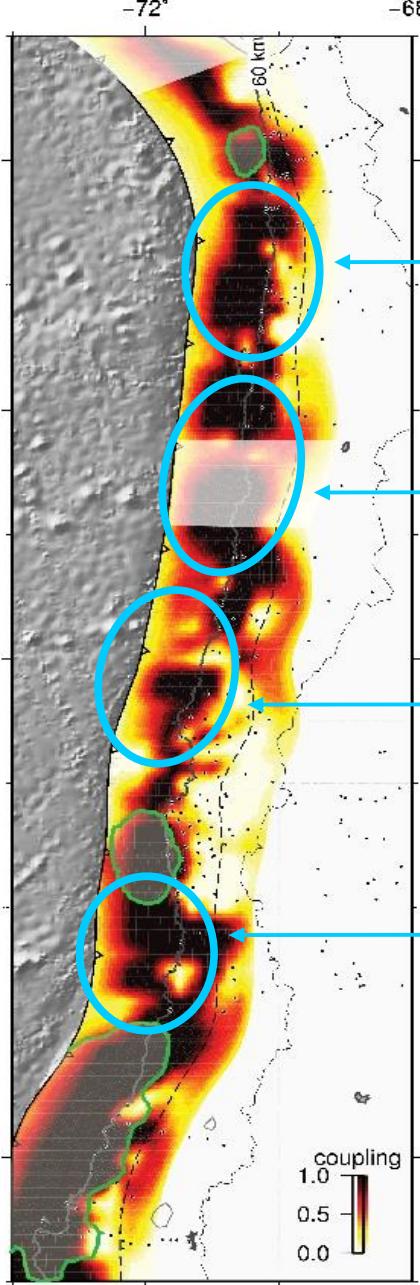
Iquique 2014 8.1



Illapel 2015 8.4



~~Plausible predictions scenarios~~



1877 north Chile Gap :

Coupled but not mature

Unlikely to break soon with $Mw \sim 9$ (like 1877)
but could produce several “smaller
magnitude(s)” $Mw 8.?$

Taltal ????

1922 Copiapo Gap :

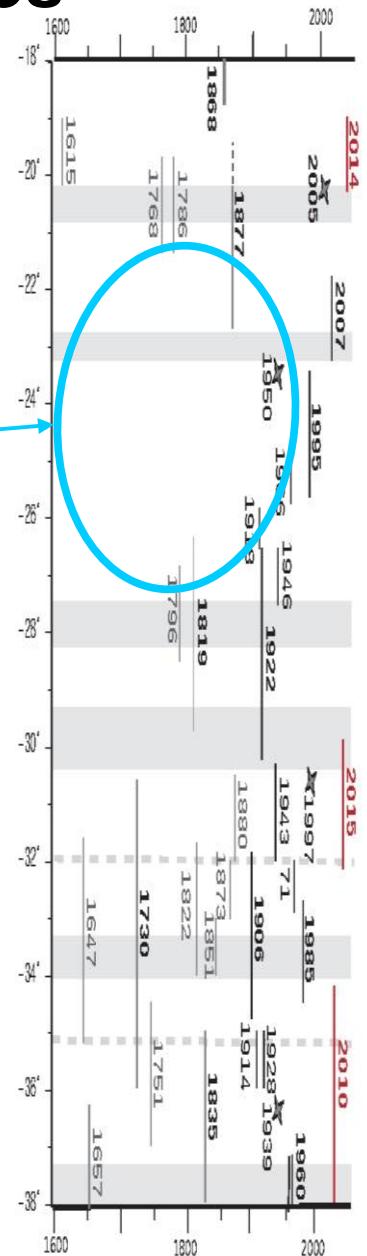
Coupled and mature and triggered

likely to break soon with $Mw \sim 8.5$ (like 1922)
But complex gap (segmented ?)

1906 Valparaiso Gap :

Coupled, may be mature but delayed by relax.

Unlikely to break soon with $Mw > 8.5$ (like 1906)

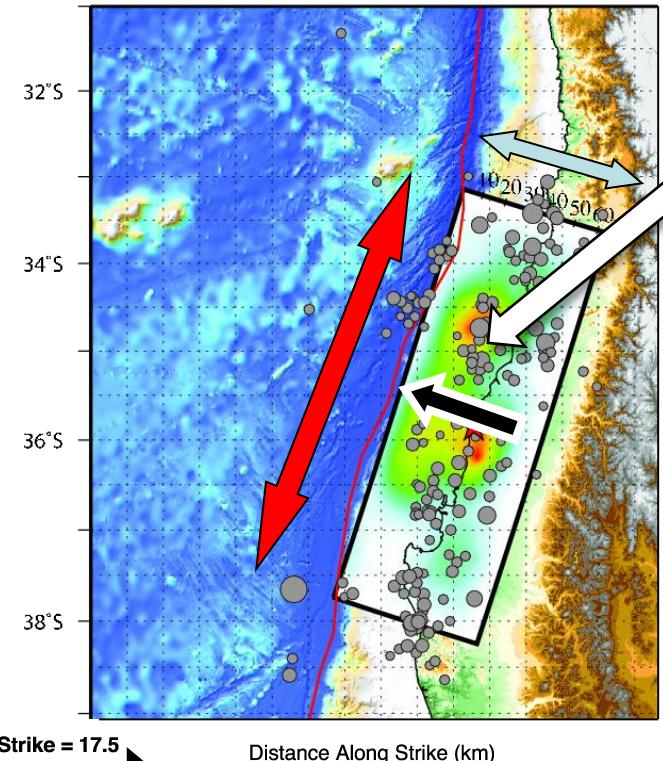


Surprises !!!

Things to learn from....

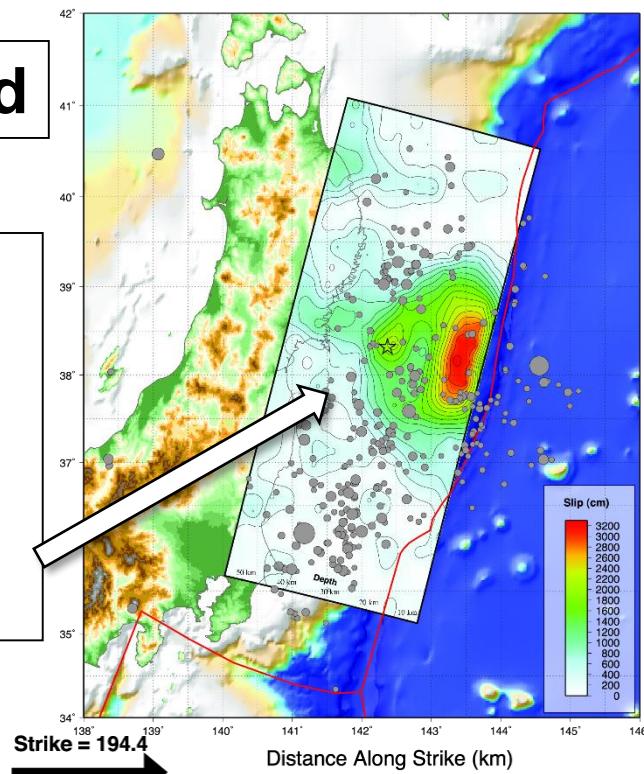
Earthquake size = surface of asperity x slip

Chile: 450 km x 150 km x 9 m => Mw 8.8 | Japan: 200 km * 150 km * 25 m => Mw 9.0



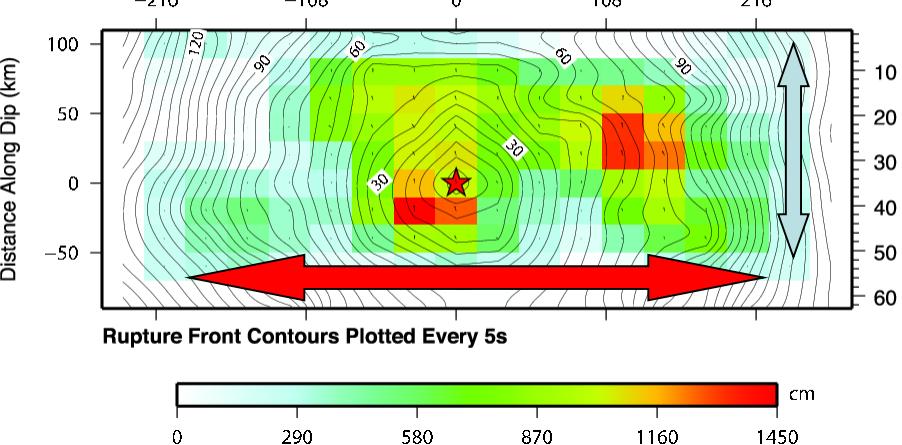
standard

Short
does
not
mean
small !

