

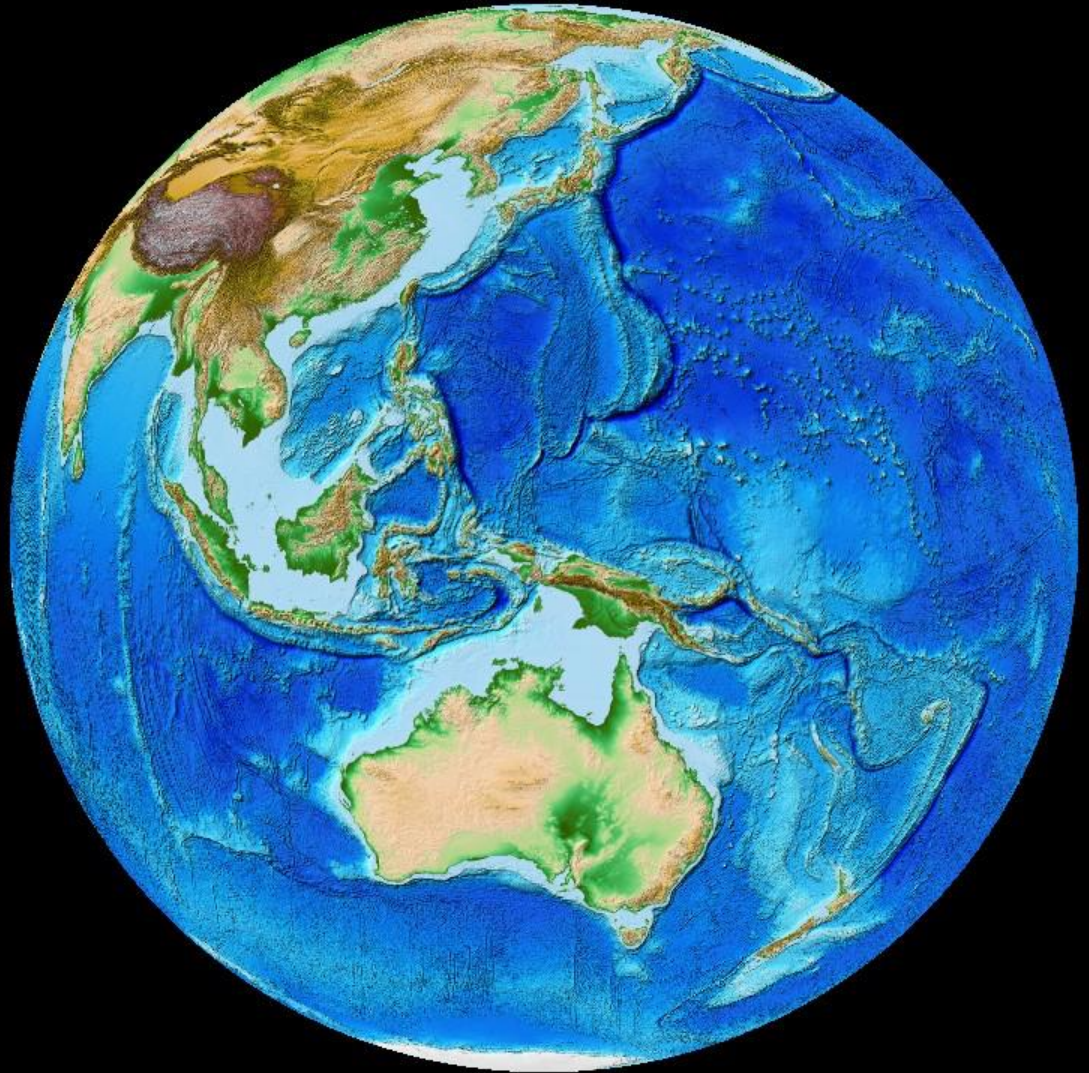
# Plate Tectonics: Linking Surface Geology to Earth's Deep Interior

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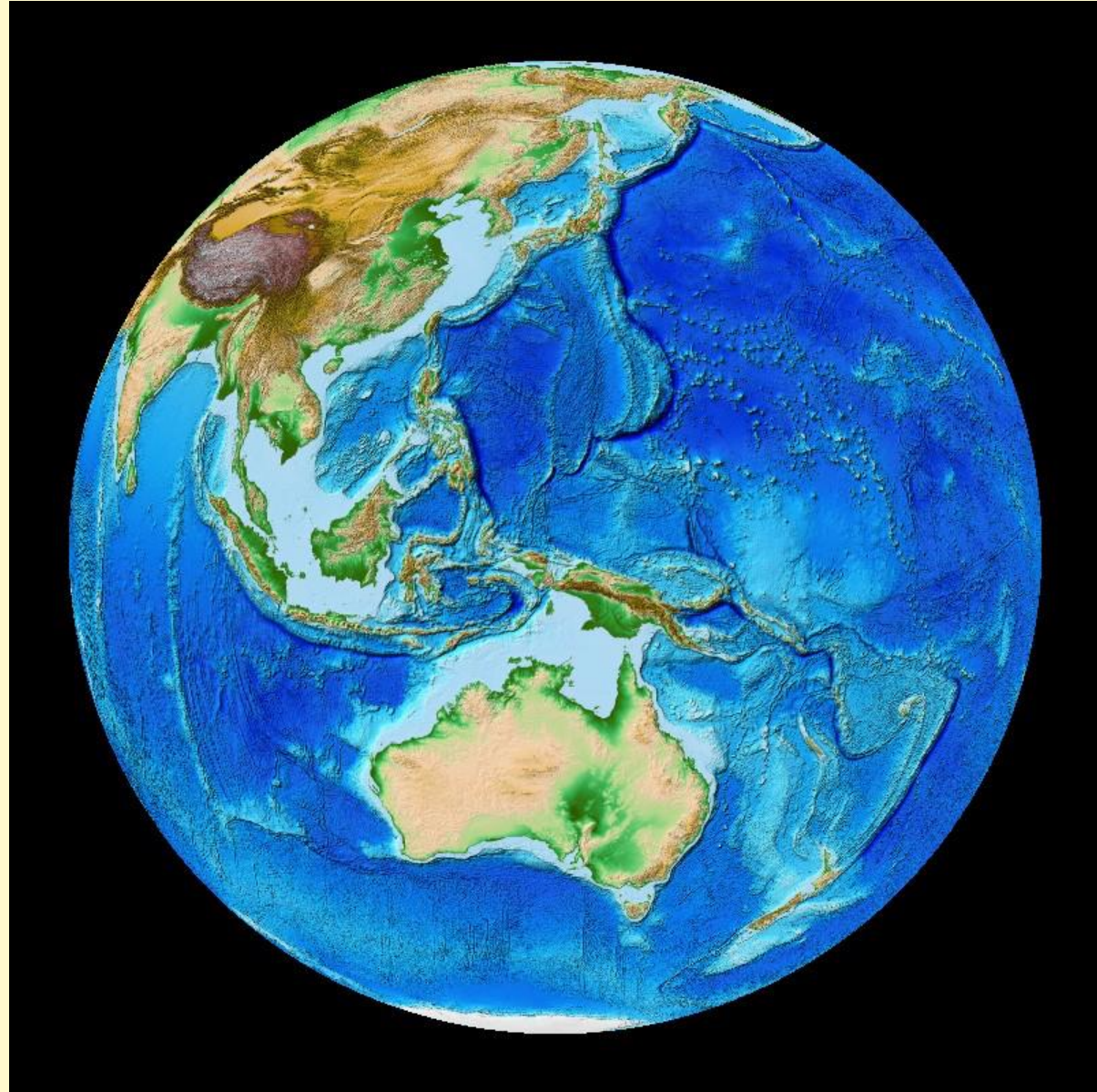


## From observations at the surface:

→ **Geologic change:**  
expresses a long  
time history

→ **Tectonic change:**  
great forces arise  
from the deep  
interior

**MAIN QUESTION:**  
How is surface  
geology linked to  
the deep interior?



Surface

← Depth

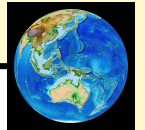
CMB

Earliest Earth:  
4.5 Billion years ago

Time →

Present

Now



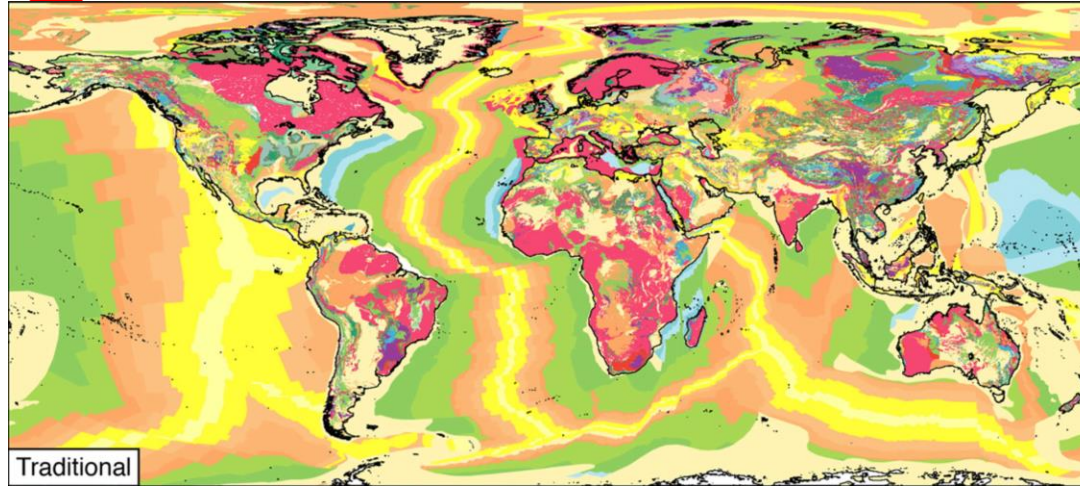
**MAIN QUESTION:**  
How is surface geology linked  
to the deep interior?

# Observations of Geological Change



Surface

← Depth



Geological Time	[m.y.]	Traditional
Quaternary	0 - 2.6	Lightest yellow
Pliocene	2.6 - 5.3	Yellow
Miocene	5.3 - 23	Light orange
Oligocene	23 - 34	Orange
Eocene	34 - 56	Dark orange
Paleocene	56 - 66	Red-orange
Late Cretaceous	66 - 100	Red
Early Cretaceous	100 - 145	Dark red
Late Jurassic	145 - 164	Light blue
Middle Jurassic	164 - 174	Blue
Early Jurassic	174 - 201	Dark blue
Late Triassic	201 - 237	Purple
Middle Triassic	237 - 247	Dark purple
Early Triassic	247 - 252	Dark red
Late Permian	252 - 260	Red
Middle Permian	260 - 272	Dark red
Early Permian	272 - 299	Red
Late Carboniferous	299 - 323	Dark red
Early Carboniferous	323 - 359	Dark red
Late Devonian	359 - 383	Dark red
Middle Devonian	383 - 393	Dark red
Early Devonian	393 - 419	Dark red
Silurian	419 - 444	Dark red
Late Ordovician	444 - 458	Dark red
Middle Ordovician	458 - 470	Dark red
Early Ordovician	470 - 485	Dark red
Cambrian	485 - 541	Dark red
Precambrian	541 - 4500	Dark red