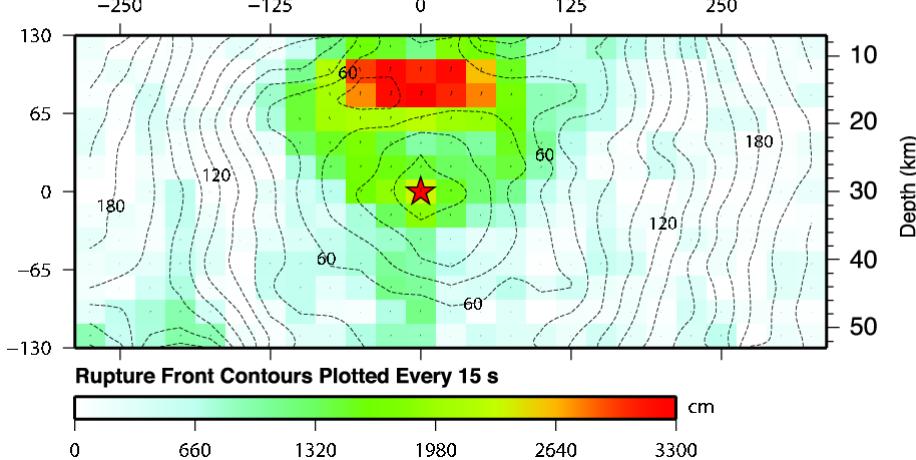


Strike = 194.4

Distance Along Strike (km)



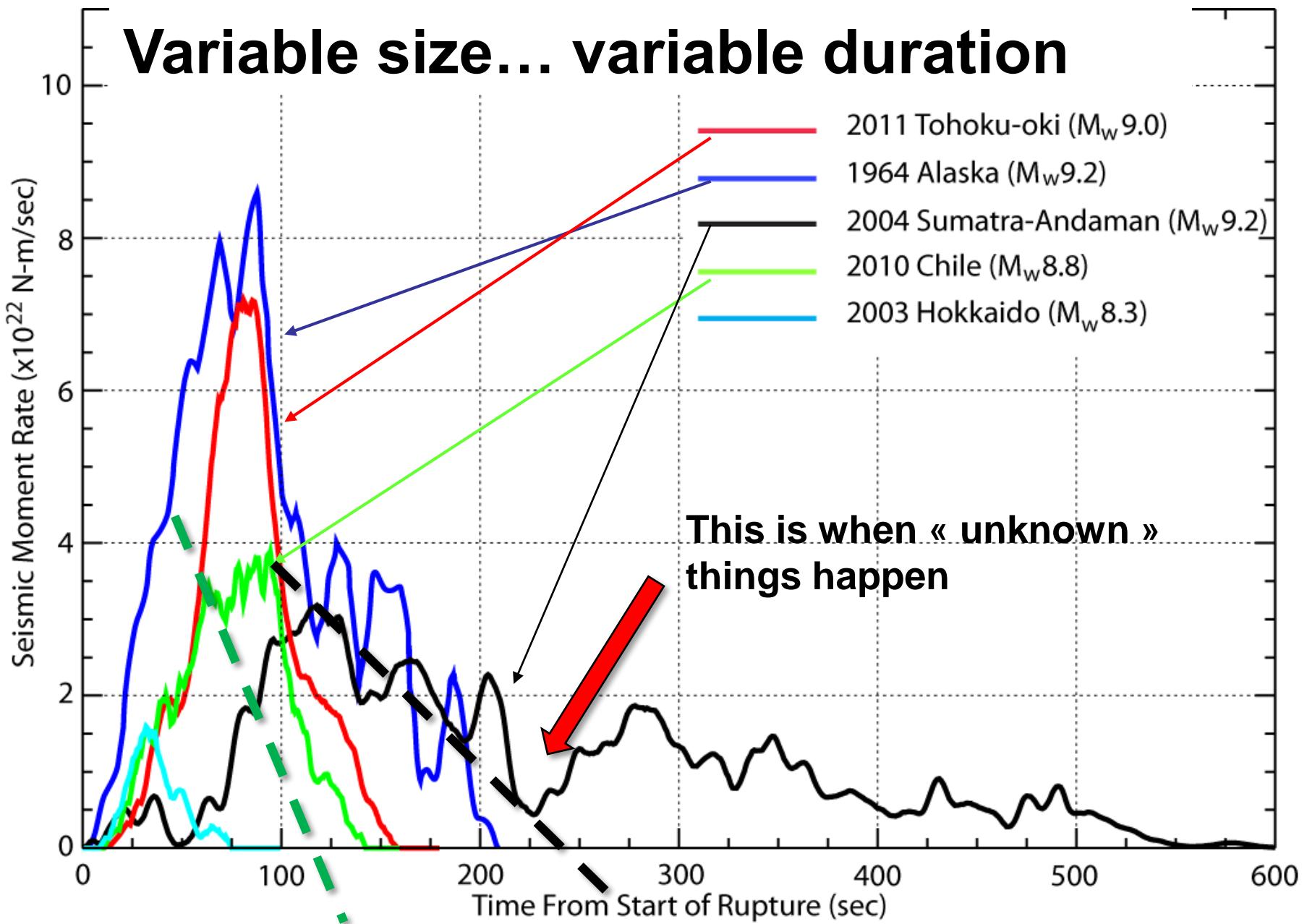
Rupture Front Contours Plotted Every 5 s