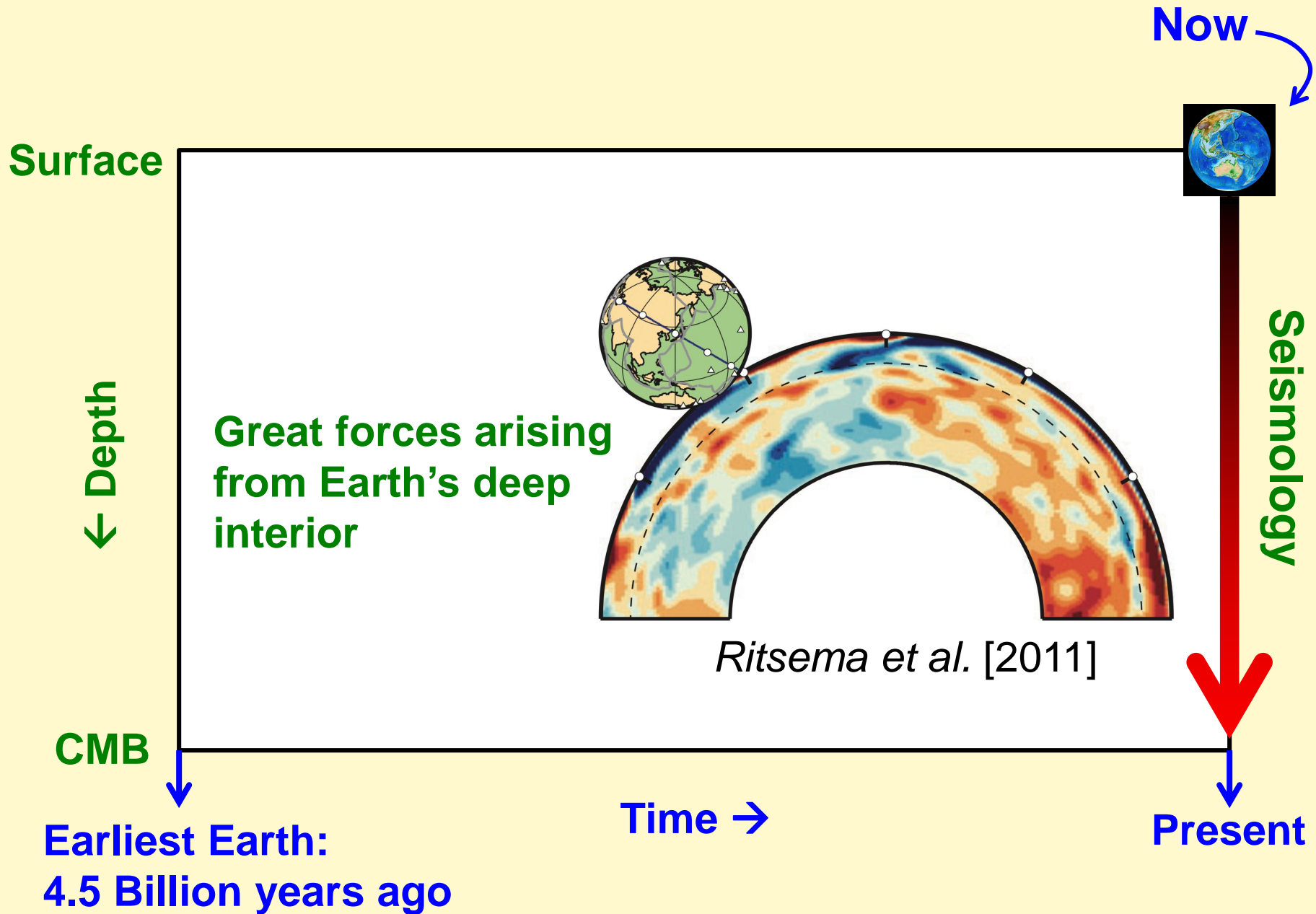
*Friedrich et al. [2018]*

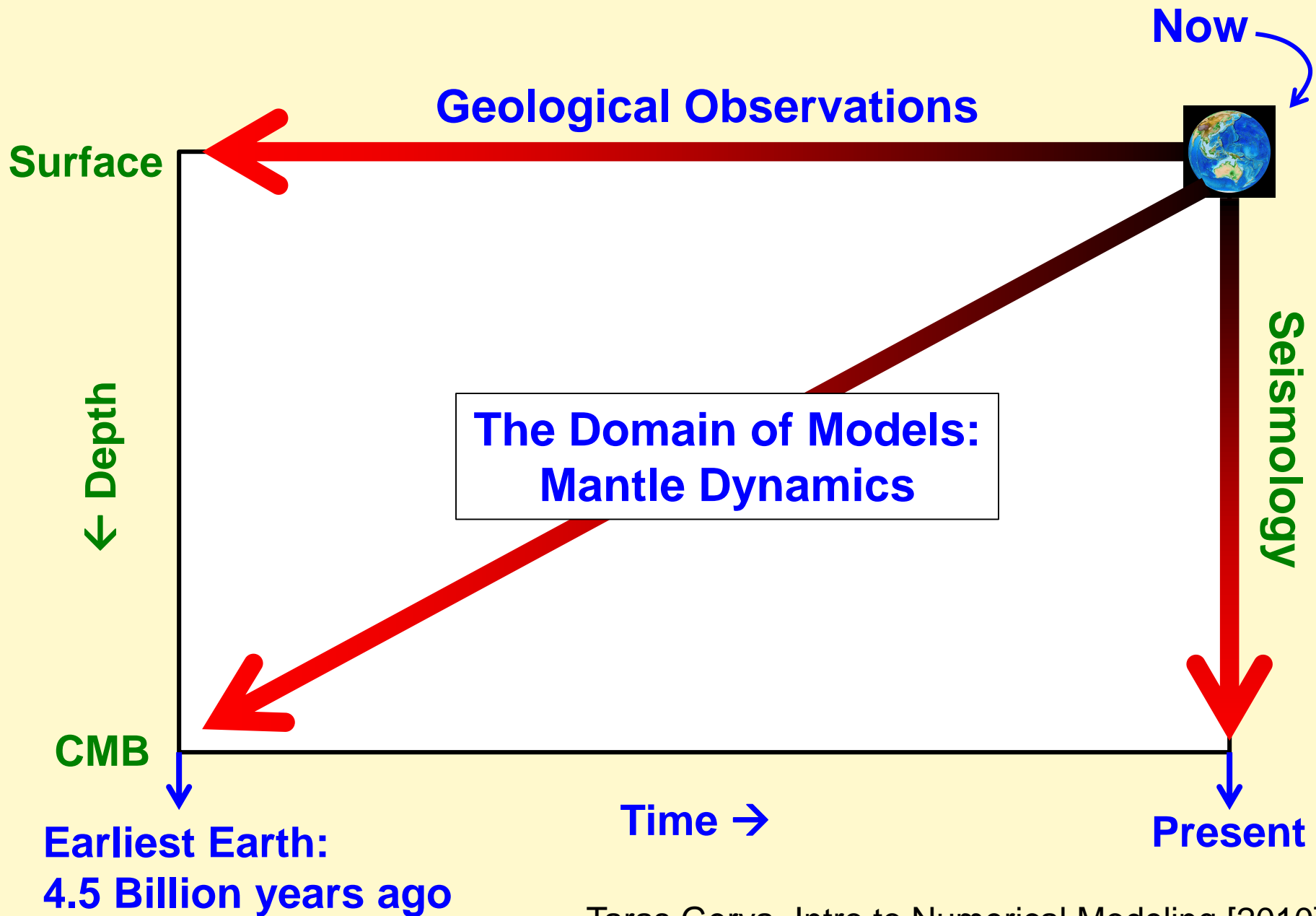
CMB

Earliest Earth:  
4.5 Billion years ago

Time →

Present

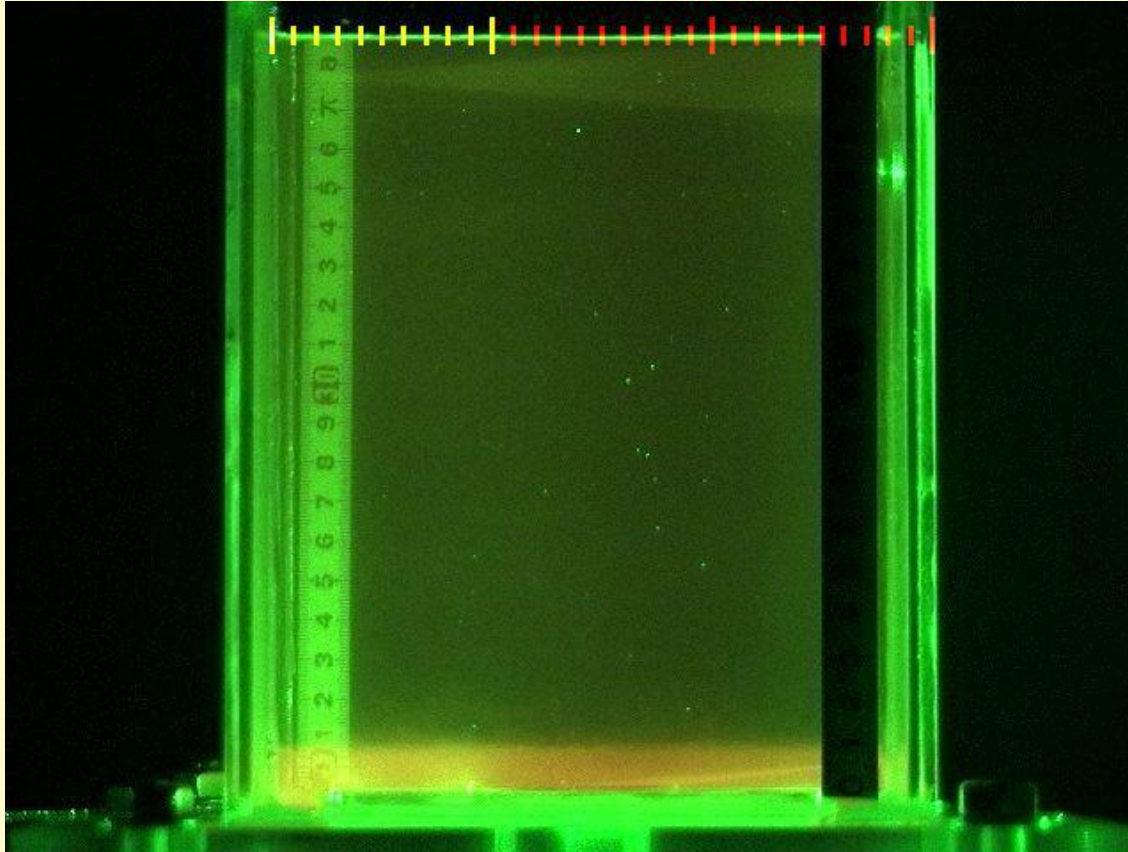




# Models of Mantle Dynamics

- A model is a *representation* of some a physical process
- A model should be *useful* for understanding the process
- A model should be *testable* in some way

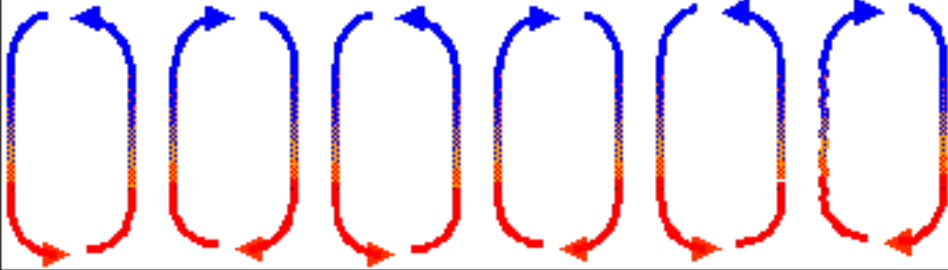
“All models are wrong but some are useful.” [George Box, 1987]



A Plume  
Experiment in  
Corn Syrup

# How Convection Works

Fluid cools by losing heat from surface



## Convection cell

Warm, low density fluid rises  
Cool, high density fluid sinks

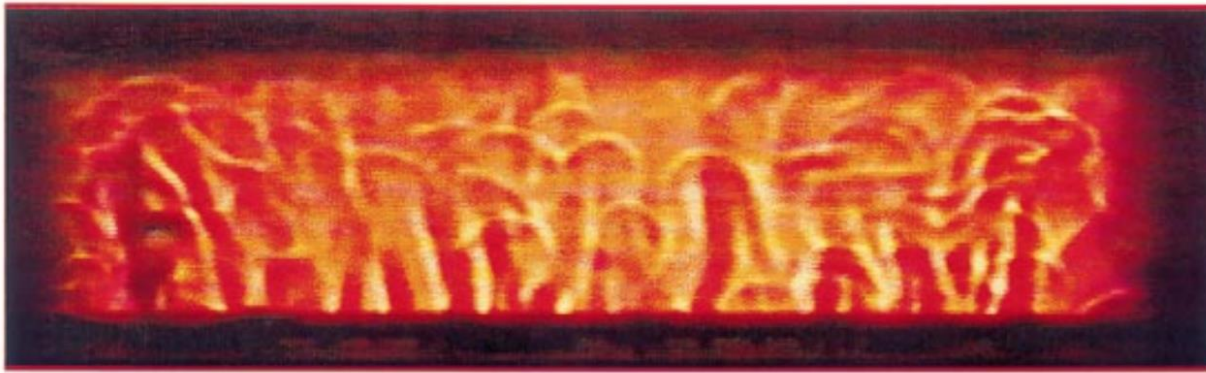
Cold Fluid  
Sinks

Hot Fluid  
Rises



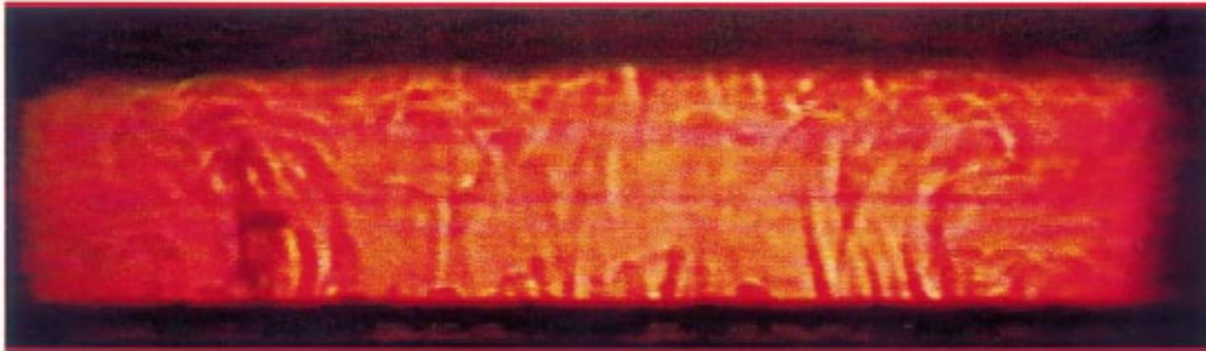


(a) 20880 s



← Base is Hot →

(b) 56220 s



(c) 88200 s



Laboratory experiment of convection in a tank of corn syrup.

*Lithgow-Bertelloni et al. [2001]*

**Tank is getting hotter with time**

The **Rayleigh Number** is a dimensionless parameter that measures the **vigor of convection**:

$$Ra = \frac{r g a D T^3}{k h}$$

$r$  = density (3300 kg/m<sup>3</sup>)

$g$  = gravity (10 m/s<sup>2</sup>)

$a$  = thermal expansivity ( $3 \times 10^{-5}$  K<sup>-1</sup>)

$DT$  = Temperature contrast across mantle (3000 K)

$D$  = Depth of Mantle (2860 km)

$k$  = Thermal diffusivity ( $10^{-6}$  m<sup>2</sup>/s)

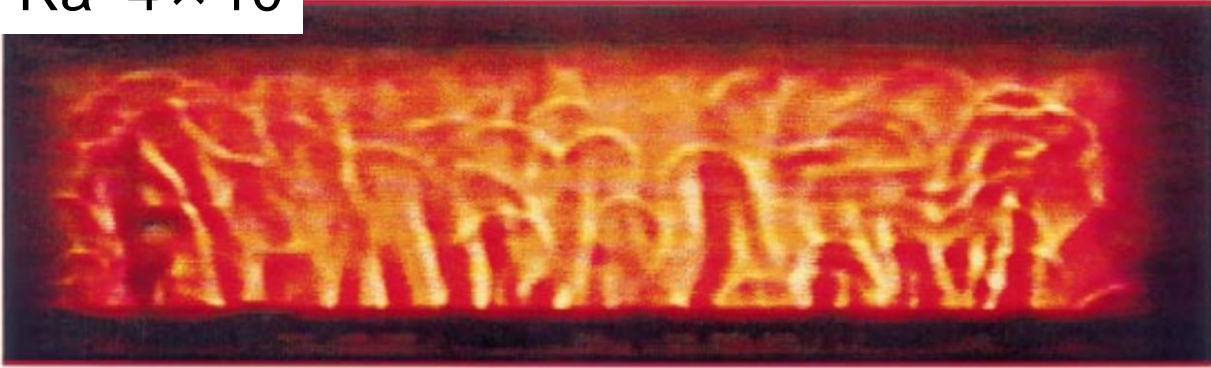
$h$  = Mantle viscosity ( $10^{21}$  Pa s)

Convection occurs if  $Ra > 657$

Using these parameters for the mantle:  $Ra_m \sim 7 \times 10^7$

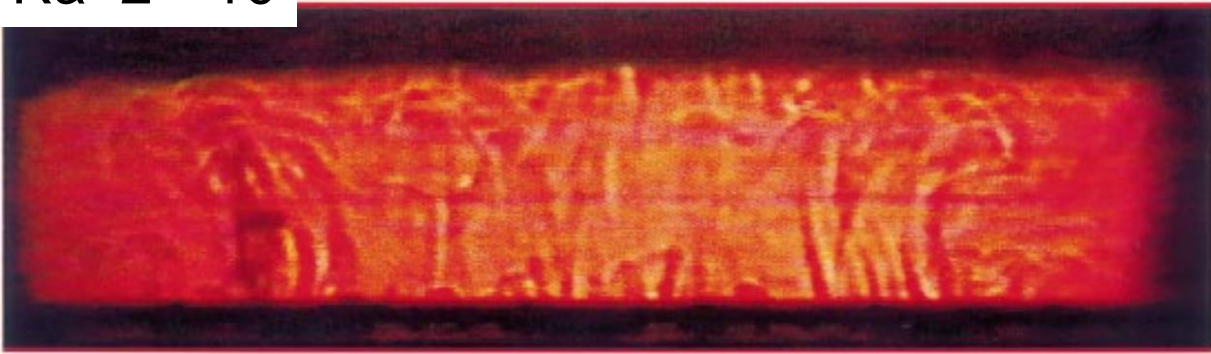
→ This “model” implies vigorous convection in the mantle

$Ra \sim 4 \times 10^6$

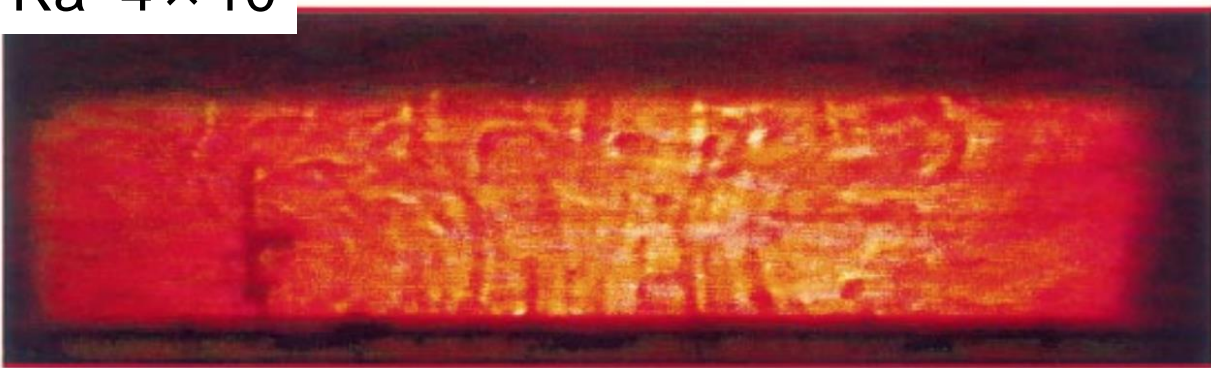


← **Base is Hot** →

$Ra \sim 2 \times 10^7$



$Ra \sim 4 \times 10^7$



Laboratory experiment of convection in a tank of corn syrup.

*Lithgow-Bertelloni et al. [2001]*

**Rayleigh number is getting larger with time**

**→ Vigorous Convection!**



## How to solve for mantle dynamics using a computer:

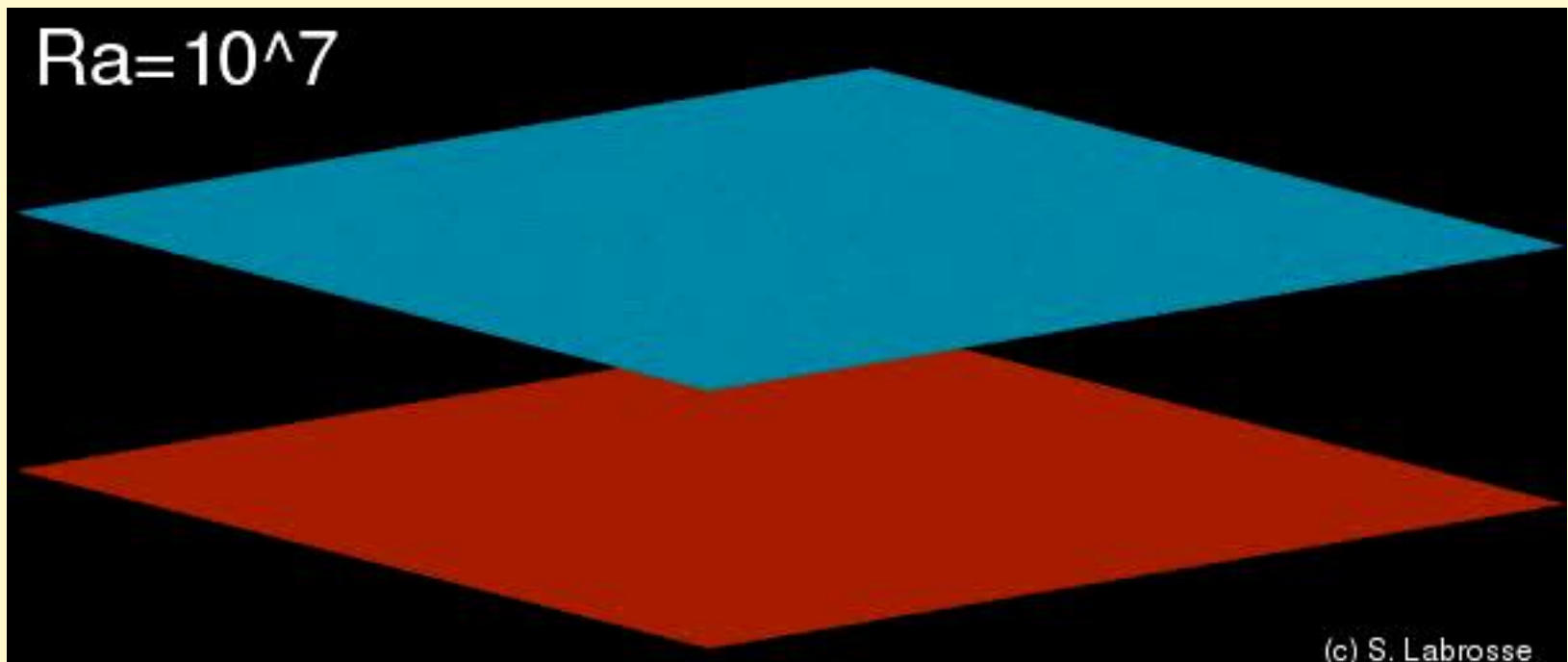
### 1. Express the physics of convection using equations

Conserve Mass:  $\nabla \cdot \mathbf{v} = 0$

Conserve Momentum:  $-\nabla p + \eta \nabla^2 \mathbf{v} + \mathbf{f} = 0$

Conserve Energy:  $\delta T / \delta t + \mathbf{v} \cdot \nabla T = \kappa \nabla^2 T + H/C$

### 2. Solve these equations for the mantle geometry



# How to solve for mantle dynamics using a computer:

## 1. Express the physics of convection using equations

Conserve Mass:

$$\nabla \cdot \mathbf{v} = 0$$

Conserve Momentum:

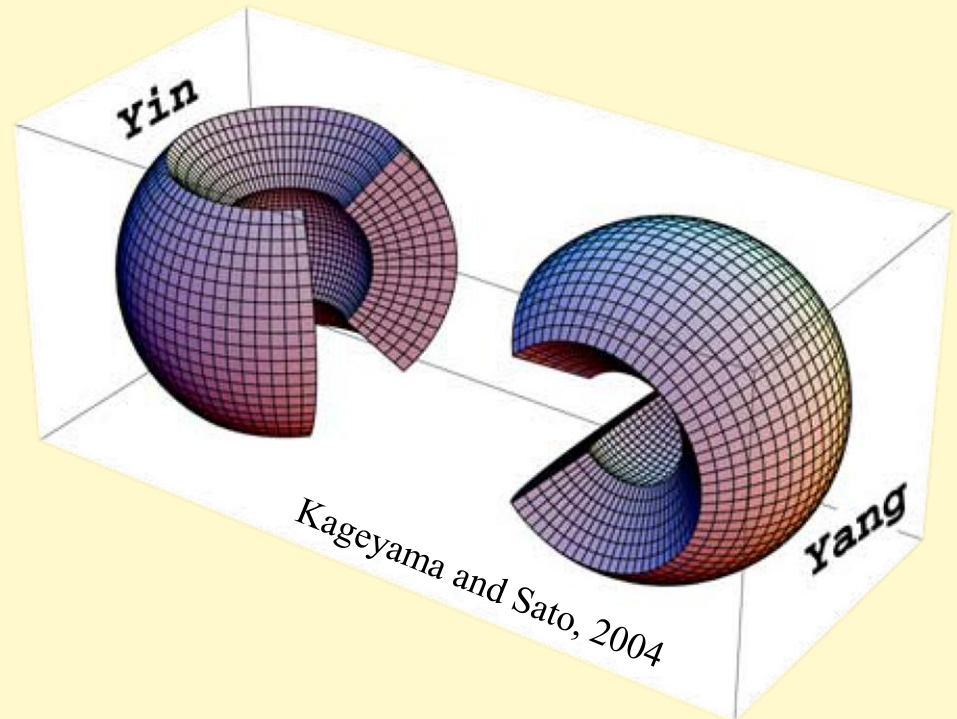
$$-\nabla p + \eta \nabla^2 \mathbf{v} + \mathbf{f} = 0$$

Conserve Energy:

$$\delta T / \delta t + \mathbf{v} \cdot \nabla T = \kappa \nabla^2 T + H/C$$

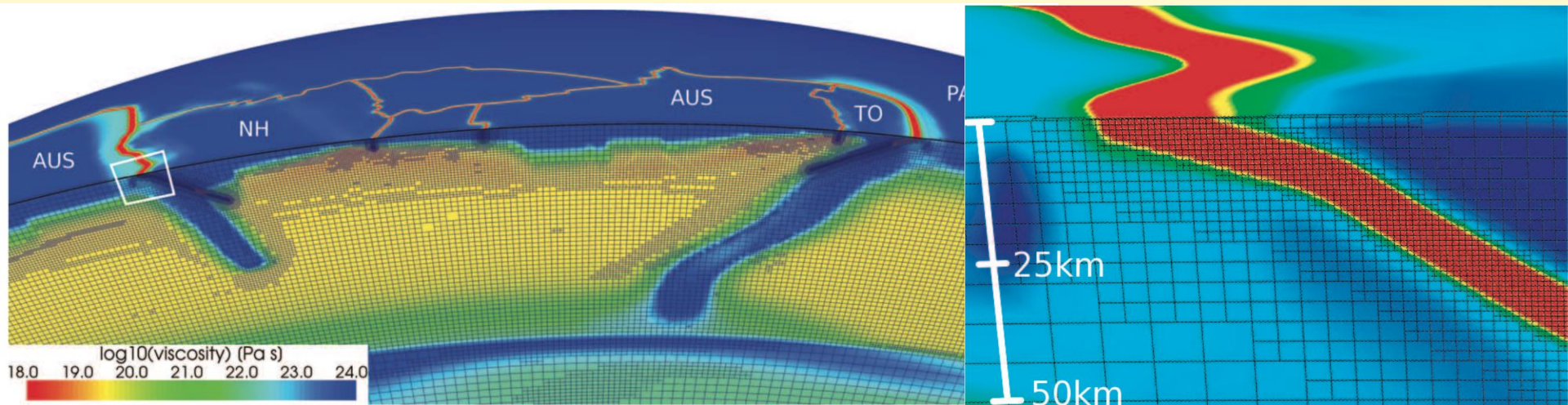
## 2. Solve these equations for the mantle geometry

An grid useful  
for the mantle



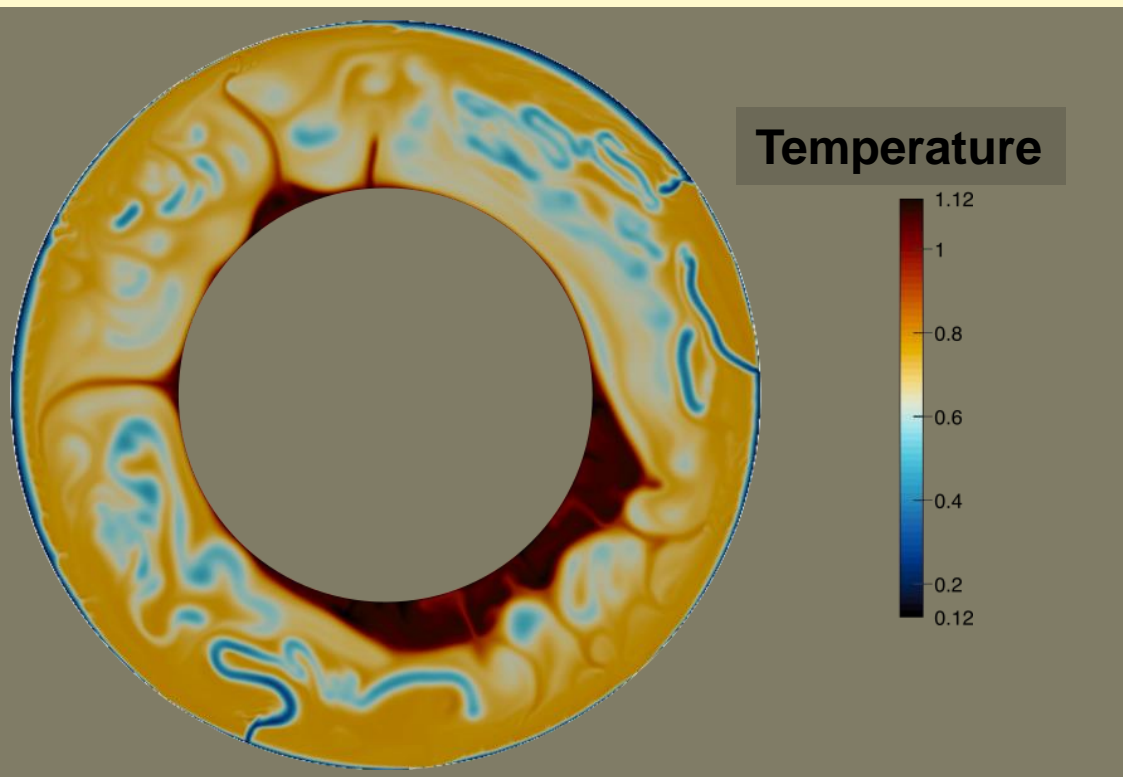
## How to solve for mantle dynamics using a computer:

1. Express the physics of convection using equations
2. Solve these equations for the mantle geometry
3. We must make some major simplifications:  
→ Our model will have less complexity than Earth



## How to solve for mantle dynamics using a computer:

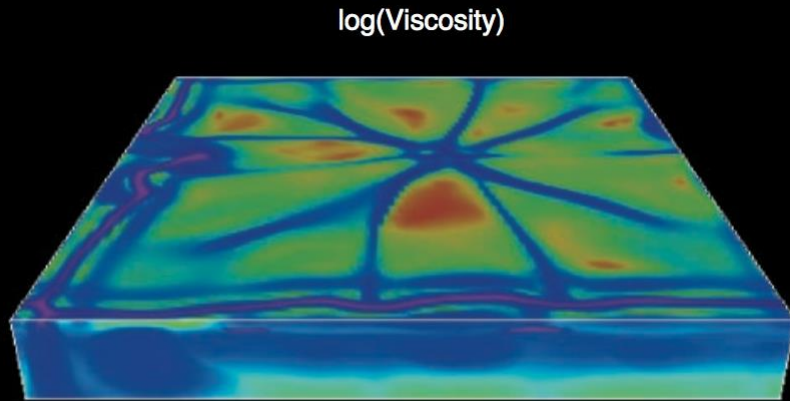
1. Express the physics of convection using equations
2. Solve these equations for the mantle geometry
2. We must make some major simplifications
3. We must make some major assumptions:
  - We can't be sure of the interior structures



## How stiff is the mantle?

- Cold temperatures make rocks stiffer
- The tectonic plates should be very stiff!

# How stiff are the plates?



**Red** → no deformation  
**Green** → slow deformation  
**Blue** → fast deformation

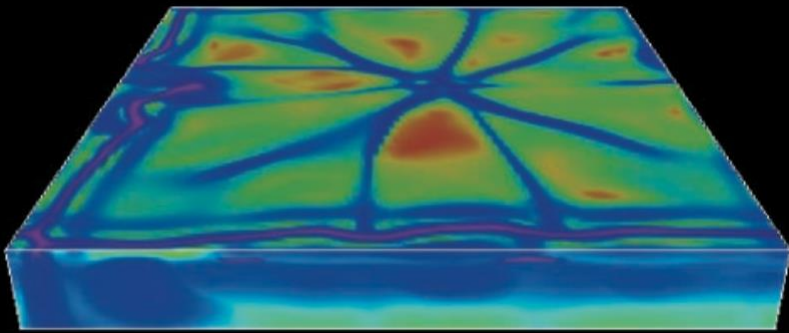
**Weak plates**

→ Deformation is distributed across the plate

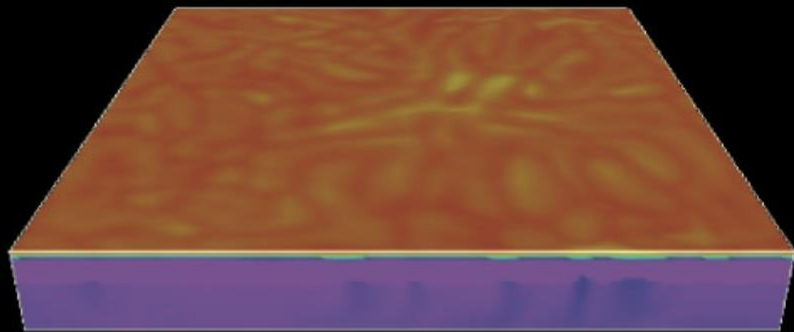


# How stiff are the plates?

log(Viscosity)



**Red** → no deformation  
**Green** → slow deformation  
**Blue** → fast deformation



*Tackley [2000]*

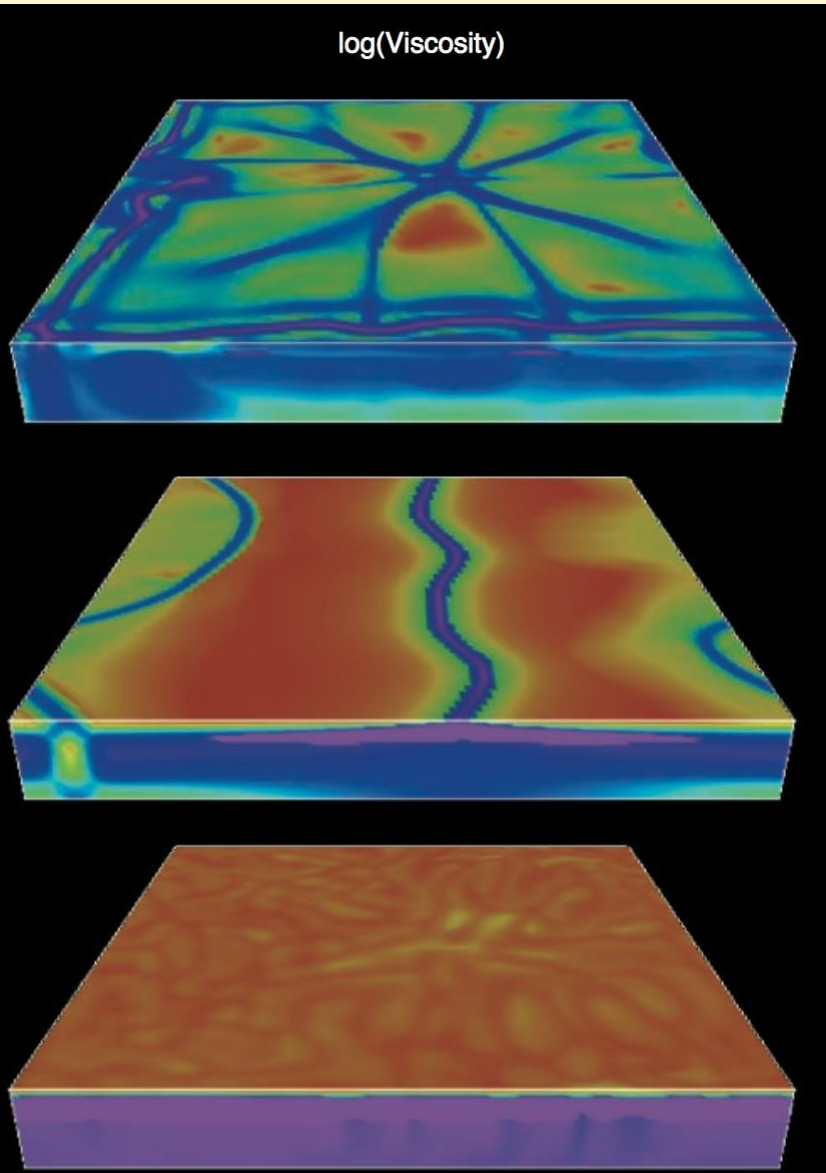
**Weak plates**

→ Deformation is distributed across the plate

**Strong plates**

→ The plates cannot deform

# How stiff are the plates?



## Weak plates

→ Deformation is distributed across the plate

## Intermediate stiffness

→ Deformation is localized at plate boundaries

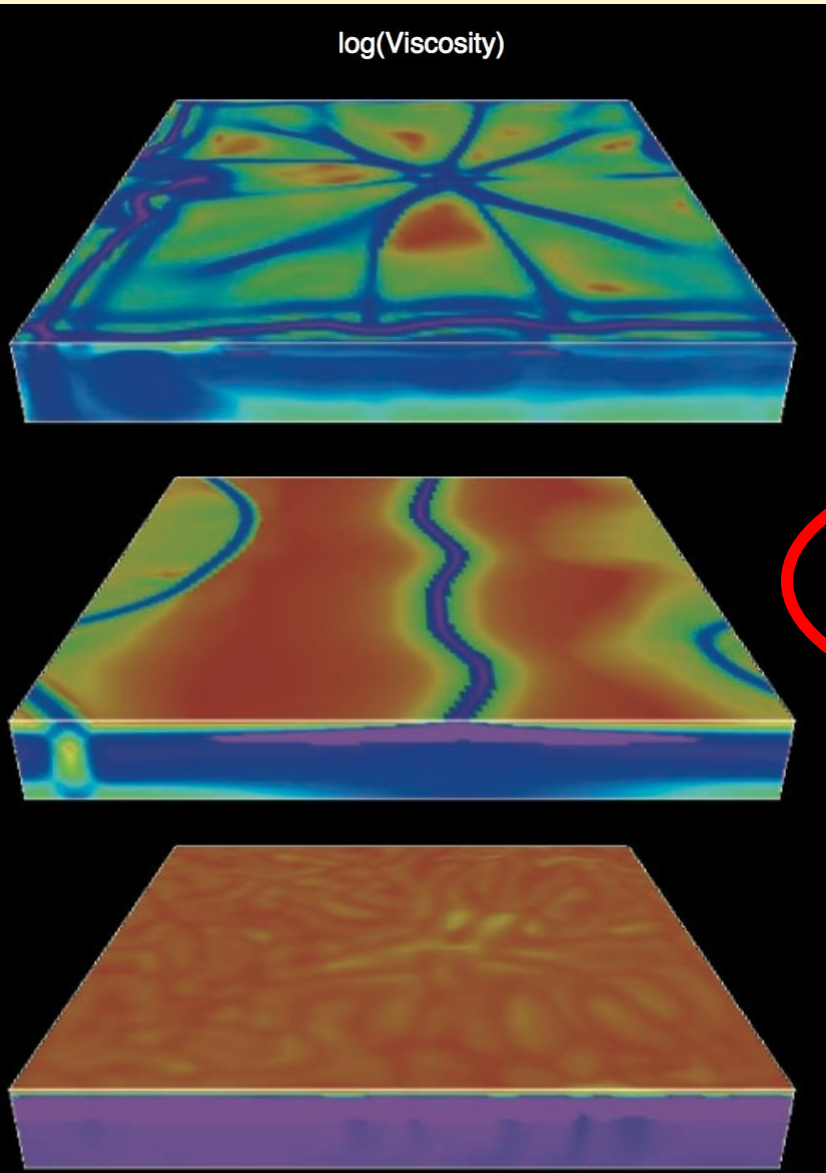
## Strong plates

→ The plates cannot deform

**Red** → no deformation  
**Green** → slow deformation  
**Blue** → fast deformation

*Tackley [2000]*

# How stiff are the plates?



Tackley [2000]