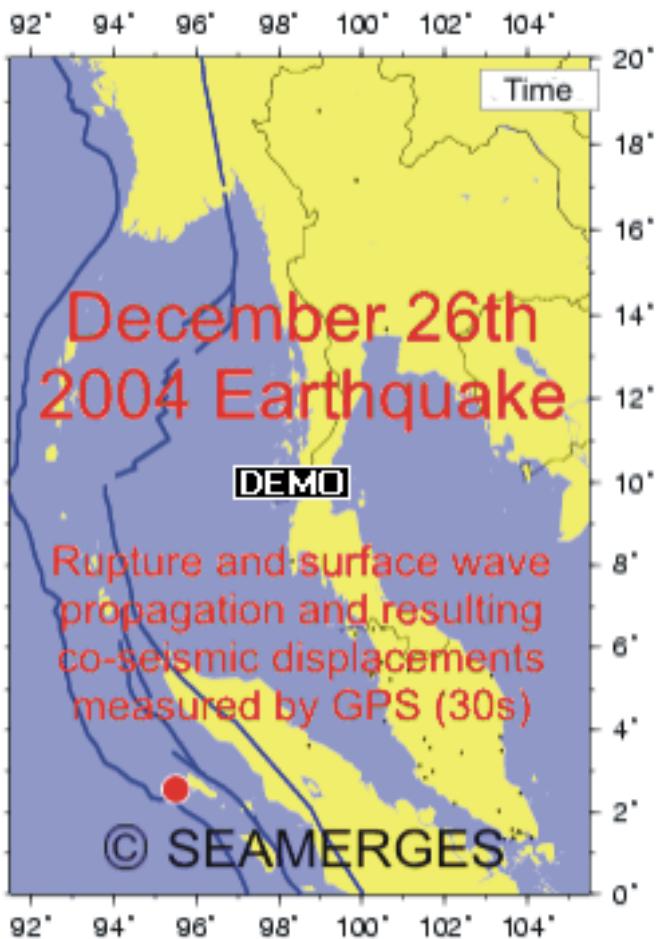


Rupture Front Contours Plotted Every 15 s



Variable size... variable duration

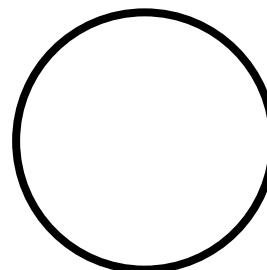




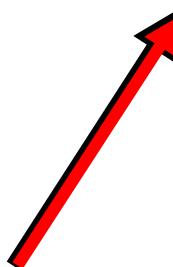
<http://www.deos.tudelft.nl/seamerges>



rupture



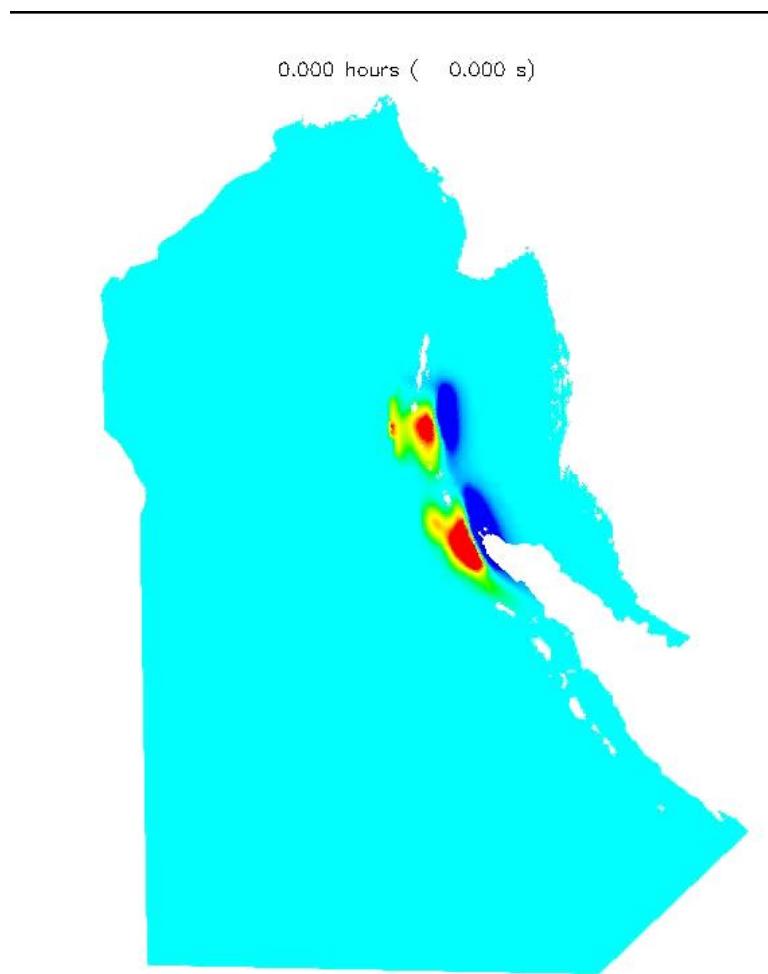
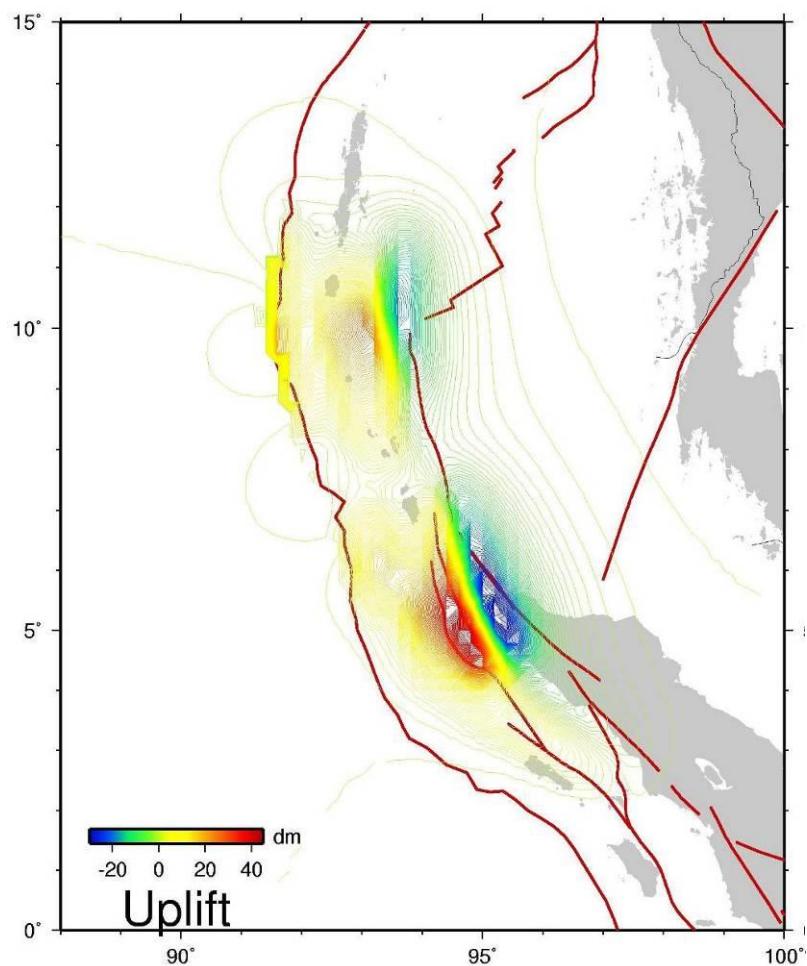
Seismic **surface**
waves propagation
(3.7 km/s)