**Weak plates**

→ Deformation is distributed across the plate

**The most Earth-like model!**

**Intermediate stiffness**

→ Deformation is localized at plate boundaries

**Strong plates**

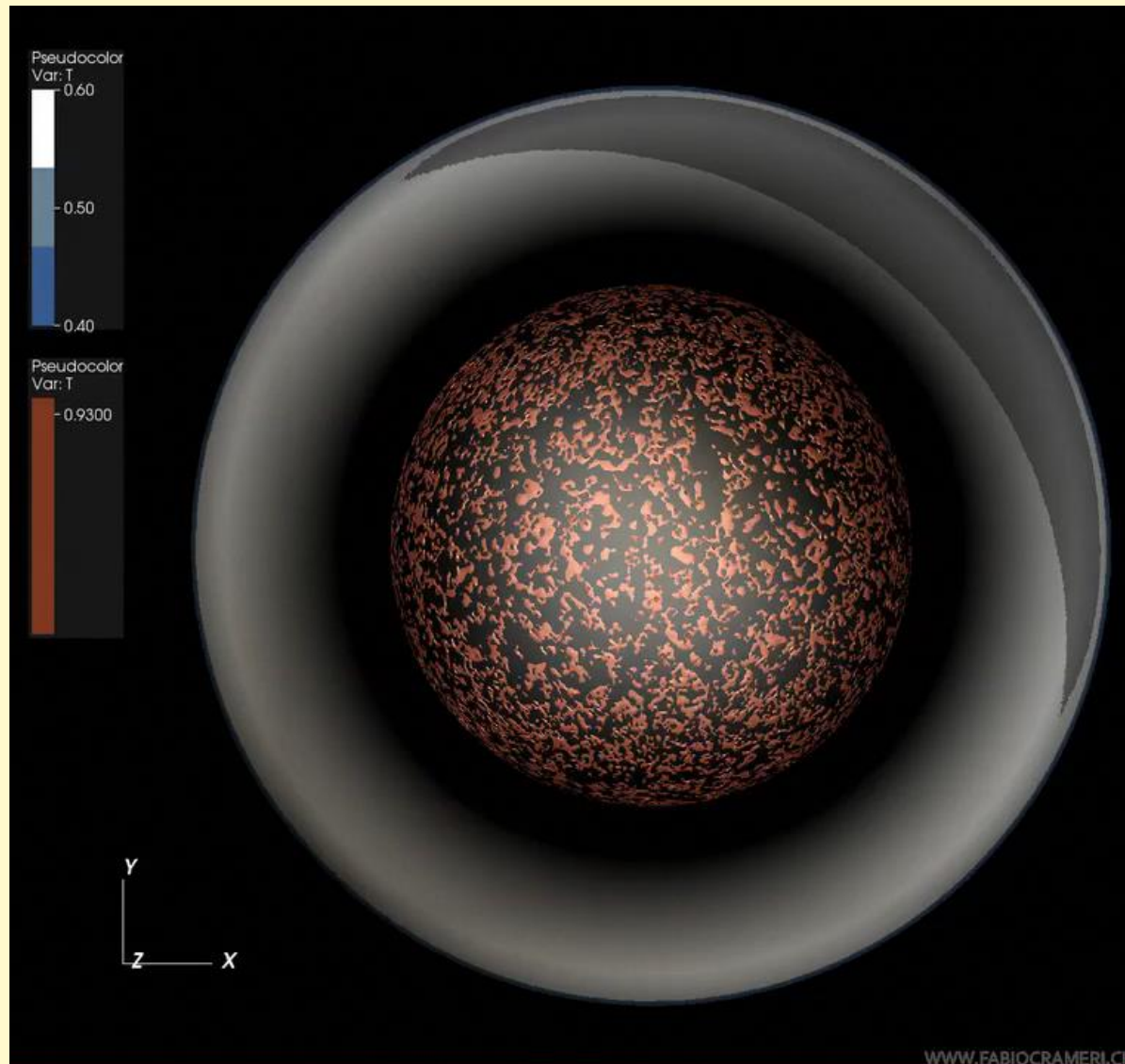
→ The plates cannot deform

**Red** → no deformation

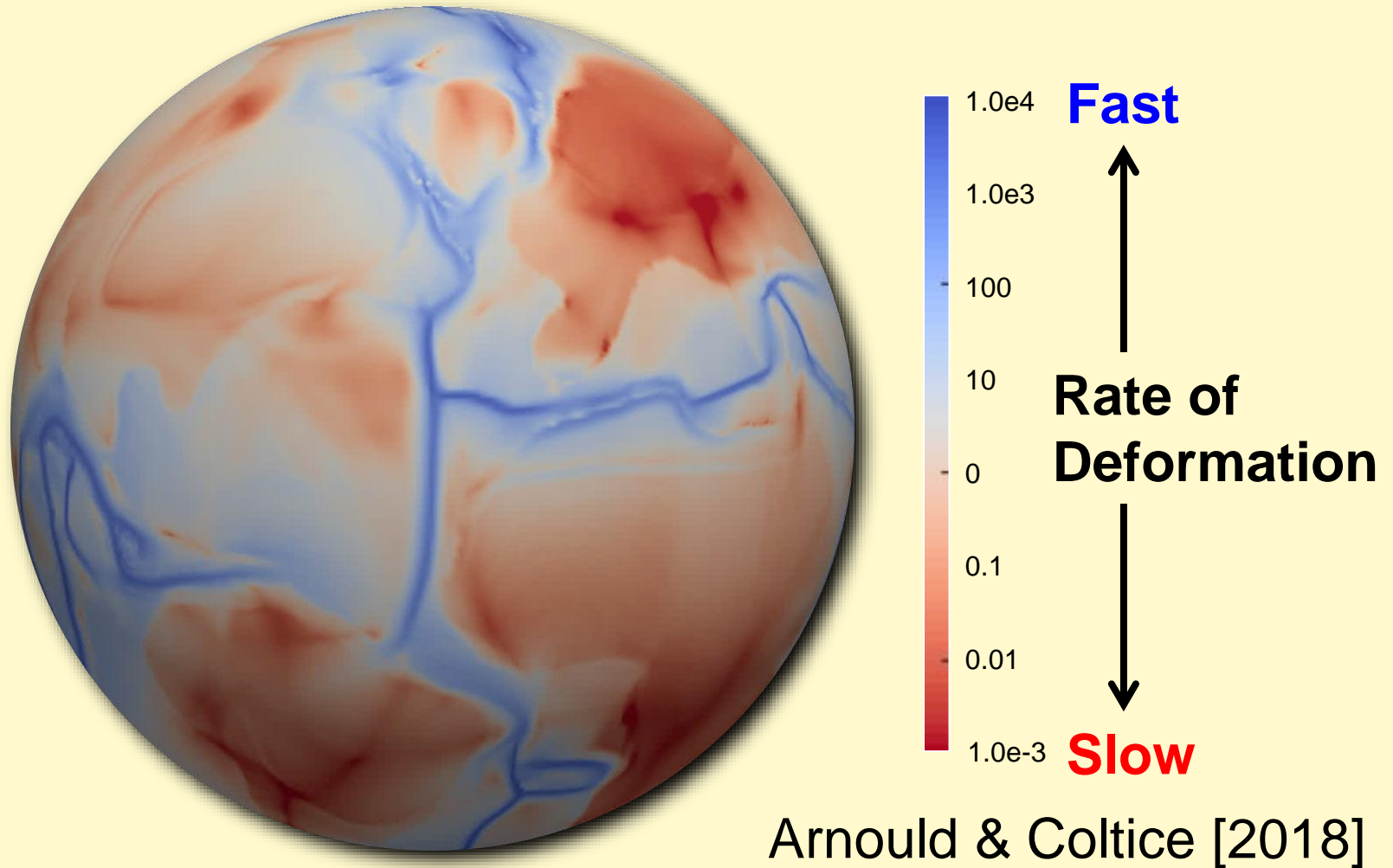
**Green** → slow deformation

**Blue** → fast deformation

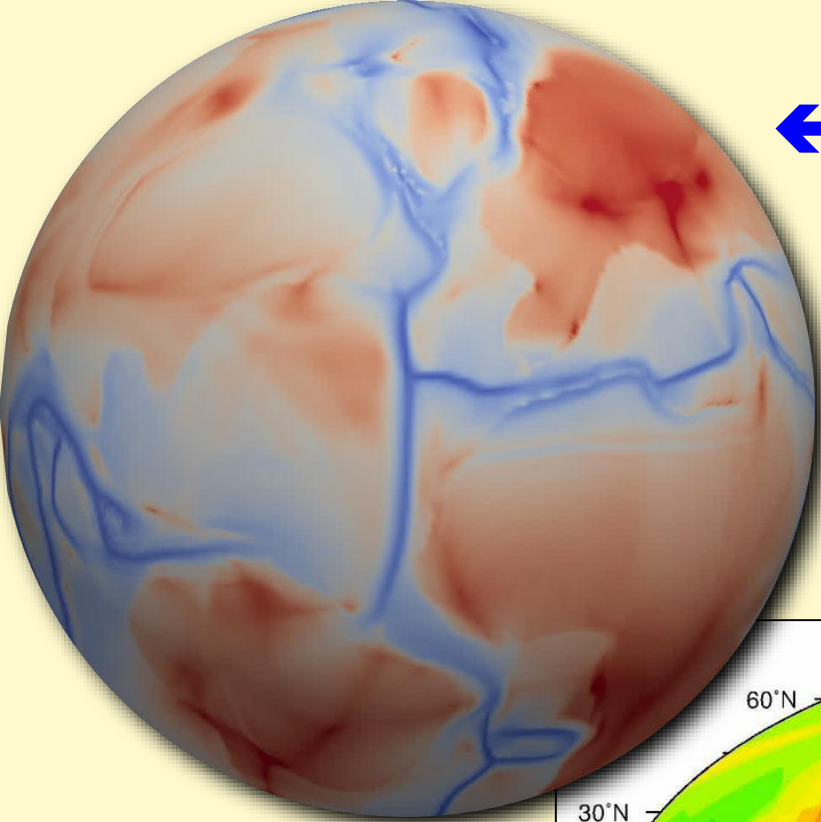
# Mantle Convection: The role of plate motions



# Testing Mantle Convection Models



**Is this a useful representation of plate tectonics?**  
→ How to test this model?

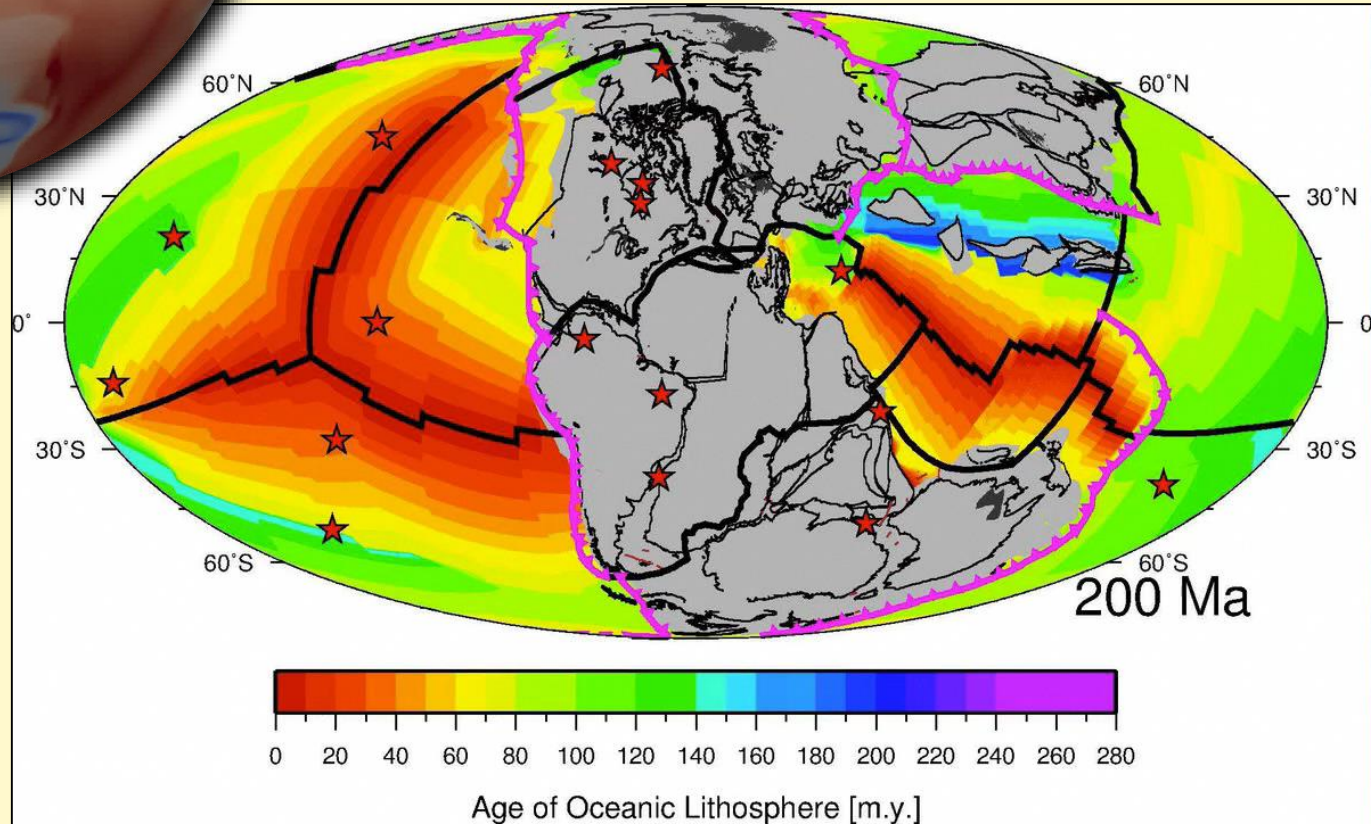


## ← Mantle Flow Model

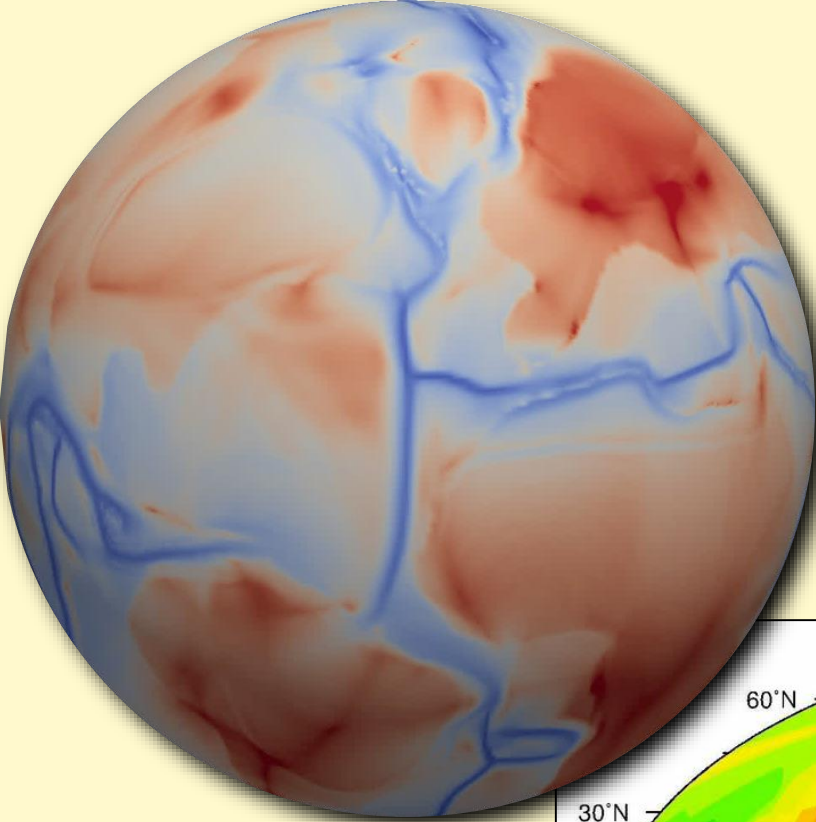
How do they compare?