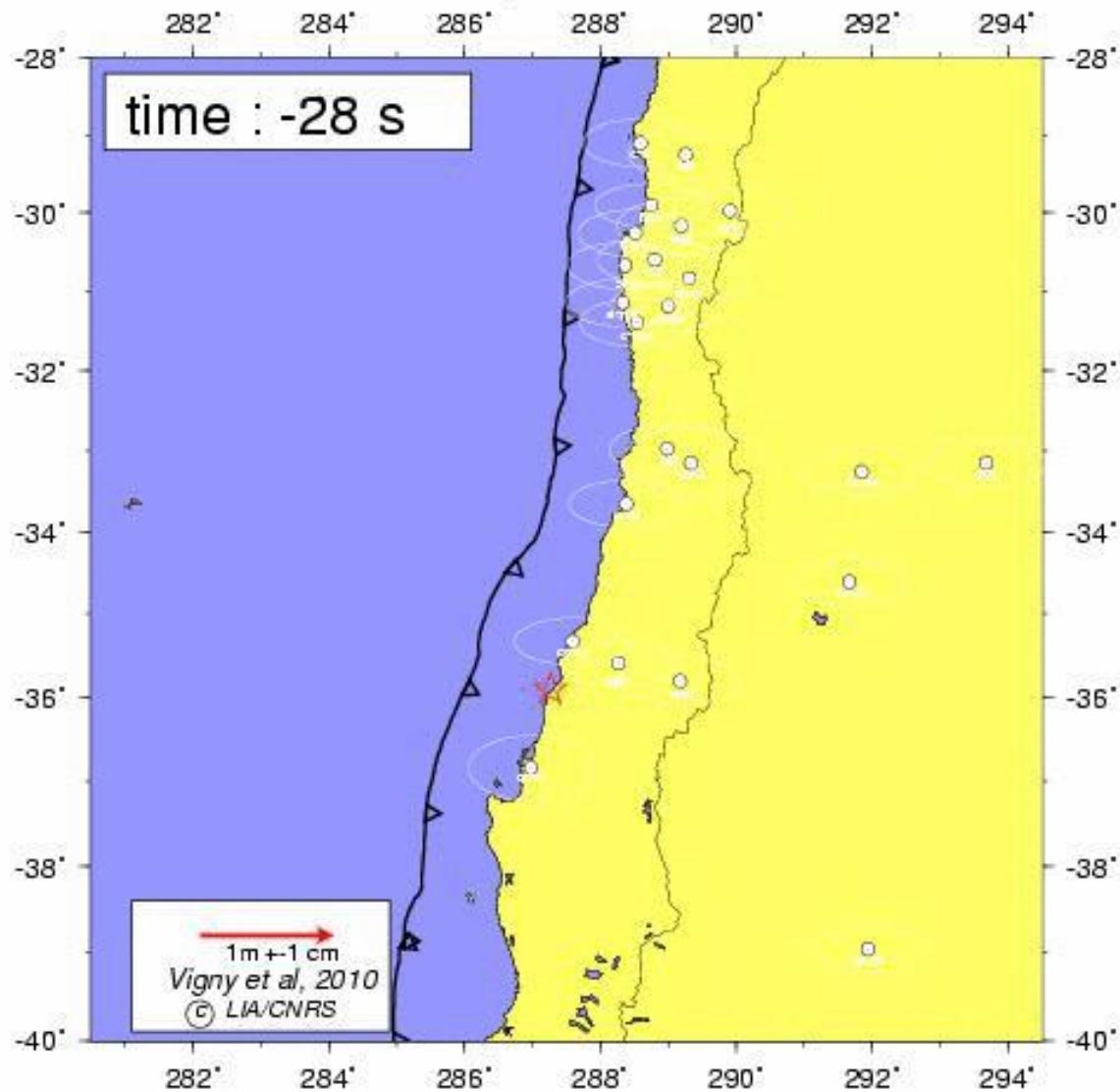
GPS stations
displacements

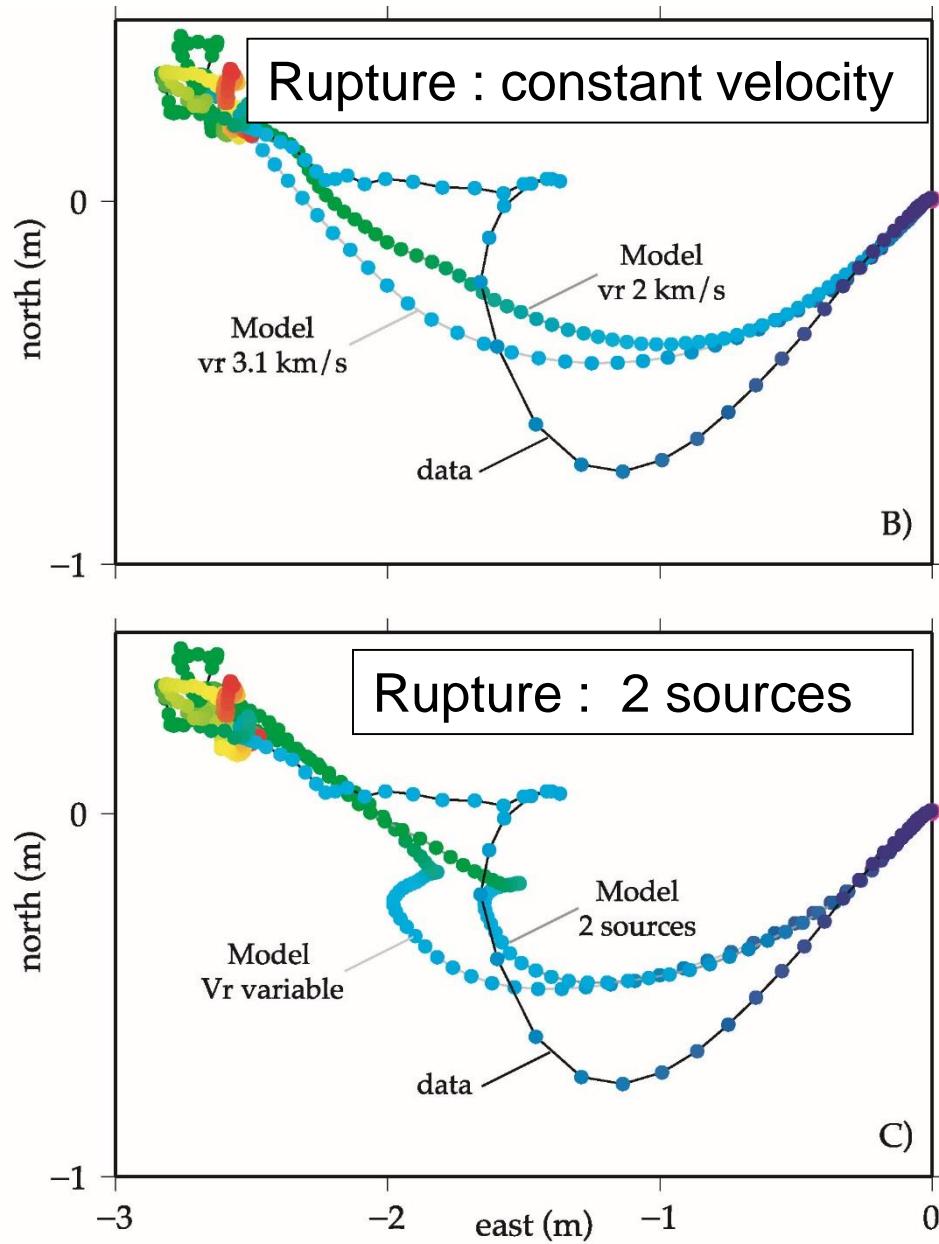
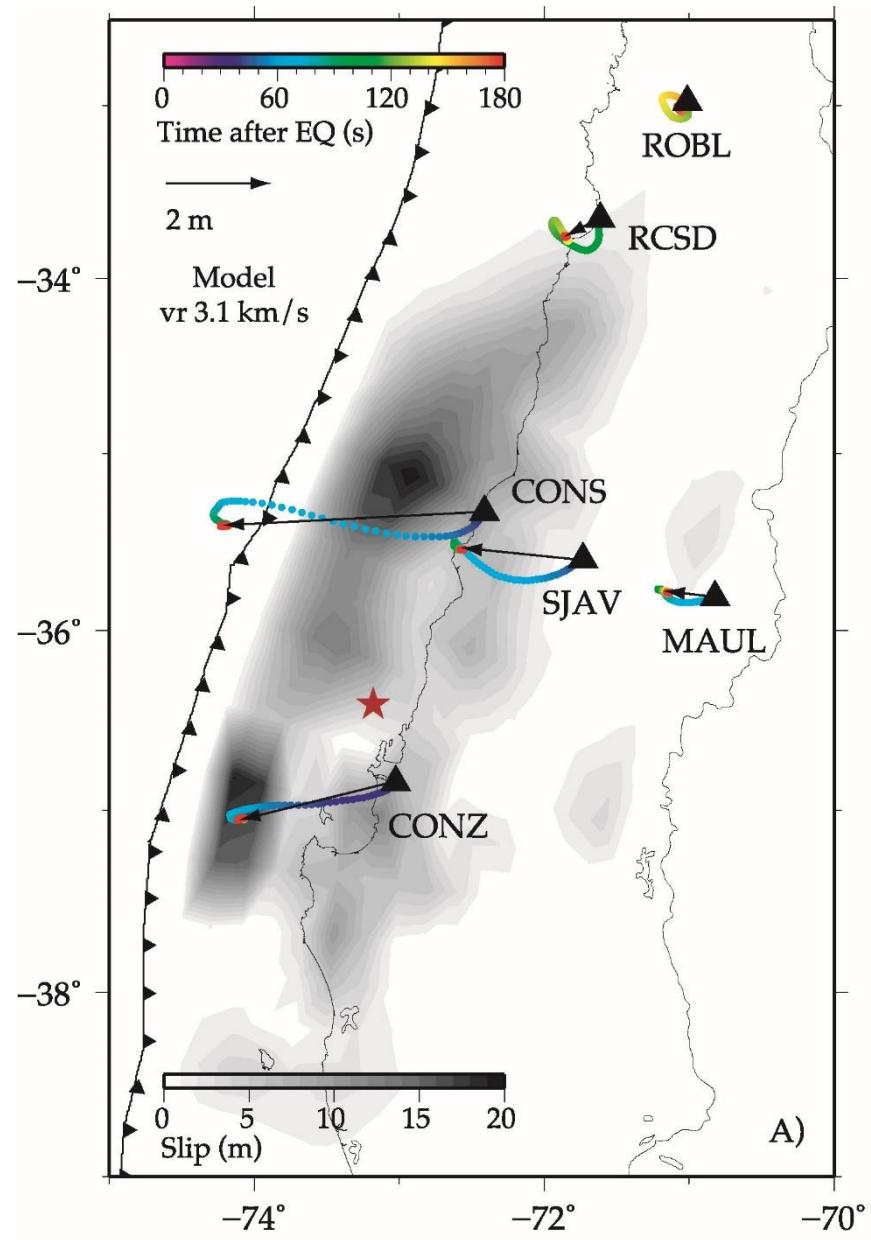
Rupture Propagation:
3.7 km/s initially (South)
30s stop $\sim 8^\circ$ lat
1.8 km/s onward (North)

Underwater ground motion => Tsunami

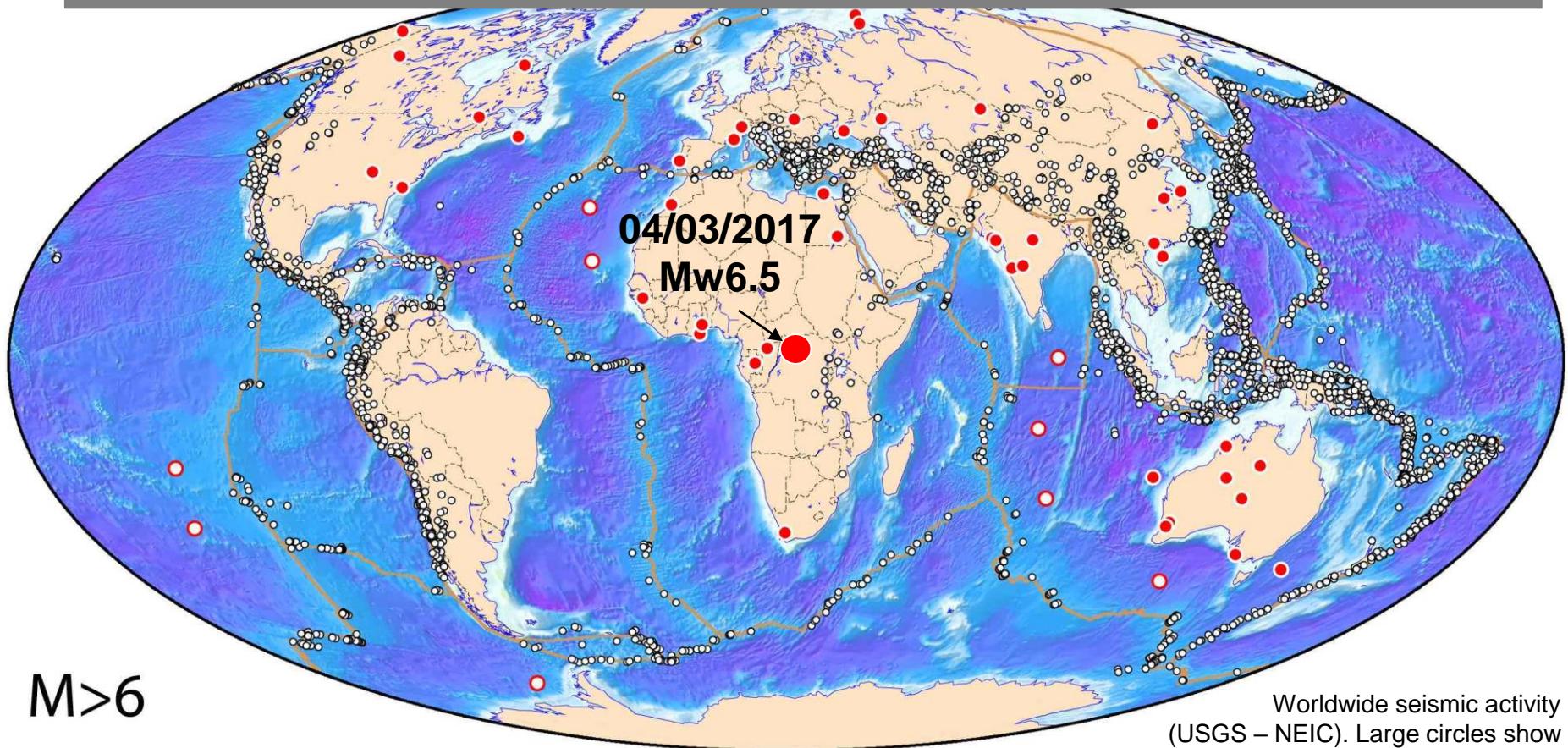


Maule Eq 27-Feb-2010



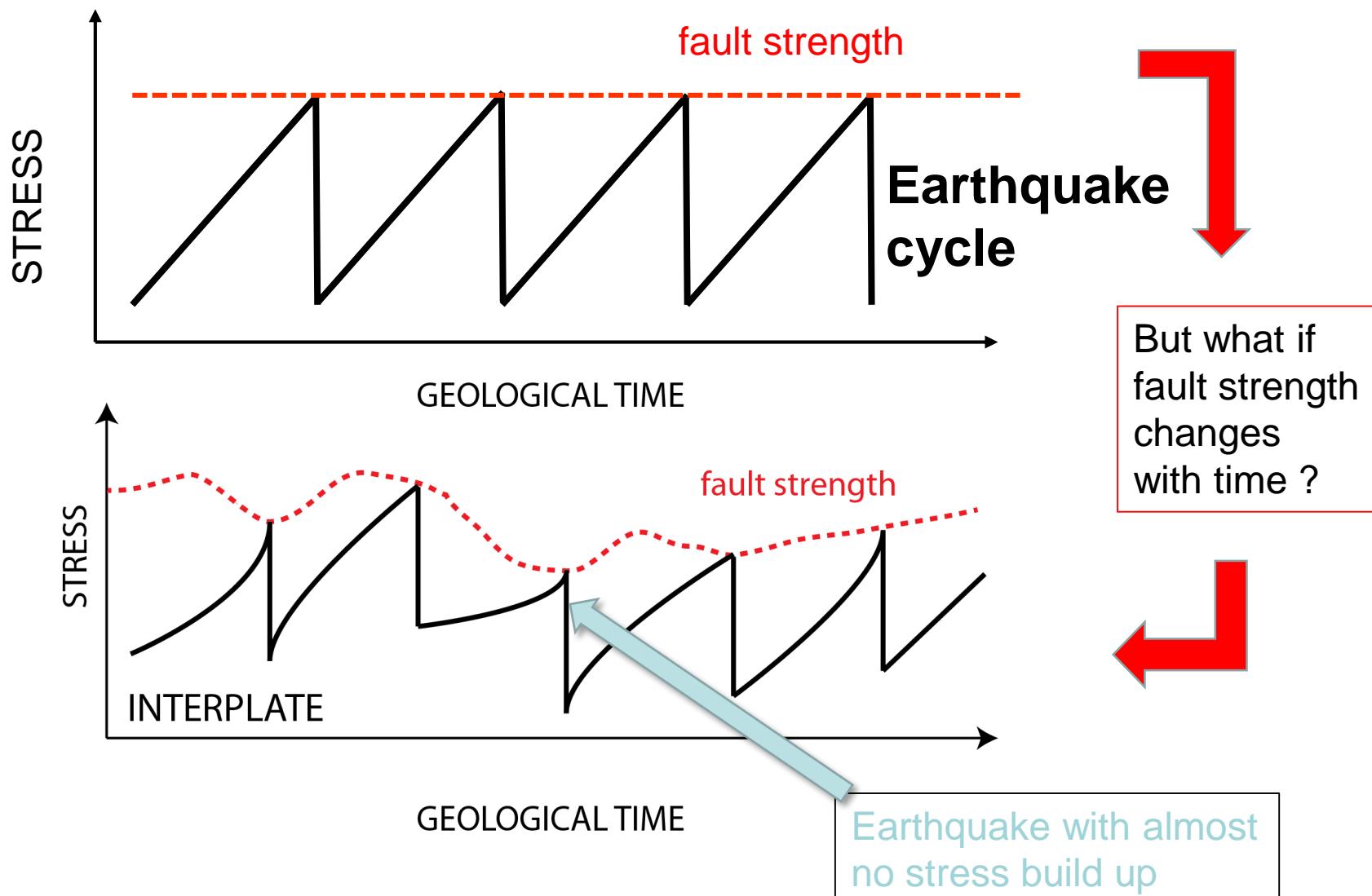


Large Earthquakes in Plate Interiors: A Challenge to Plate Tectonics

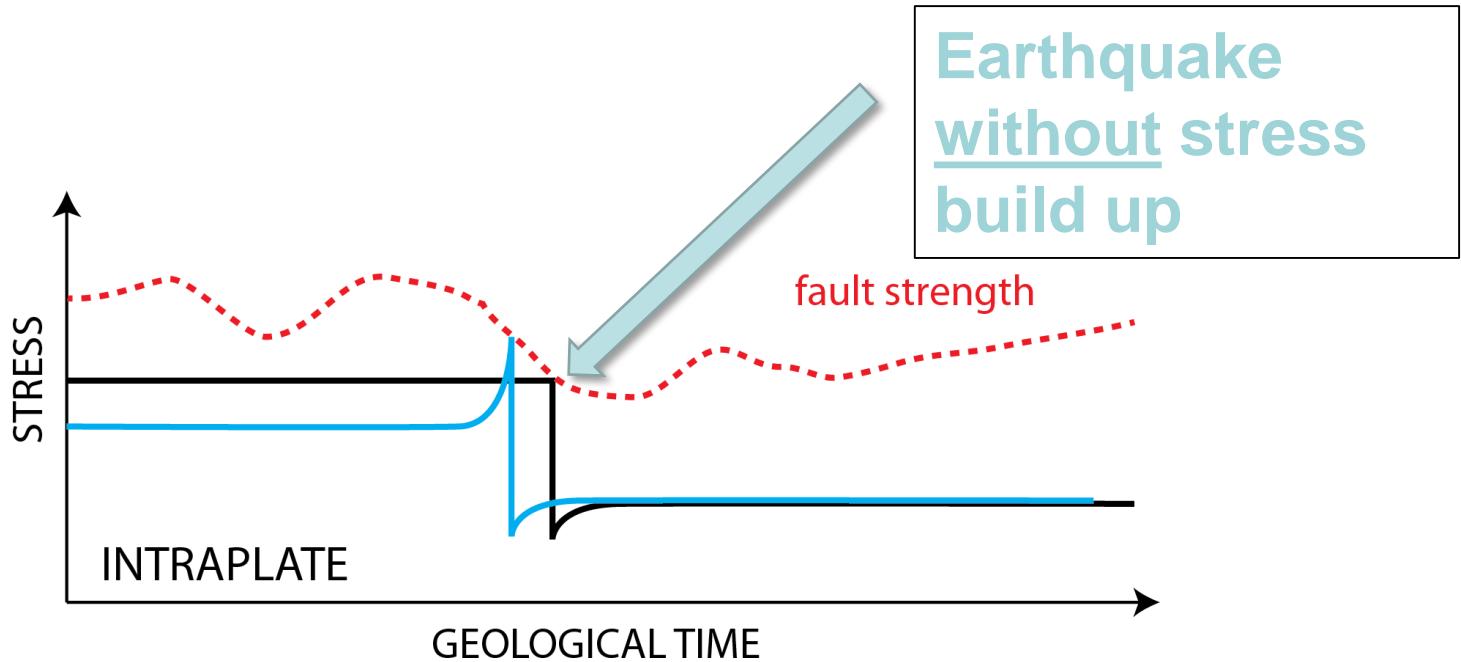


Stable continental regions (SCRs): “areas where the continental crust is largely unaffected by currently active plate boundary processes” (Johnston, 1989)

“Earthquakes occur as a result of global plate motion”
(Kanamori and Brodsky, Rep. Prog. Phys., 2004)



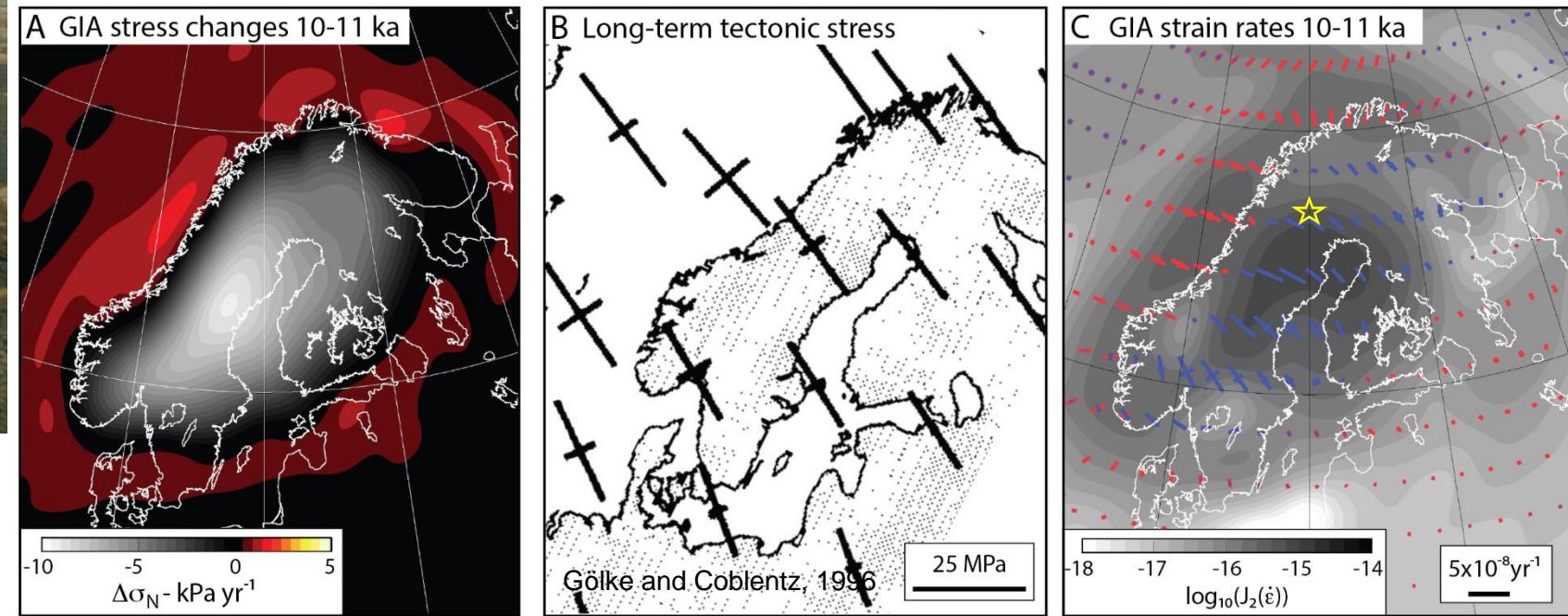
While long-term tectonic stress provides the energy that is released during large SCR earthquakes, they are **triggered** by transient perturbations of local stress that release elastic energy from **a pre-stressed crust** where faults are at failure equilibrium.



SCR earthquakes occur as a result of transient stress/strength perturbations and release elastic energy from a pre-stressed crust.

Parvie scarp, 3-10 m, reverse faulting, ~10 ka, 155 km long

Current release of paleo stress in Fennoscandia => new seismic hazard



End-glacial faults did not release strain building up at the time

=> SCR crust stores elastic energy over time scales that are longer than observable by geodesy or paleoseismology

...more [difficult] work ahead..

Thanks for attention



21 4 2007

Size of an Earthquake

Earthquake « size » or released Moment M_0 , is proportional to :

- Quantity of slip (U)

 fault velocity (V) \times time between earthquakes Δt

- Size of ruptured surface (S)