Observation:  
Plate Reconstruction ↓

[Arnould &  
Coltice, 2018]



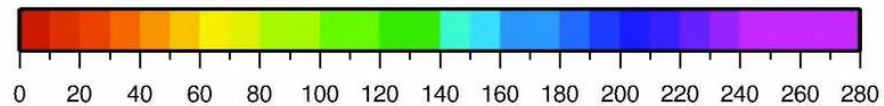
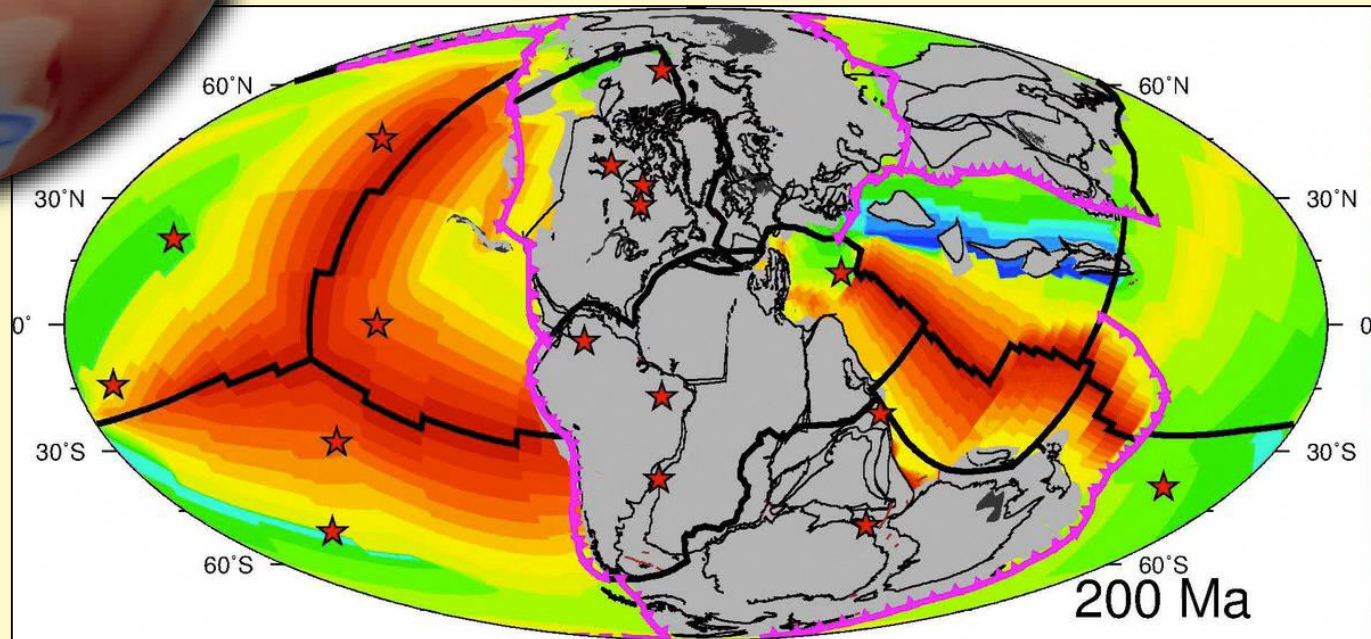
Seton et al. [2012]



## 7 comparisons:

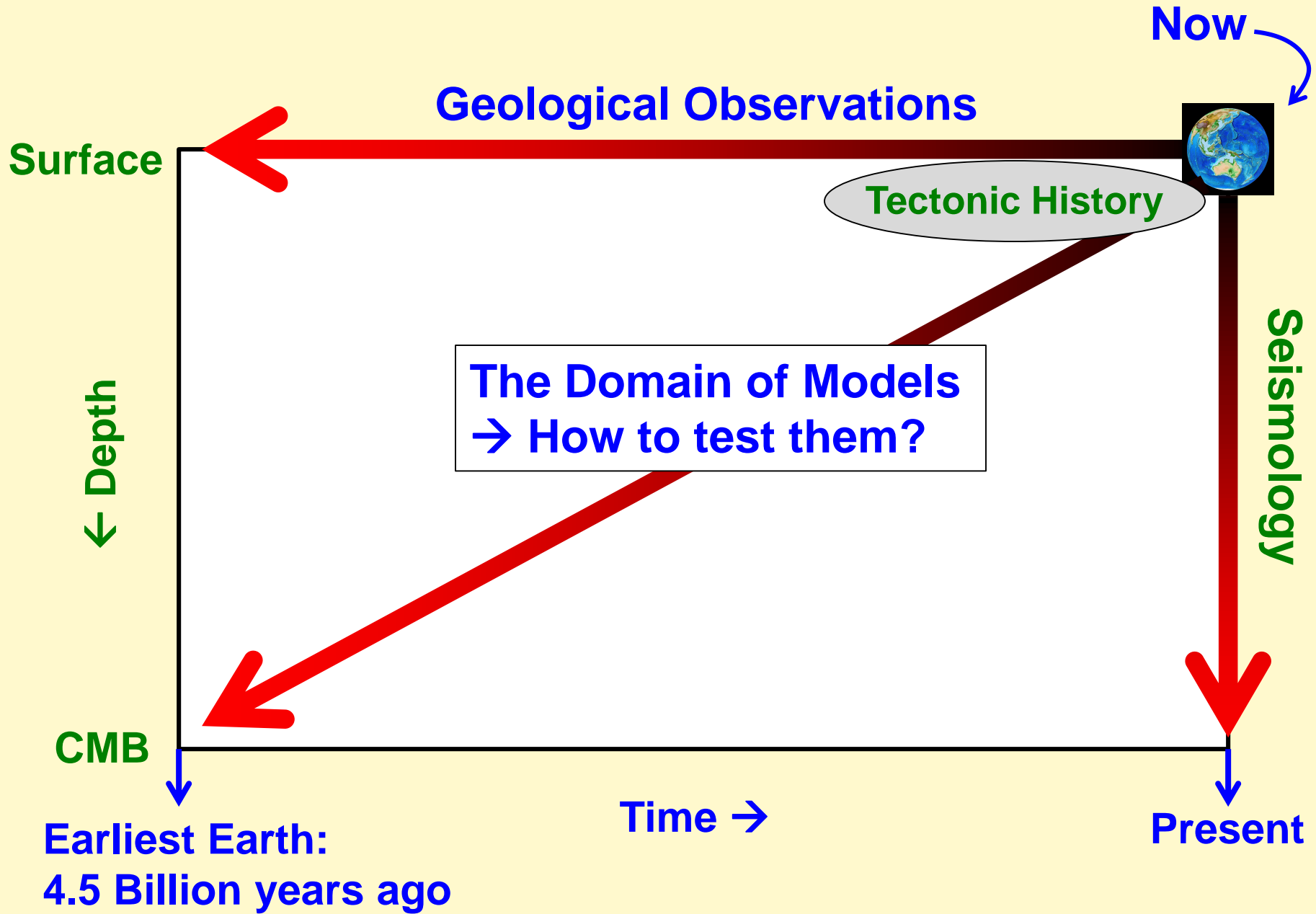
- Spacing of boundaries
- Length & Width of boundaries
- Duration of boundaries
- Deformation rate
- Migration rate of boundaries
- Boundary type

[Arnould & Coltice, 2018]

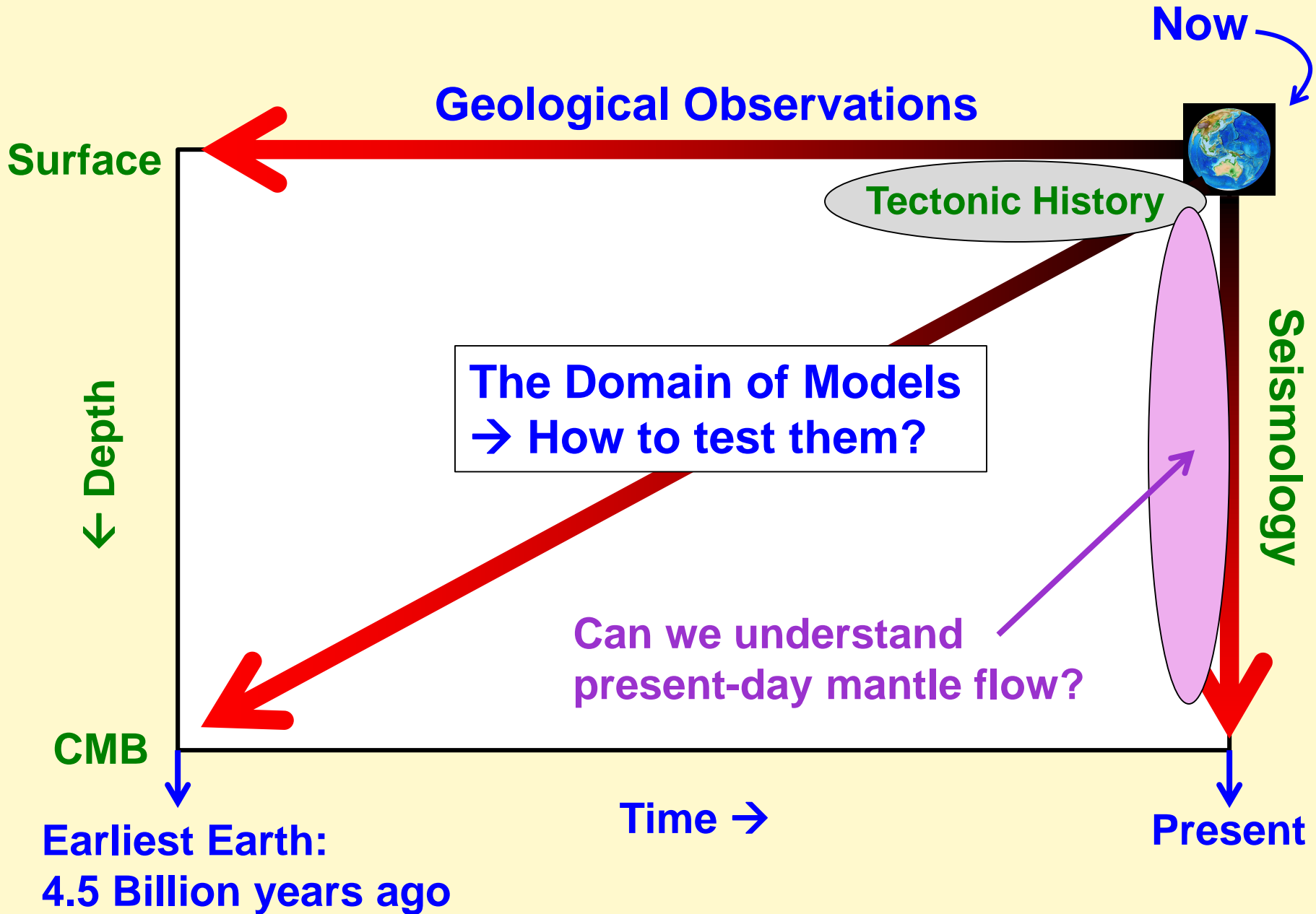


Age of Oceanic Lithosphere [m.y.]

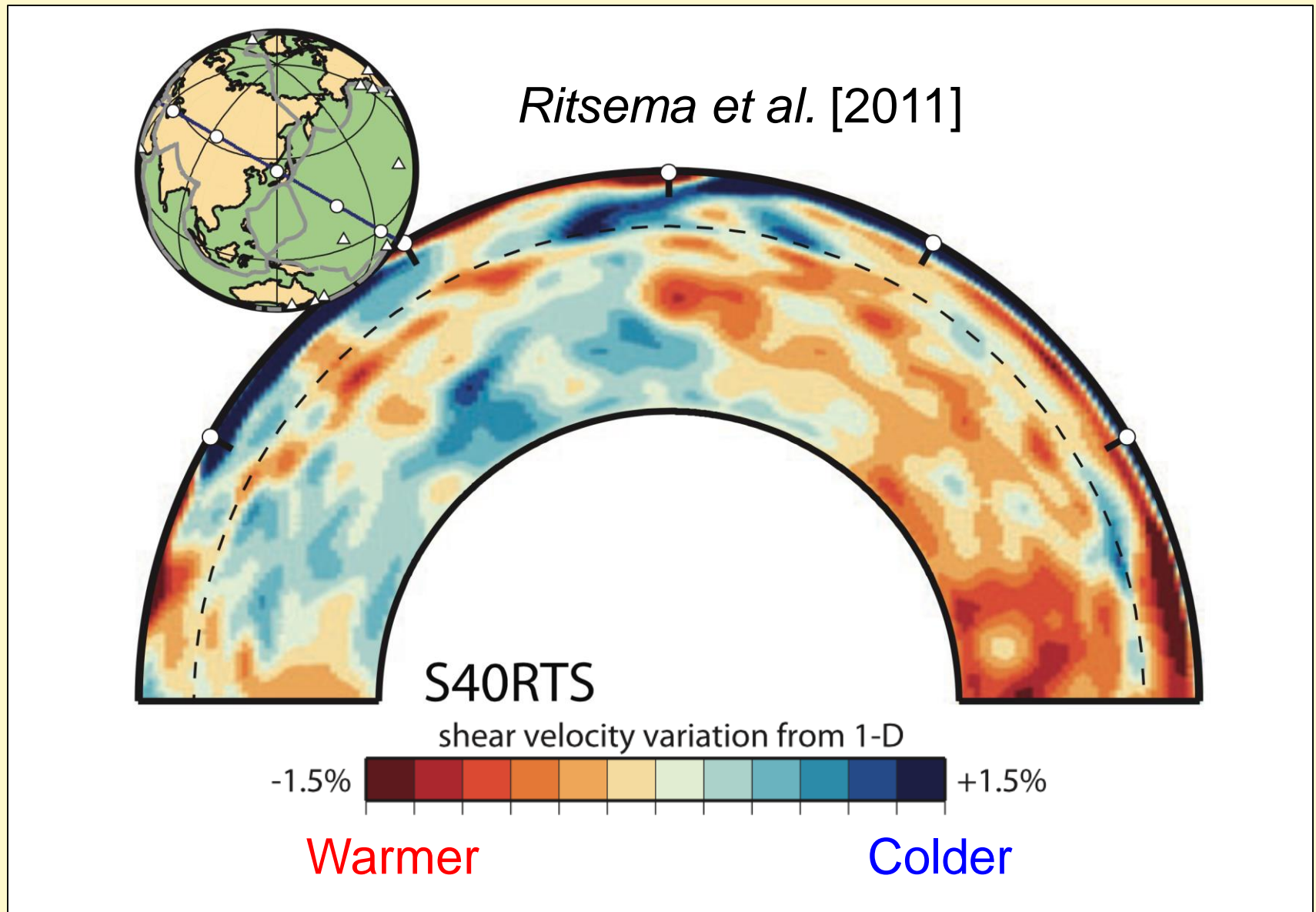
Seton et al. [2012]







# Seismology shows us a snapshot of mantle structures → What is the flow pattern?



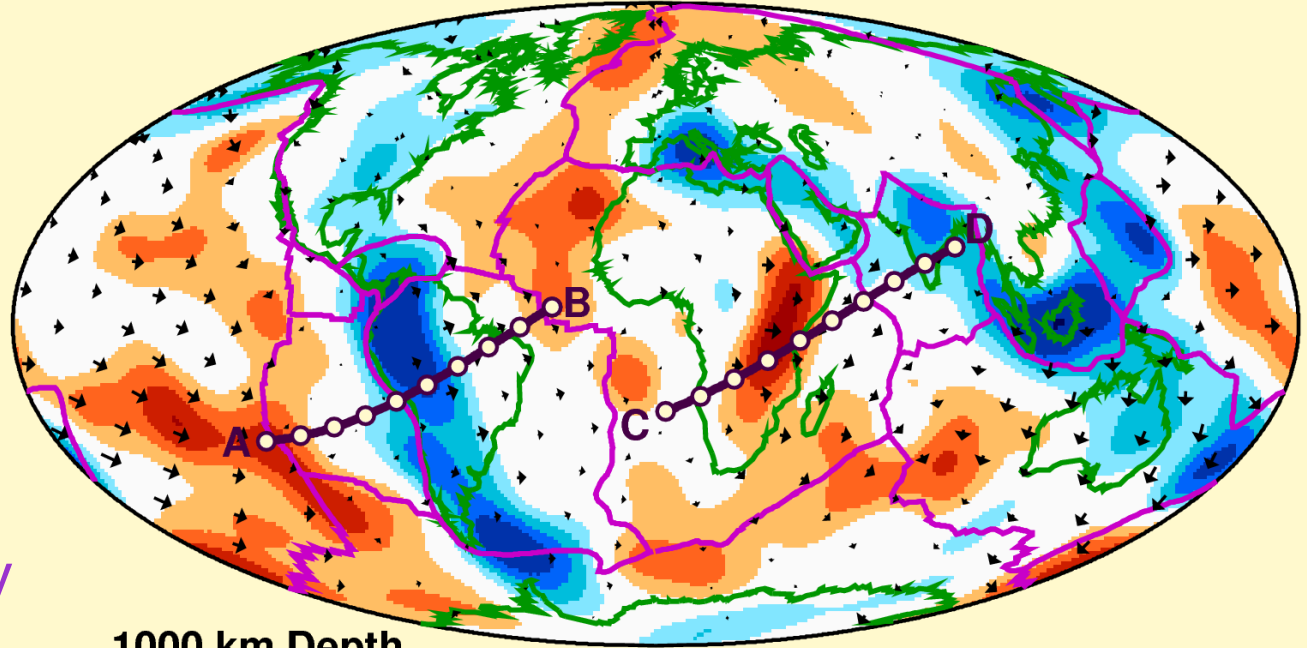
# Global Mantle Circulation Models

## Input:

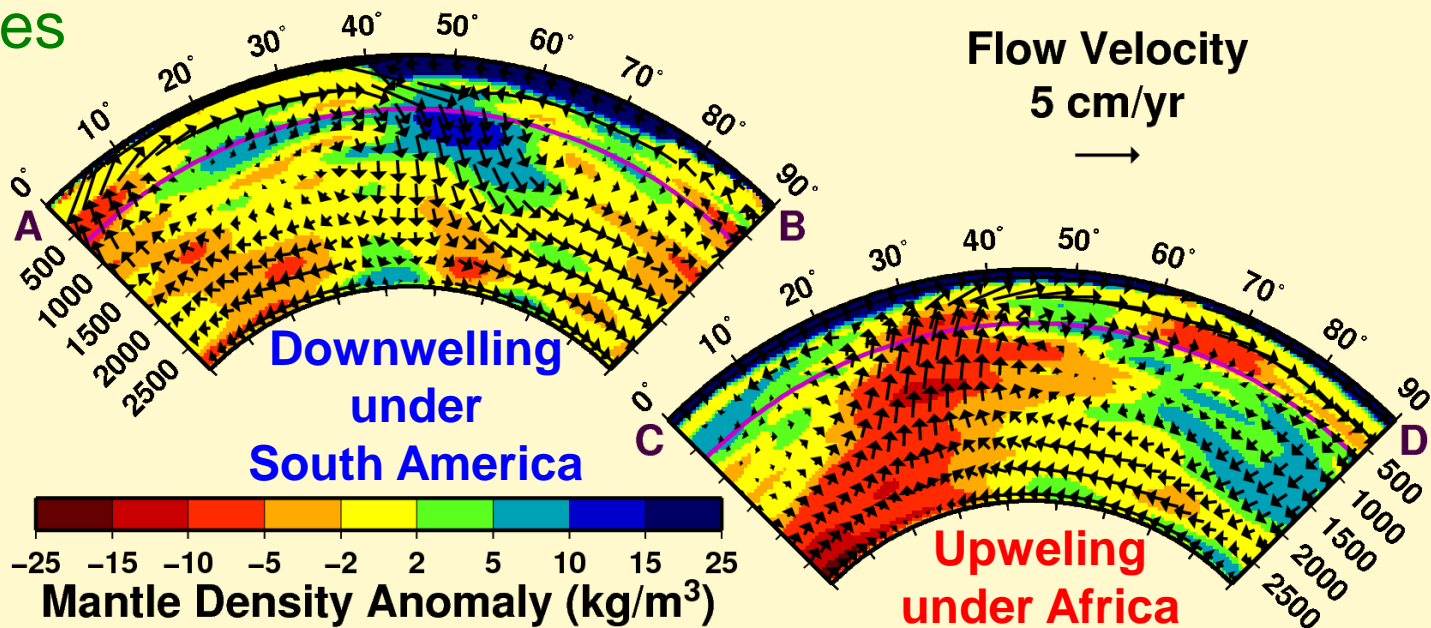
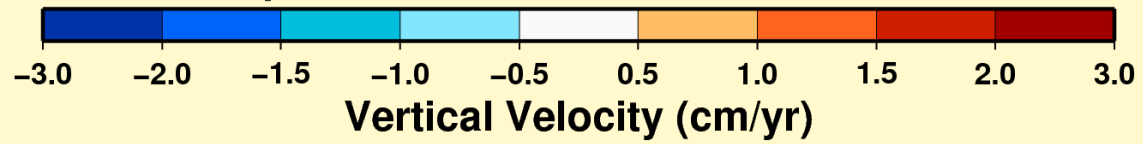
- Mantle Densities
- Mantle Viscosity