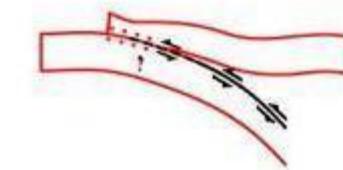
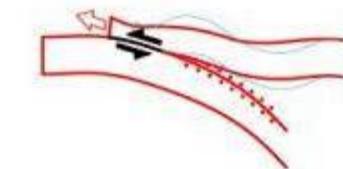
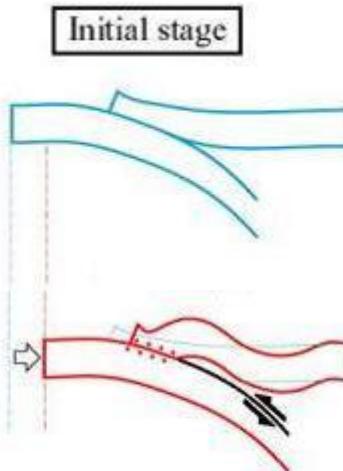
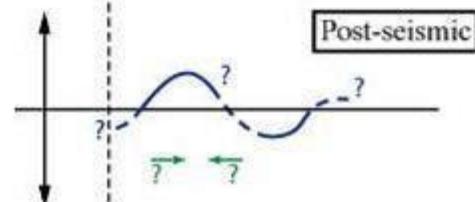
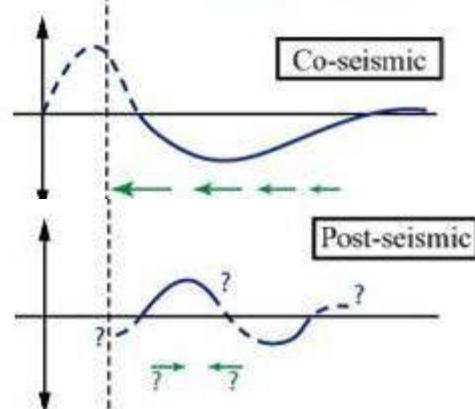
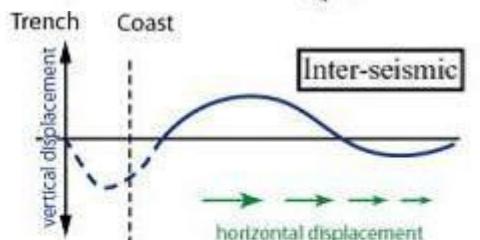
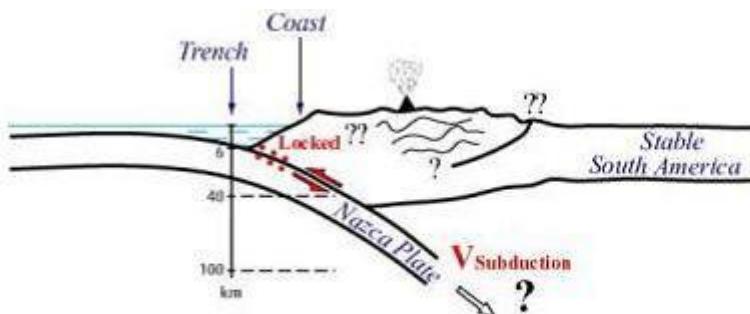
 Length of rupture (L) \times Locking depth of fault (d)

$$\Rightarrow M_0 = \mu \times S \times U = \mu \times L \times d \times V \times \Delta t$$

Magnitude of an Earthquake (Hanks & Kanamori):

$$M_w \sim \log(M_0) = 2/3 \log(M_0) - 10.7$$

La Subduction : Un très très gros ressort !!!



Centaines
d'années

Quelques secondes
quelques minutes

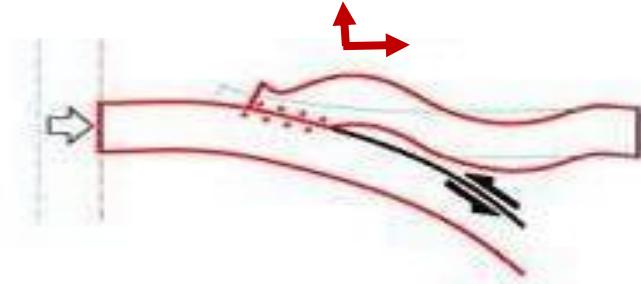
Des mois
des années

Seismic cycle on subduction zone

Interseismic stage :

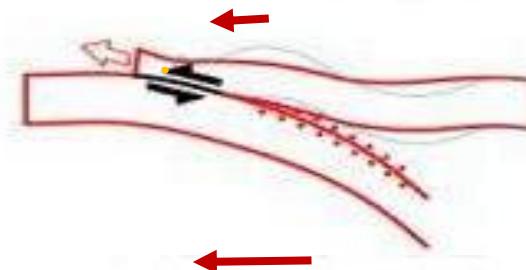
Interface locked over several decades

Strain accumulation *over decades to century*



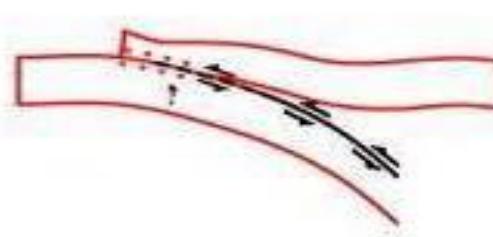
Coseismic stage : The earthquake

Sudden stresses release *in a few seconds/minutes*



Postseismic stage :

Response of coseismic stresses transfert on deep layers *over decades*



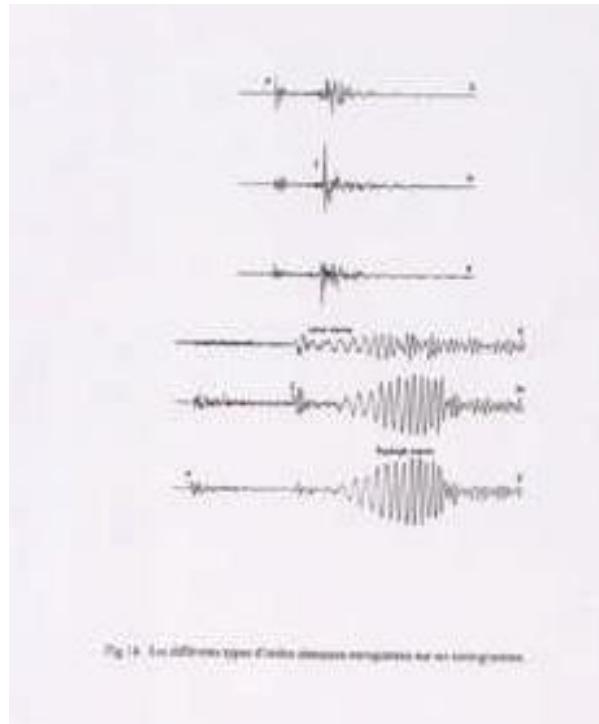
Different mechanisms:

Modified from Chlieh *et al.*, 2004

- Poroelastic rebound → < 50km from the trench / several months
- Elastic afterslip → first hundreds km / few months to few years
- Viscoelastic relaxation → very large scale / few decades

Pourquoi un séisme produit-il des dommages ?

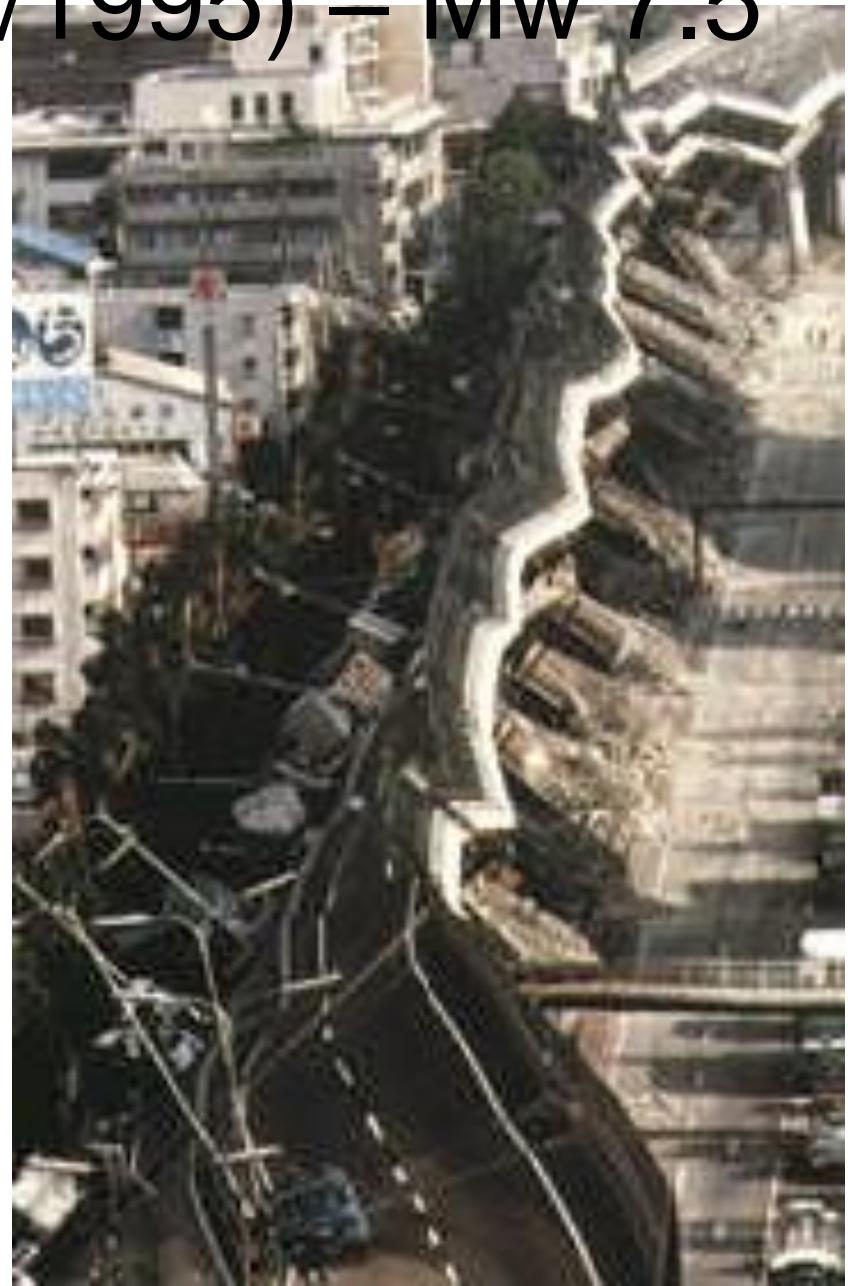
À cause de l'énergie rayonnée sous forme d'ondes sismiques



L'énergie émise est bien sûr reliée à la taille de la rupture (au moment sismique), *mais pas seulement*. Elle dépend aussi de la vitesse à laquelle la « déchirure » se propage.....

Kobé, Japon (17/01/1995) – Mw 7.5

- dommages au bâti sont surtout causés par mouvements horizontaux du sol (ondes S) dans la gamme de période de 1 seconde à 0,1 seconde qui correspond aux résonnances de la plupart des bâtiments



- La chute opposée de deux rangées de vélos permet de déterminer la demi longueur d'onde du mouvement horizontal (onde S)



- Effands indirects : liquéfaction du sol
- Caracas, Venezuela 29/07/1967 Mw 6.6

Photomontage de la phénomène de liquéfaction responsable du basculement et de l'enfoncement d'un immeuble à Caracas (Venezuela), lors du séisme du 29 juillet 1967 (magnitude 6,6).
(Collection NTSB)



Risque sismique = Aléa x Vulnérabilité

- **Aléa sismique:**

Probabilité pour un site de subir un séisme donné (localisation, profondeur, taille, type, andc...) à un instant donné

- **Vulnérabilité:**

Dépend 1/ des personnes and 2/ des biens exposés à l'aléa:

1/ De la Population :

- De son niveau d'éducation
- De son niveau économique and social
- De l'existence ou non d'une "culture sismique" des individus and des pouvoirs publics (réflexes, organisation, secours, andc...)

2/ Des constructions:

- Du type and de la qualité des matériaux
- Du mode de construction
- De la qualité du sol

- Dans le désert, le risque sismique est négligeable même si l'aléa est élevé.
Exemple: le séisme des îles Balleny, Antarctique, 2004.
- Dans régions à forte densité de population (and aux constructions précaires), le risque sismique est très élevé même si l'aléa est modéré.
Exemple séismes de Bâle (Suisse) au 16eme siècle ou de Lisbonne au 17eme siècle.
- La vulnérabilité, **donc le risque** est en augmentation constante dans pays en voie de développement (surpopulation and paupérisation).

Séisme d'El Asnam (Algérie), le 10 octobre 1980

M = 7,3



Risque sismique négligeable



Risque sismique élevé

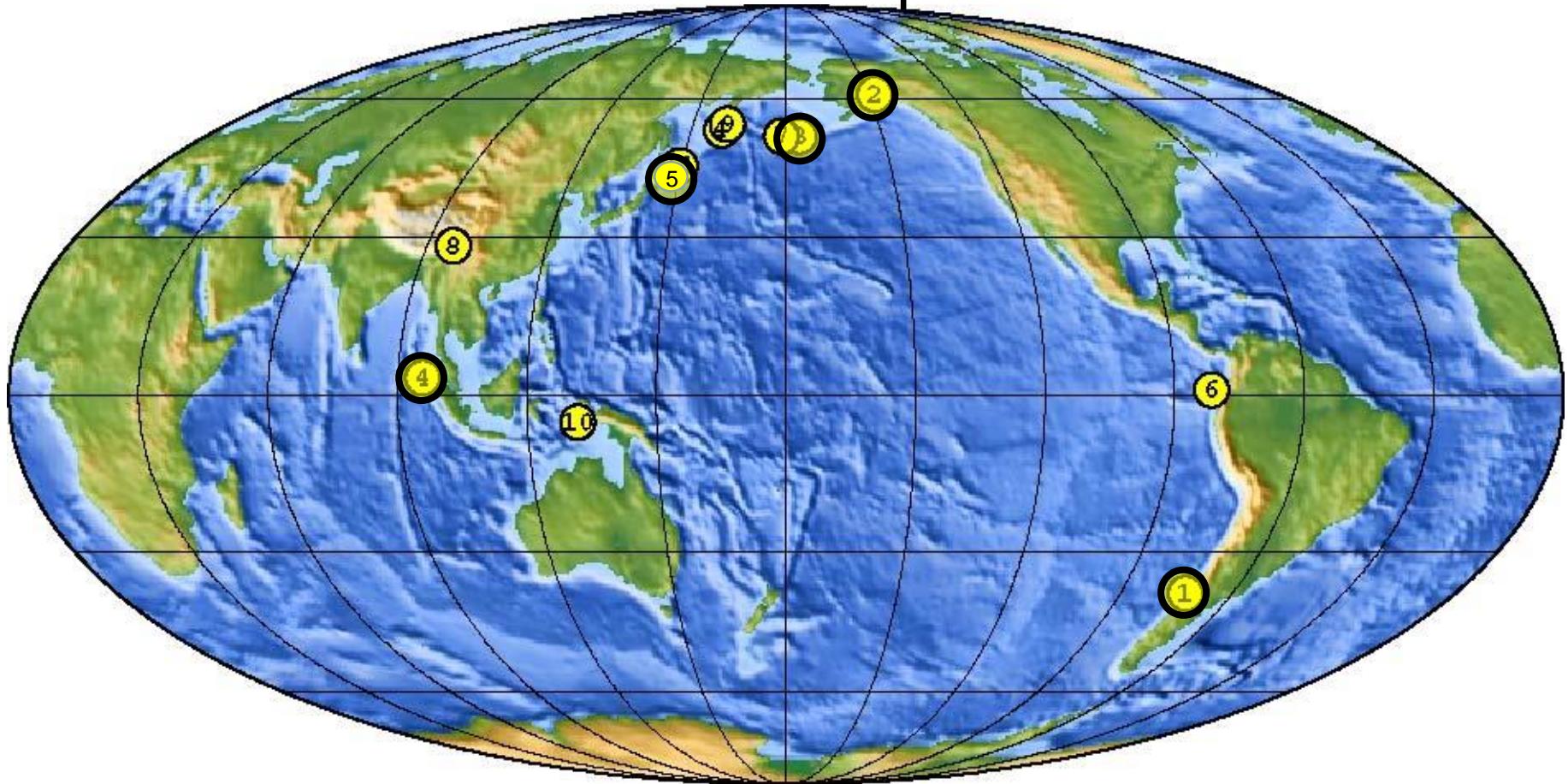
Exemp de vulnérabilité du bâti: étage souple, piliers sous dimensionnés, maçonnerie non chaînée....

Zémmouri-Boumerdes, Algérie Mai 2003 Mw 6.5



plus gros séismes en 50 ans

5: c'est bien environ 1 par décennie



1 : Valdivia 1960: 9,6

3 : I Aléoutiennes 1957: 9,2

2 : Alaska 1964: 9,4

4 : Sumatra 2004: 9,2

5 : Japon 2011: 9,0

Chili : Déjà 4 séismes majeurs en quelques années

20^{eme} siècle (10)

Antofagasta	1995	8.0	
Valparaiso	1985	8.0	
Valparaiso	1971	7.8	
Valdivia	1960	9.6	
Antofagasta	1950	8.2	
Coquimbo	1943	8.1	
Chillan(*)	1939	7.8	
Copiapo	1922	8.3	
Atacama	1918	7.8	
Valparaiso	1906	8.2	

21^{eme} siècle (4) and counting

Tocopilla	2007	7.8
Maule	2010	8.8
Iquique	2014	8.2
Illapel	2015	8.4