## Output of Model:

- Mantle Flow
- Tectonic Forces
- Surface Deformation



1000 km Depth



# Global Mantle Circulation Models

## Input:

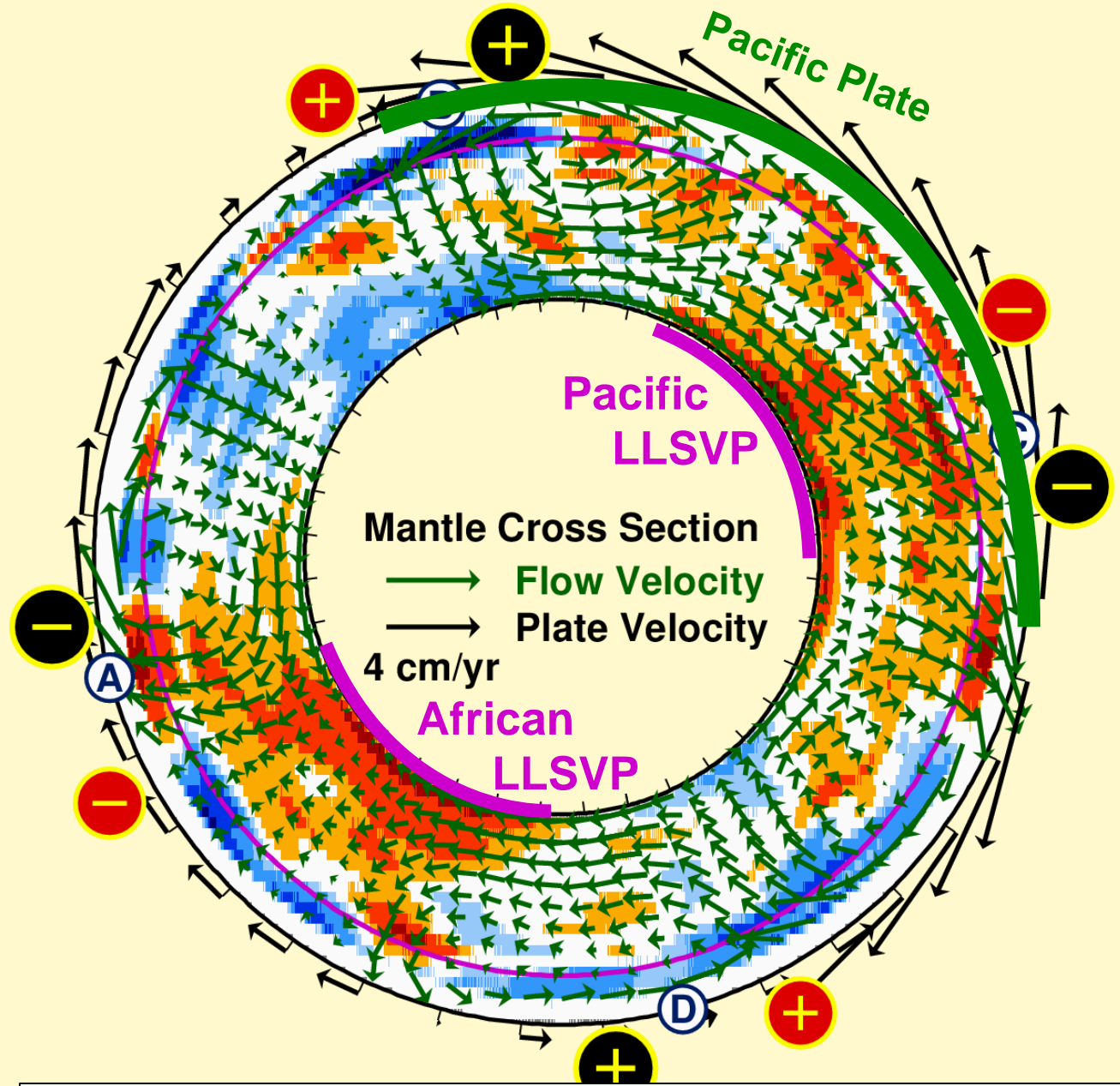
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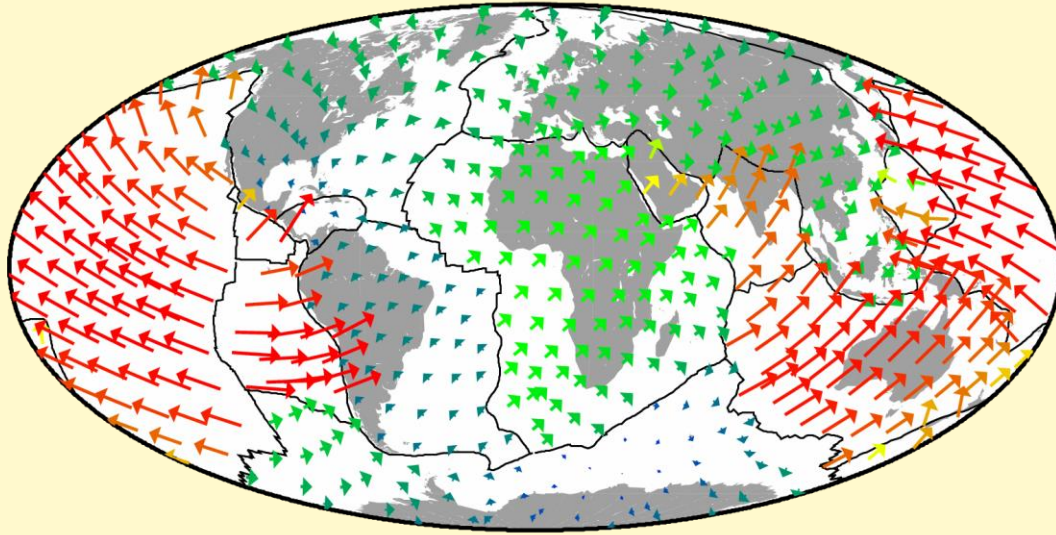


Compare to Present-day Observations



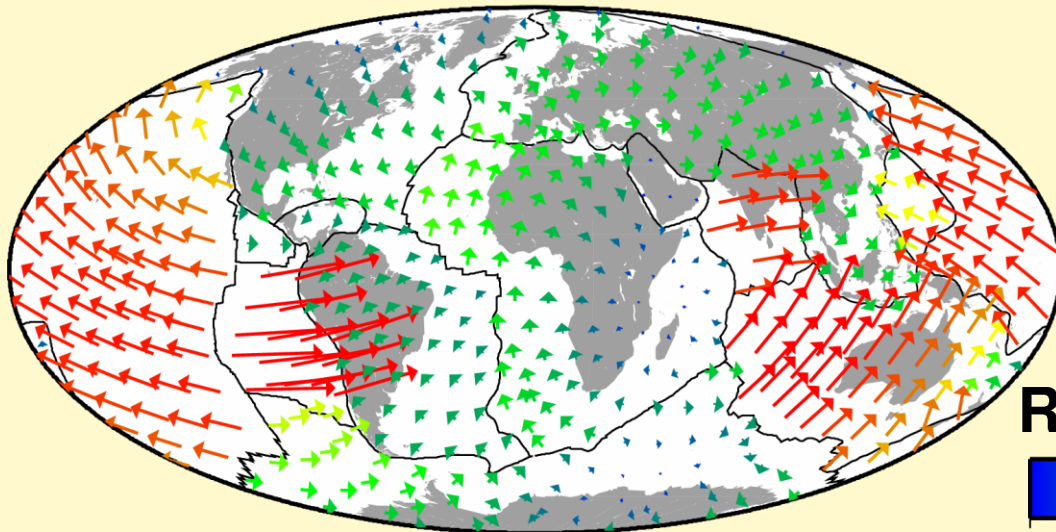
LLSVP = Large Low Slow-Velocity Province  
→ Continent-sized structures on the CMB

# Observed Plate Motions

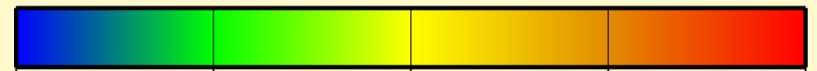


The flow model predicts plate motions!

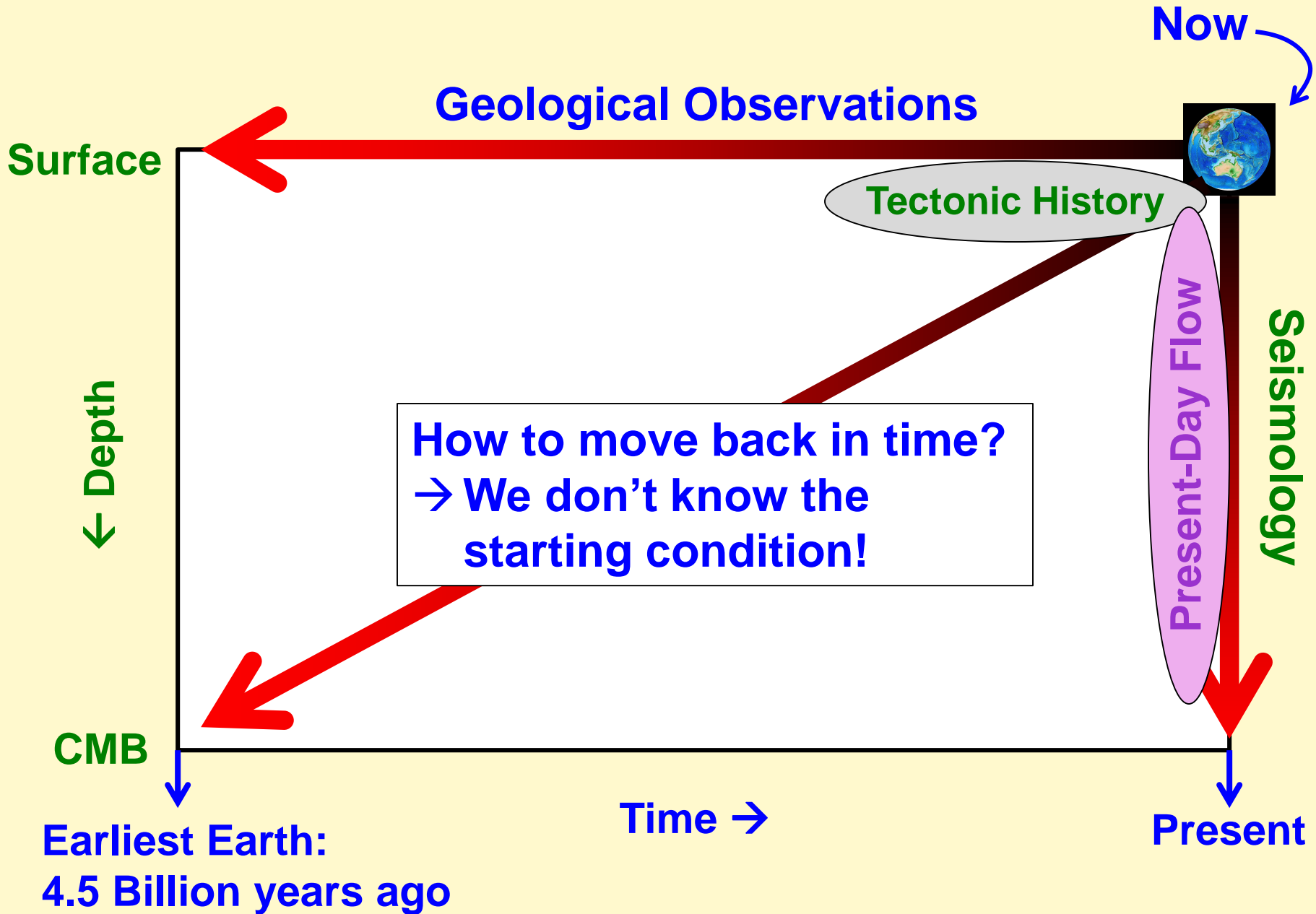
# Predicted Plate Motions

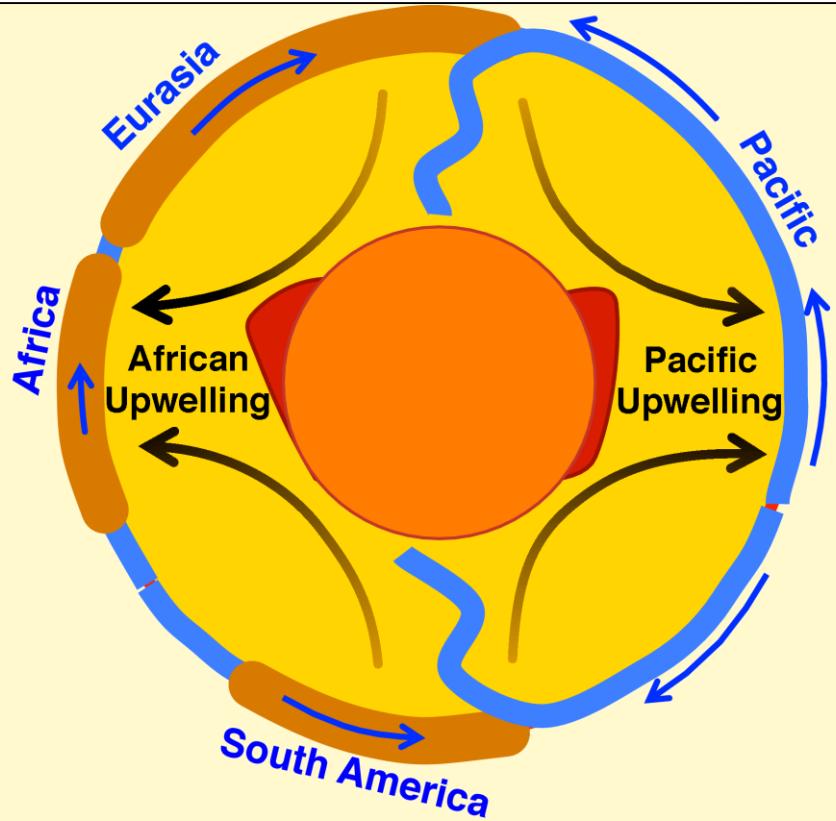
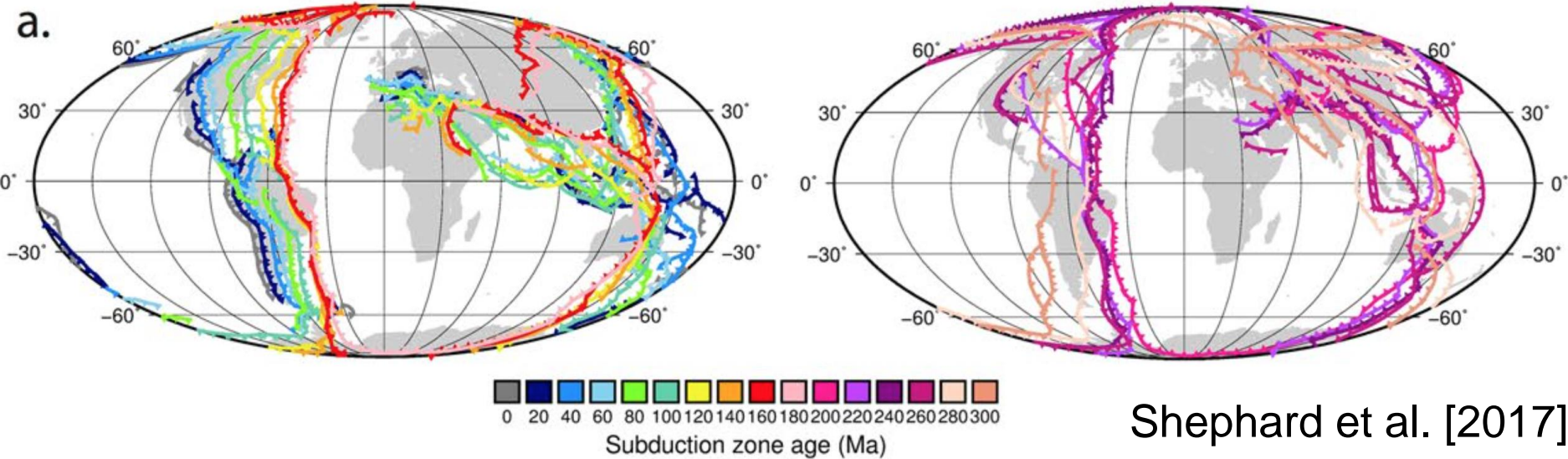


Relative Velocity Magnitude



0.00 0.75 1.00 1.25 2.00



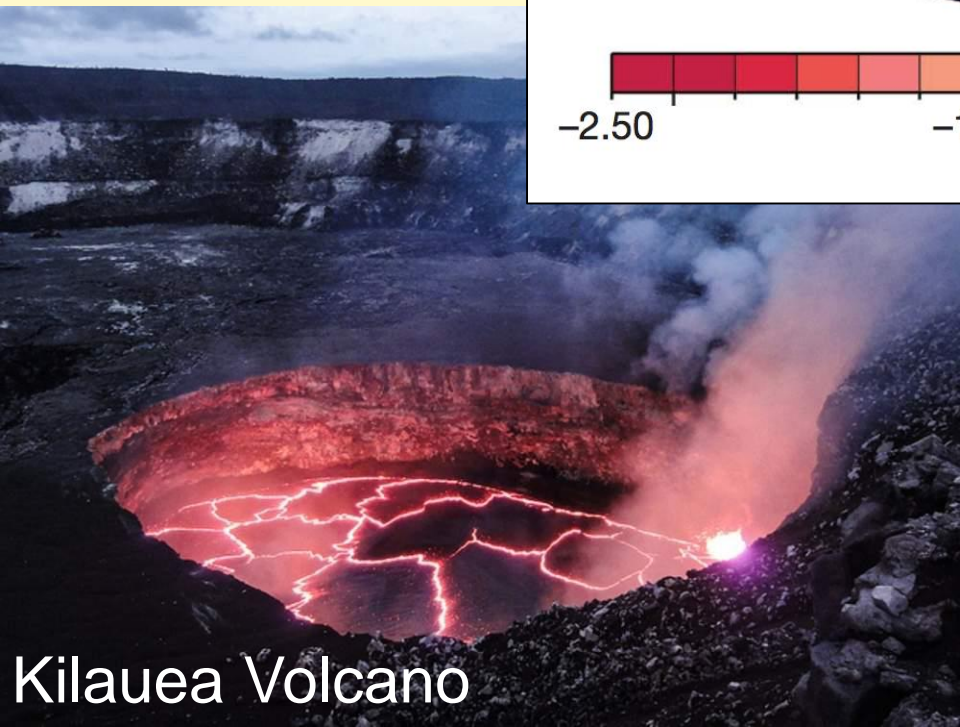
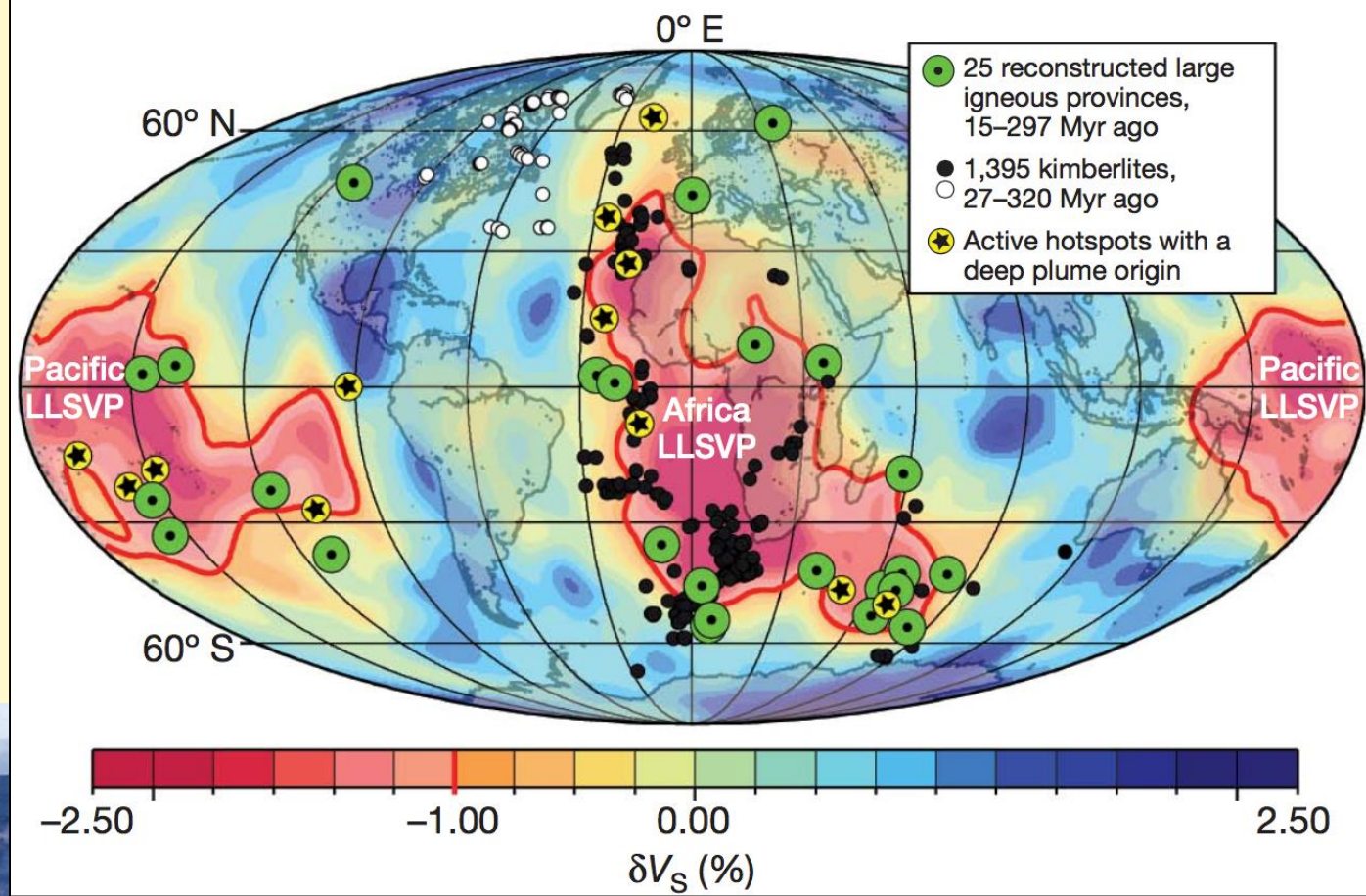


Looking backward in time

→ Subduction locations have mostly remained the same for 300 Myr!

# Long-term stability for intraplate volcanism

- Hotspots
- Flood Basalts
- Diamonds



Kilauea Volcano

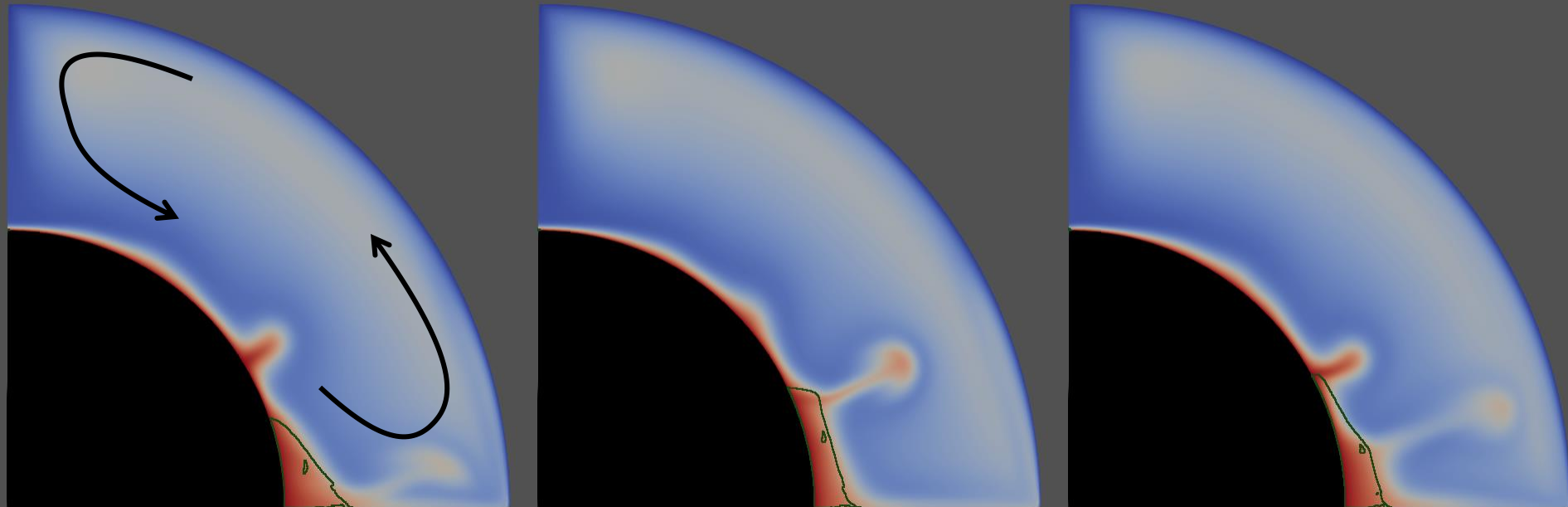
**Plumes rise from the edges of the LLSVPs (dense regions on the core-mantle boundary)**

*Torsvik et al., [Nature, 2010]*



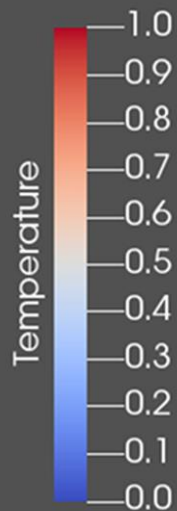
# Why do plumes rise from the LLSVP edges?

*Heyn et al. [2018]*

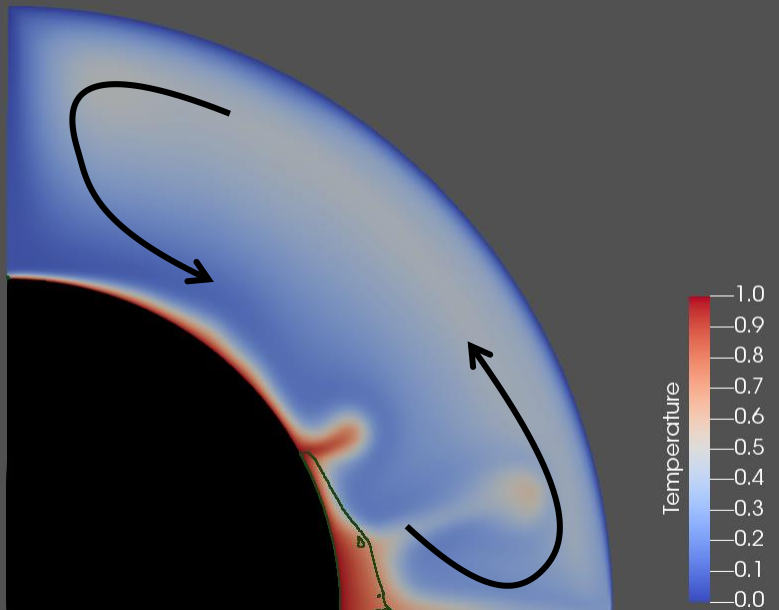
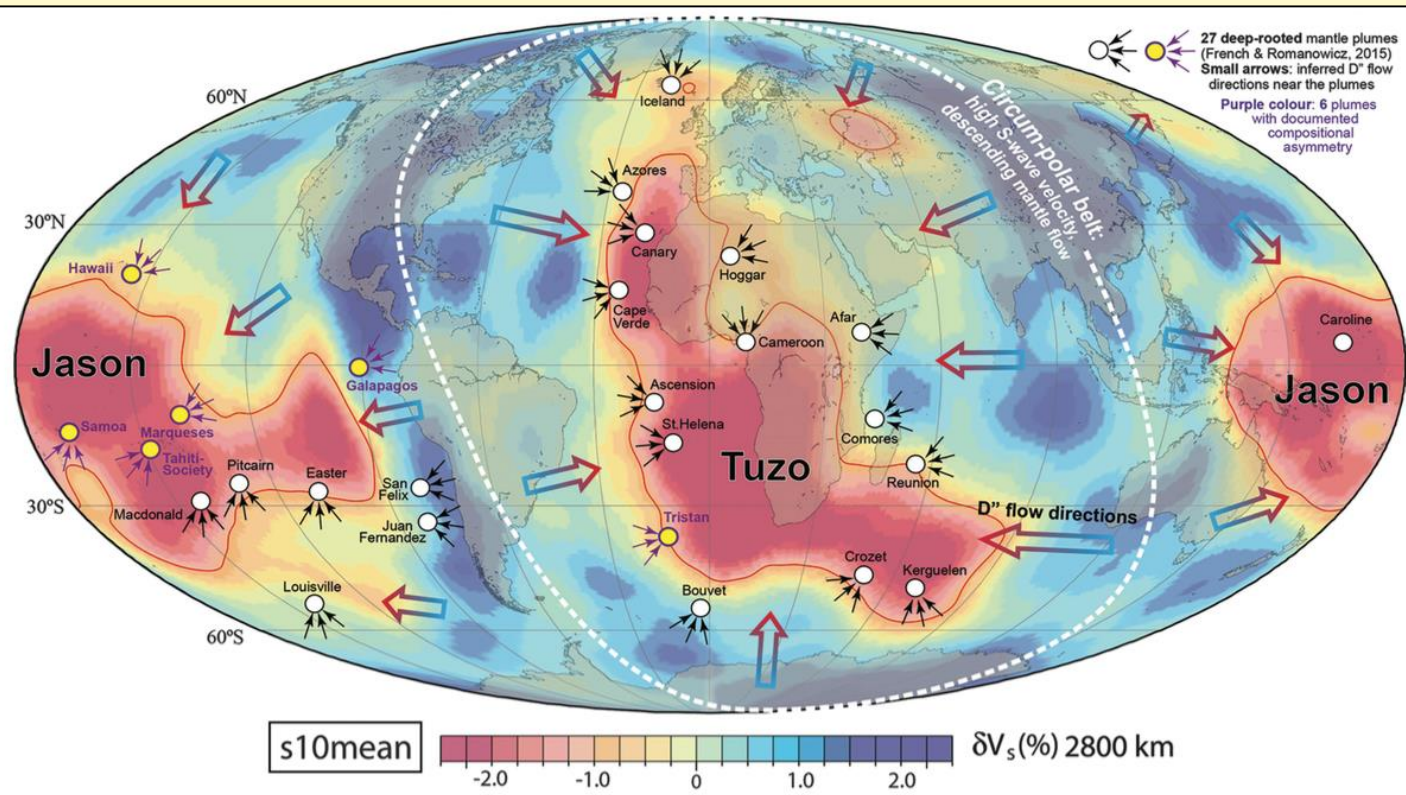


Progression of time →

- Plumes erupt from the LLSVP edges
- LLSVPs remain stable (!)



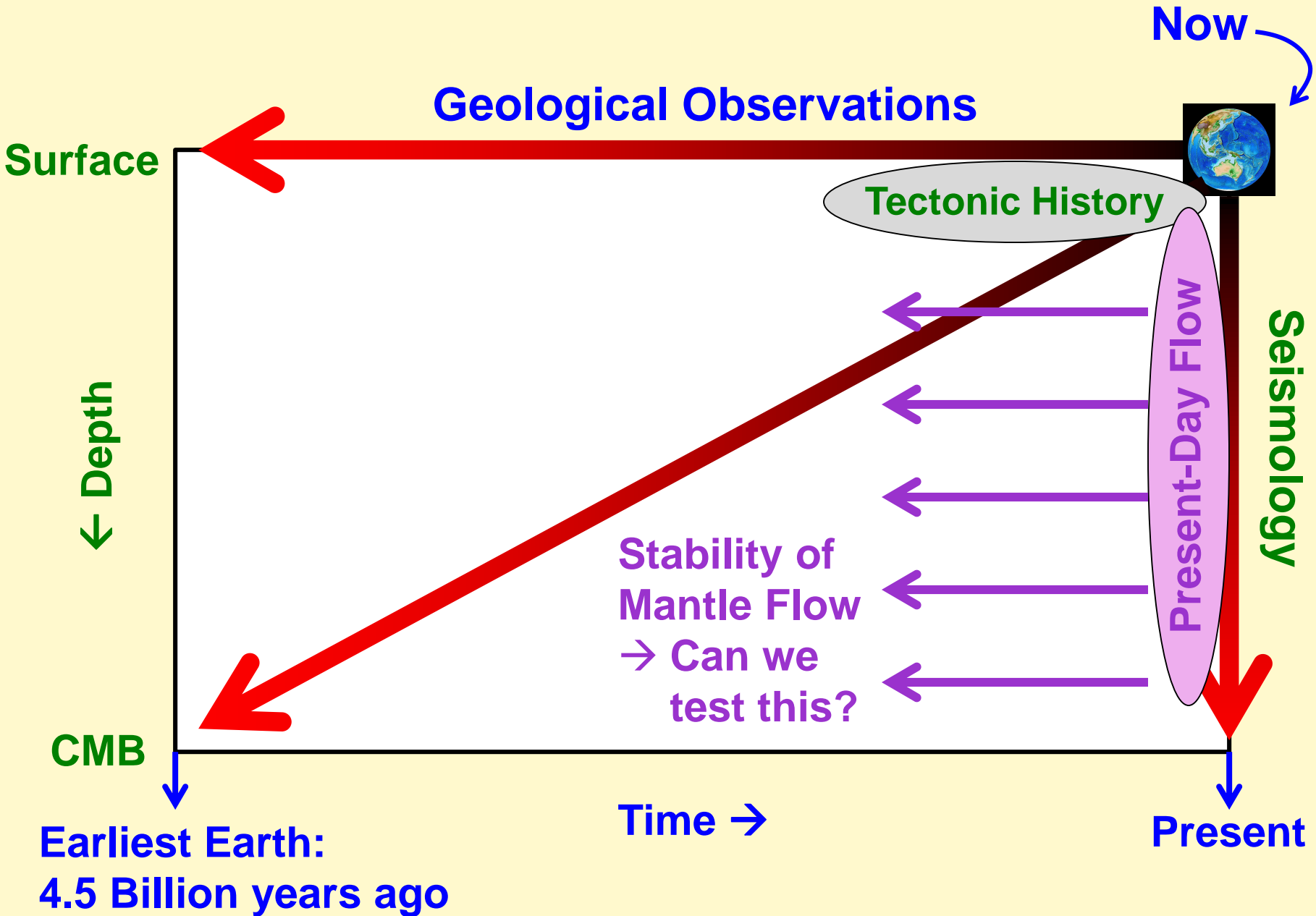
Torsvik et al. [2016]

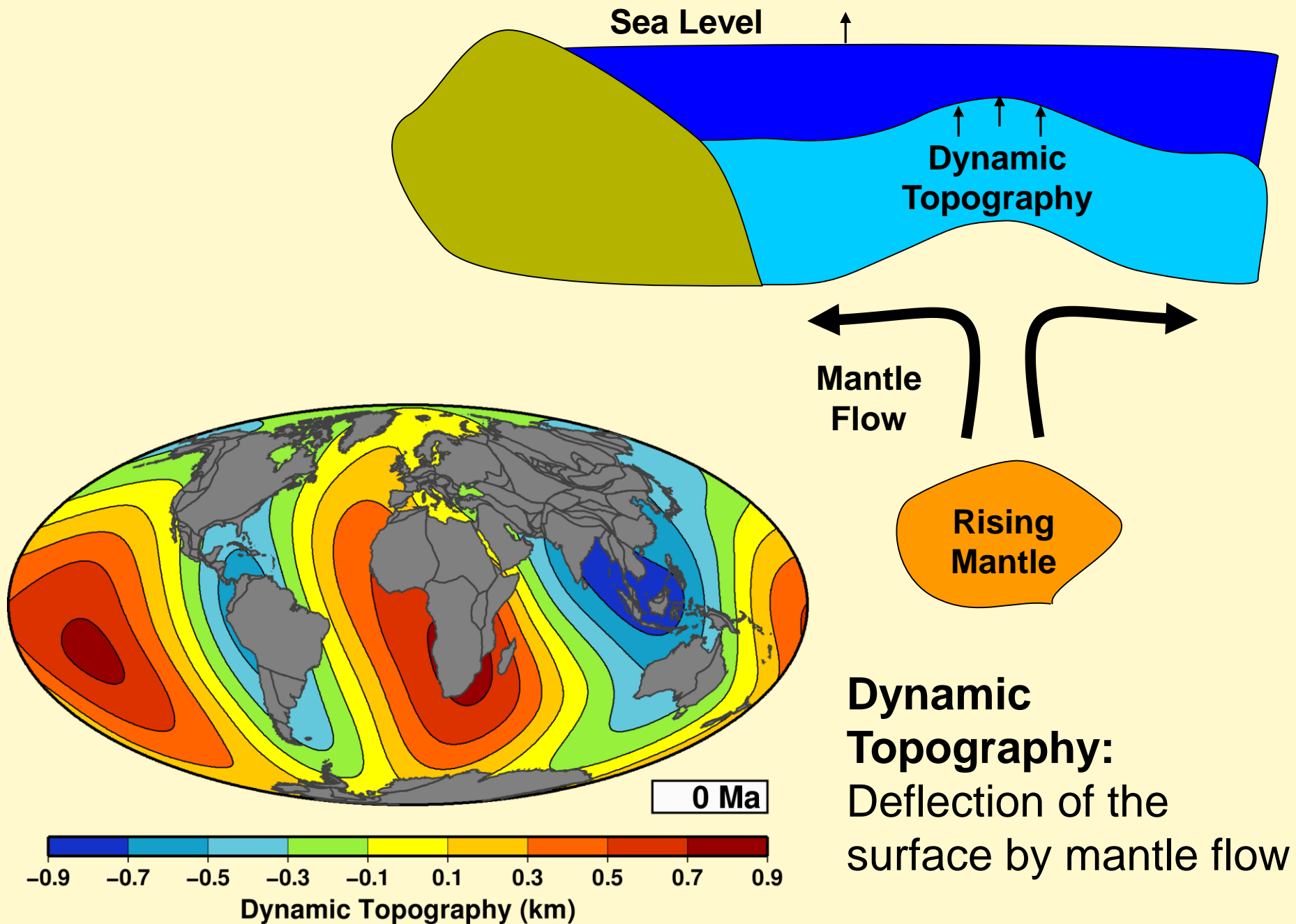


## Long-Term Stability for Mantle Dynamics (!!!)

- Stable LLSVPS
- Stable circumpolar belt of subduction

→ **Stable flow pattern:**  
**For how long?**

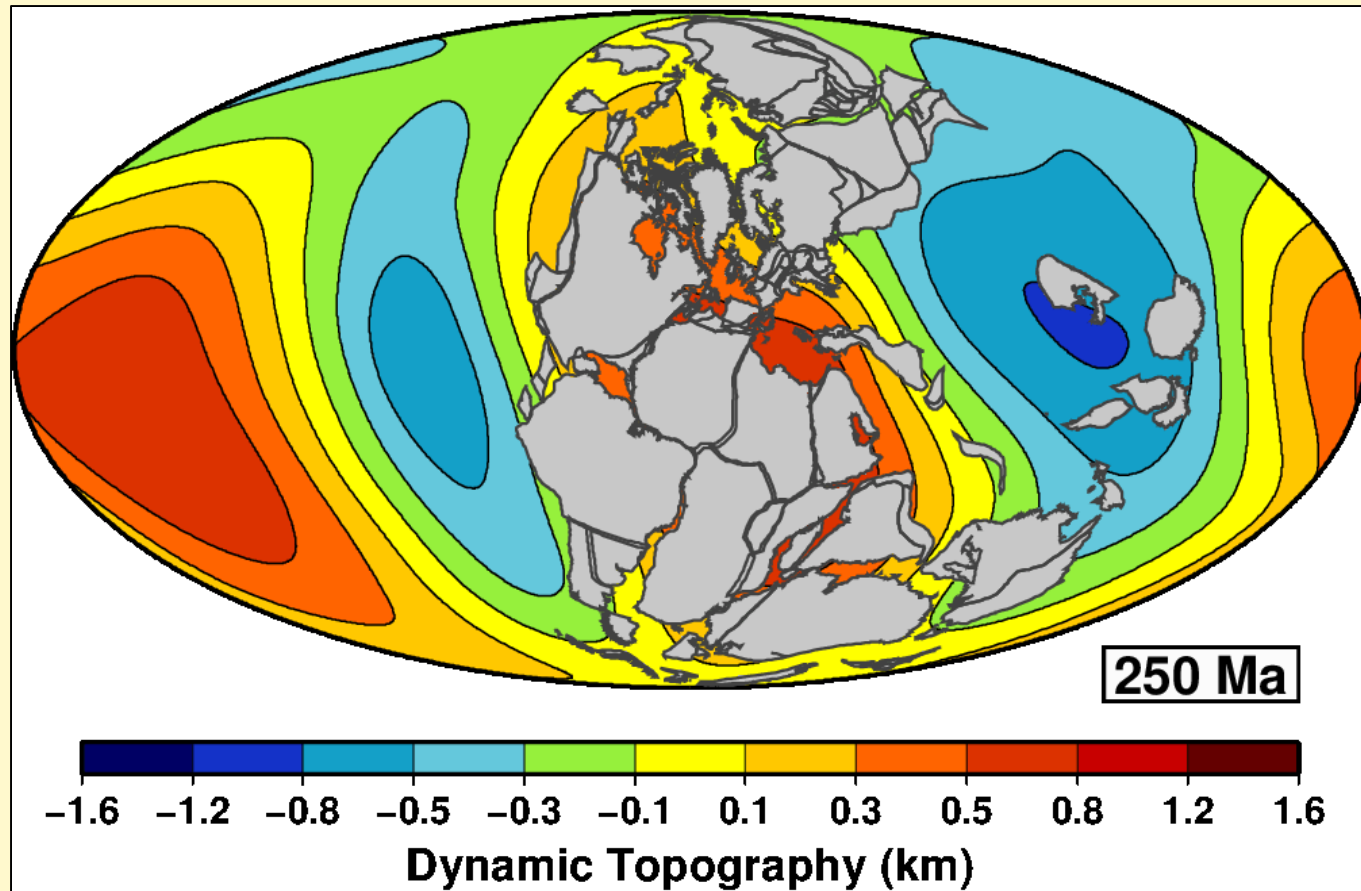




# The continents move relative to the stable mantle flow.

→ They should uplift or subside as they drift over the flow pattern.

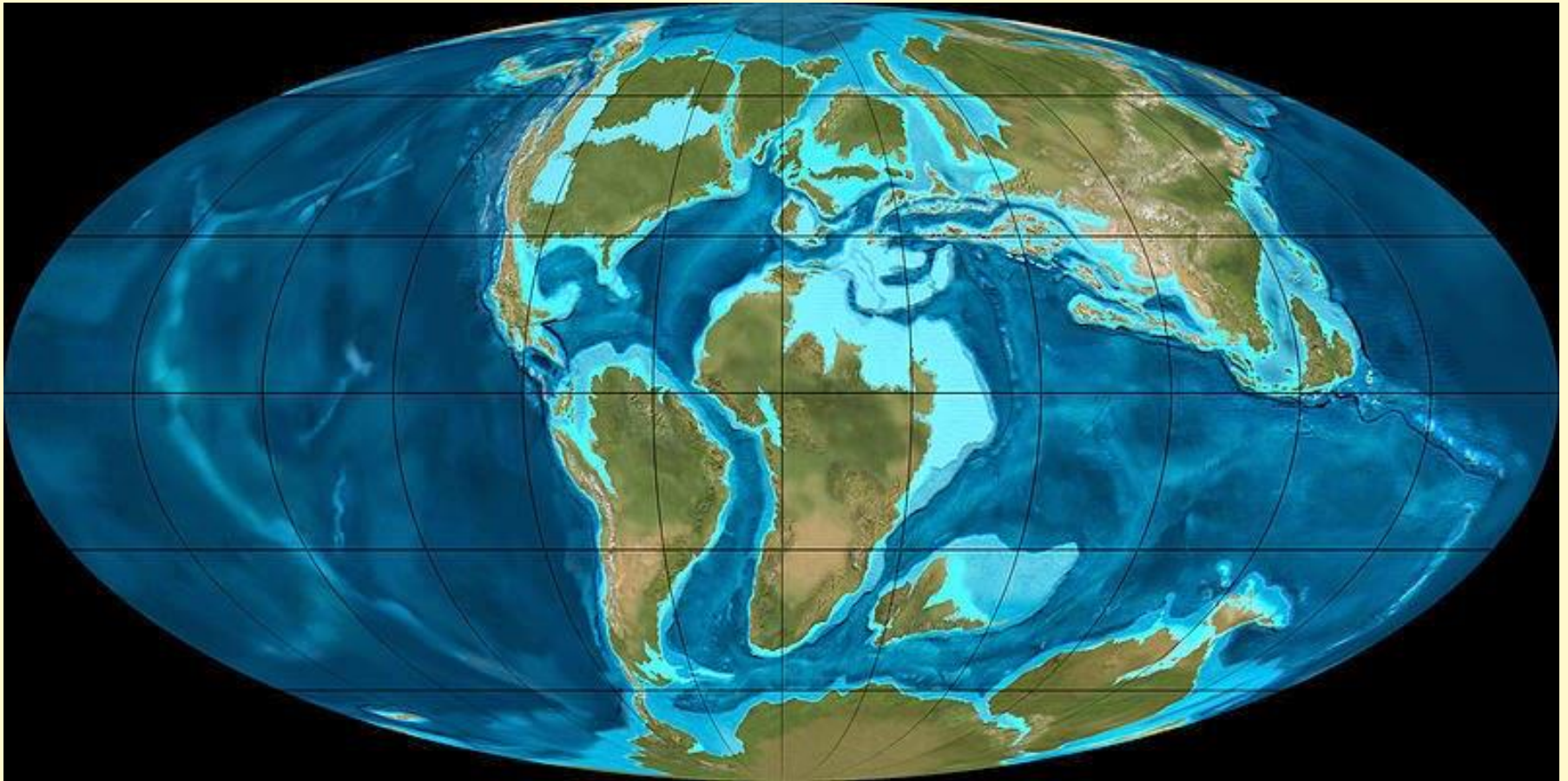
→ Compare to flooding observations on continents.



**The continents move relative to the stable mantle flow.**

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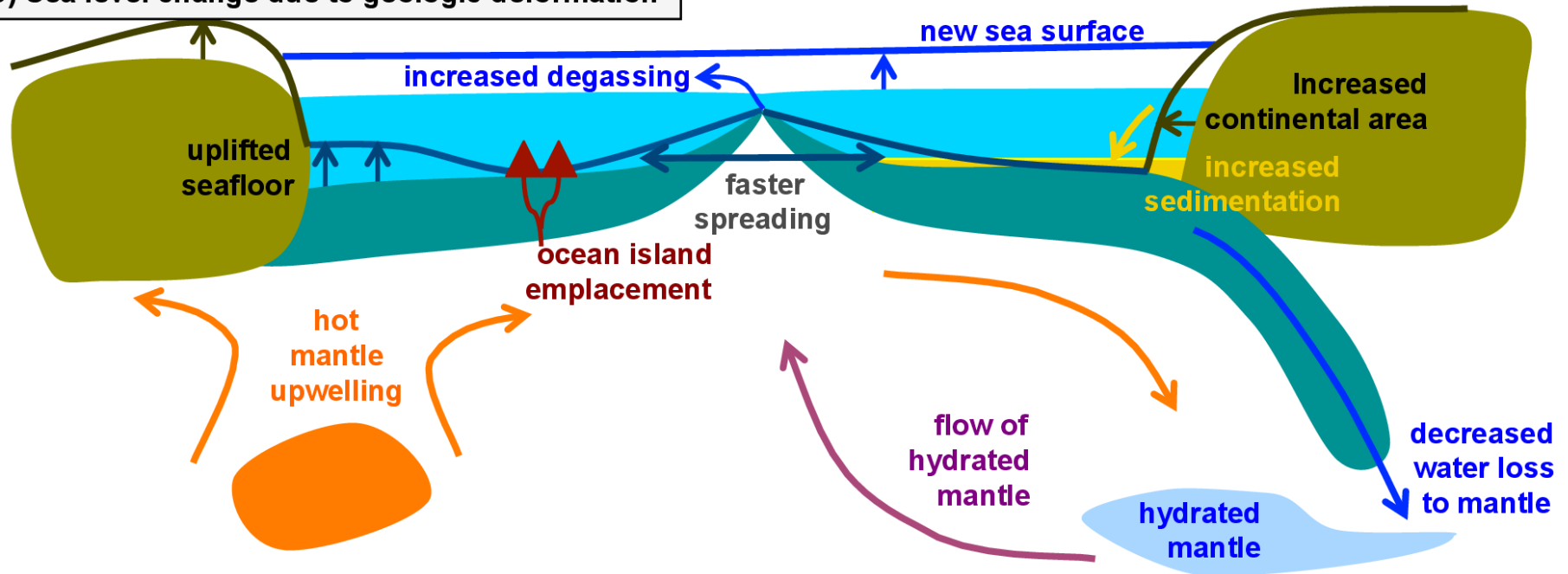
→ Compare to flooding observations on continents.



Global topography at 105 Ma [Blakey]

**Caution: Geologic constraints can be complex!**  
**Example: Many factors affect sea level change**

C) Sea level change due to geologic deformation



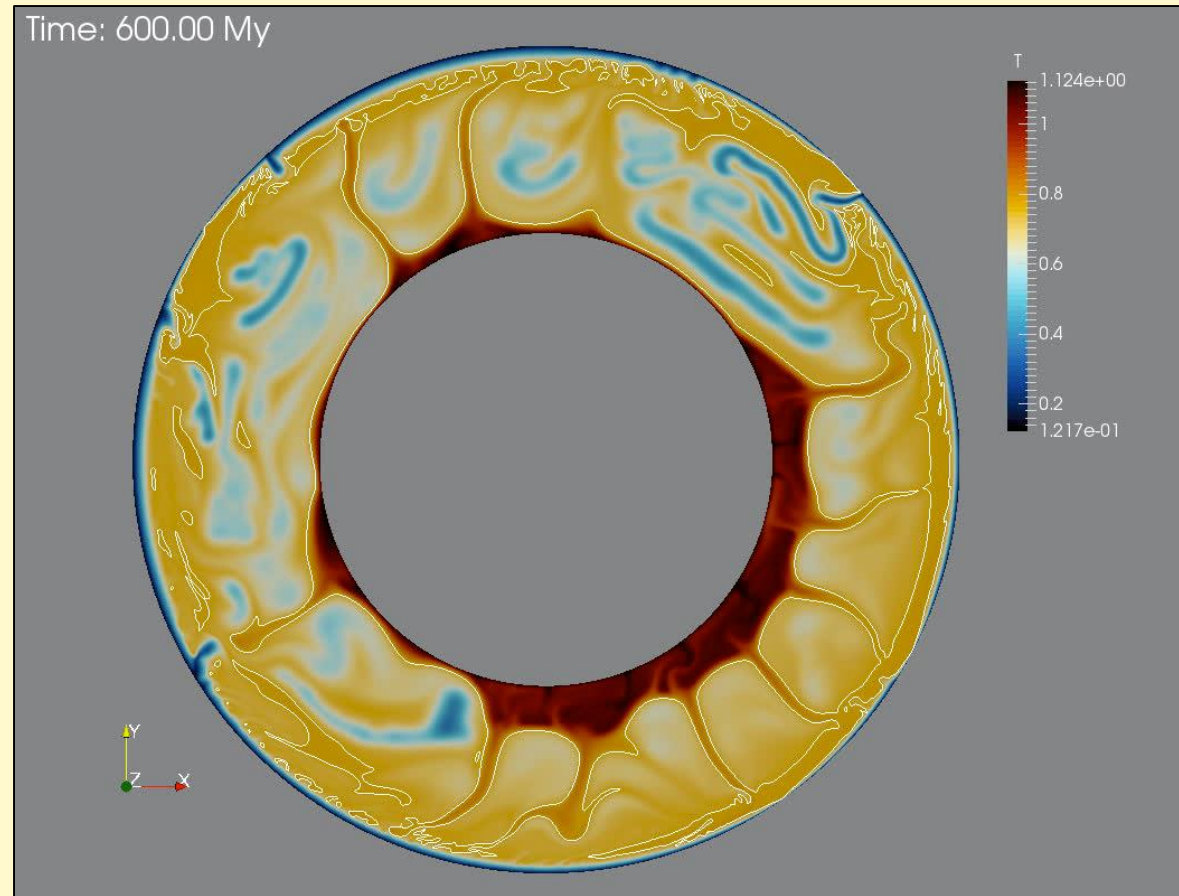
Conrad [2013]

# Conclusions

**Models** can help us understand mantle flow.

**These models can explain:**

- Plate tectonic behavior arising from mantle flow
- Patterns of present-day mantle flow



**The next challenge: Mantle flow for past times**

- How stable is the mantle flow pattern?
- Can we relate past flow to geologic observations?